



# The Universal Cloud and Aerosol Sounding System (UCASS): a low-cost miniature optical particle counter for use in dropsonde or balloon-borne sounding systems.

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**Abstract.** A low-cost miniaturized particle counter has been developed by The University of Hertfordshire (UH) for the measurement of aerosol/droplet concentrations and size distributions. The Universal Cloud and Aerosol Sounding System (UCASS) is an Optical Particle Counter (OPC), which uses wide-angle elastic light scattering for the high precision sizing of fluid-borne particulates. The UCASS has up to 16 configurable size bins, capable of sizing particles in the range 0.4–40  $\mu\text{m}$  diameter. Unlike traditional particle counters, the UCASS is an open-geometry system which relies on an external air flow. Therefore the instrument is suited for use as part of a dropsonde, balloon-borne sounding system, as part of an Unmanned Aerial Vehicle (UAV), or on any measurement platform with a known air flow. Data can be logged autonomously using an on-board SD card, or the device can be interfaced with commercially available meteorological sondes to transmit data in real time. The device has been deployed on various research platforms to take measurements of both droplets and dry aerosol particles. Comparative results with co-located instrumentation in both laboratory and field settings are used to assess the performance of the UCASS.

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## 1 Introduction

Atmospheric aerosols are a key component in the Earth's radiative system as they modify the local and planetary albedo by way of direct and indirect effects. Aerosols directly impact the radiation budget via the scattering and absorption of solar radiation, and to a lesser extent, the scattering, absorption and emission of terrestrial radiation (Li et al., 2010; Zhou and Savijärvi, 2014). At the Top Of Atmosphere (TOA), radiative forcing due to direct aerosol effects is estimated at  $-0.35(-0.85$  to  $+0.15)$   $\text{W m}^{-2}$  (Myhre et al., 2013). Indirect effects arise from aerosols acting as Cloud Condensation Nuclei (CCN) or Ice Nuclei (IN), thus influencing the formation and evolution of clouds (Twomey, 1977). Measurements from satellites observe the Cloud Radiative Effect (CRE) to be  $-50$   $\text{W m}^{-2}$  in the short-wave and  $+30$   $\text{W m}^{-2}$  in the long-wave, with estimates varying by 10 % (Loeb et al., 2009). The significant uncertainties of both direct and indirect effects, combined with the subsequent



rapid adjustments and feedbacks, prompted the 2013 Intergovernmental Panel on Climate Change (IPCC) to cite Aerosol-Radiation Effects (ARE) and Aerosol-Cloud interactions as the largest sources of uncertainty in predicting climate change today (IPCC2013). The uncertainties stem from a range of causes, including the inadequate characterisation of aerosol optical properties, the inaccurate representation of their spatial and temporal coverage, and the complex nature of interactions between cloud and aerosol (Ramanathan et al., 2001; Twomey, 1977; Charlson et al., 2001; Storelvmo, 2012).

To constrain uncertainties, a varied approach to aerosol measurement is required. To cover large geographical scales, there exists a multitude of measurements from ground based sun photometer networks (Holben et al., 1998; Che et al., 2009; Bokoye et al., 2002) to satellite based lidar (MODIS, MISR, POLDER, PARASOL)(Huete, 2004; Chu et al., 2003; Kaufman et al., 1997; Zhang and Christopher, 2003; Jiao et al., 2018). This combination of instruments provide near-continuous measurements world-wide and yield retrievals of aerosol properties such as Aerosol Optical Depth (AOD), number concentration and size distribution. This large-scale coverage is crucial in capturing the spatial extent of atmospheric aerosol, and the high temporal resolution permits the monitoring of long-term diurnal, seasonal and annual trends. However, there exist considerable discrepancies between satellite products due to uncertainties in calibration, assumed aerosol microphysics, sampling and cloud screening (Kokhanovsky et al., 2010).

Many other model studies have demonstrated the sensitivity of ARE to the aerosol layer height (Vuolo et al., 2014; Mishra et al., 2015), thus highlighting the need to faithfully represent the the vertical distribution of aerosol in global models which remains a serious challenge. Therefore, in-situ studies are necessary for obtaining high-quality, comprehensive datasets for the microphysical characterization of aerosol, the testing of retrieval algorithms and the quality assurance of remote sensing data. Vertically resolved in-situ data are typically gathered with aircraft based instrumentation during research campaigns. These campaigns employ a variety of instruments for particle measurement: from single-scattering particle probes for the counting and sizing of small particles; optical array probes for the imaging of larger particles; and filters to collect samples for in-depth chemical analysis. Despite this assemblage of techniques, there remain deficiencies in measurement capabilities. Aircraft campaigns are expensive and subject to accessibility issues, thus limiting the time, space and location each campaign may cover. Whilst co-located remote and in-situ measurement campaigns are invaluable for the validation of retrieval algorithms, the geographical extent of in-situ measurements remains limited. Alternative height-resolved in-situ measurement techniques are required to bridge this gap between the comprehensive, but spatially limited aircraft campaigns, and the expansive remote sensing networks.

This paper discusses a novel instrument, the Universal Cloud and Aerosol Sounding System (UCASS): a low-cost, lightweight, open path Optical Particle Counter (OPC), designed for use as a dropsonde or as part of a balloon-borne sounding system. Routine meteorological soundings are performed daily at a number of research stations worldwide, gathering information about temperature, pressure, humidity and wind. These radiosoundings can be combined with further instrumentation to measure additional data such as ozone distribution and electric field (Jenkins et al., 2015; Nicoll, 2012). However these soundings do not currently take measurements of atmospheric particles. The UCASS offers the ability to incorporate particle measurements into these soundings, giving vertical distributions of aerosol/droplet concentrations and size distributions. The UCASS may be used as both a dropsonde or upsonde and can therefore provide complimentary data to aircraft based campaigns. Furthermore, due

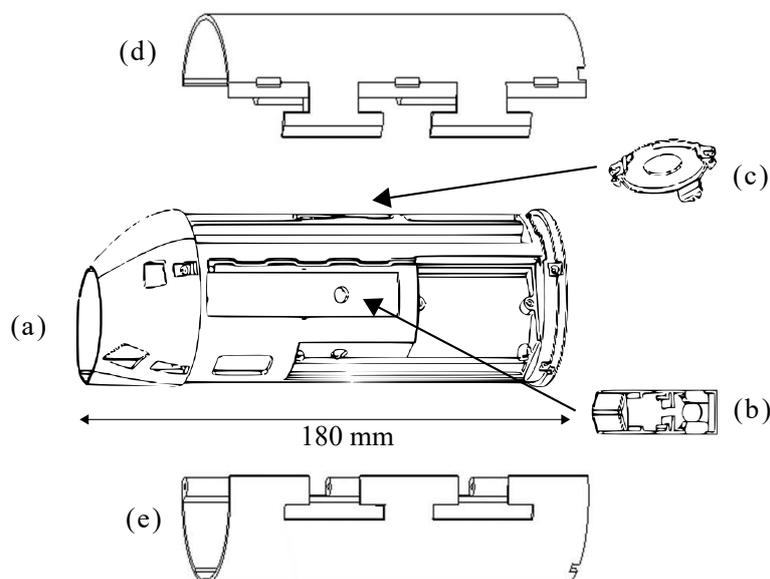


to the relative ease and affordability of radiosoundings, the UCASS also offers an alternative to aircraft based measurements with fewer time and space restrictions.

## 2 Instrument Design

### 2.1 Assembly and optical set-up

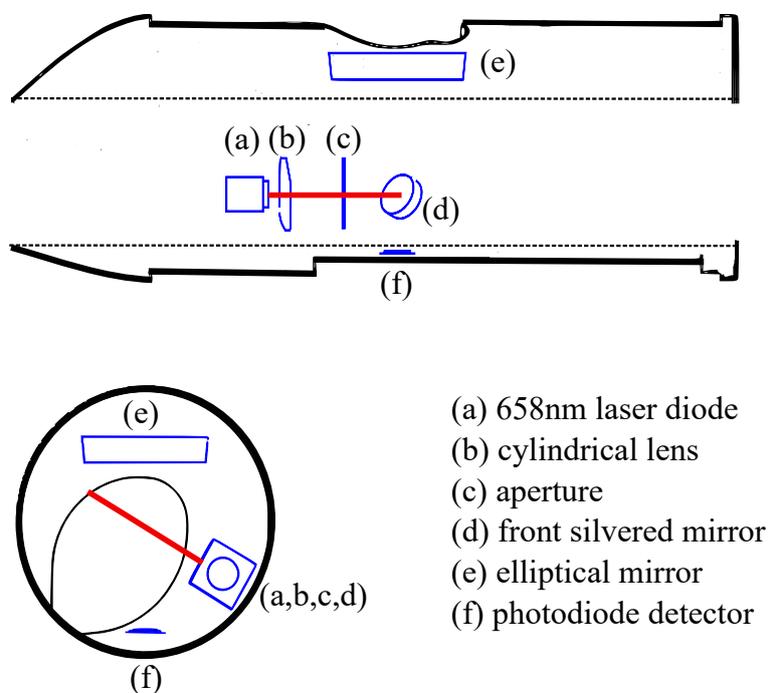
5 The UCASS casing is made primarily from 3D printed parts as shown in Fig. 1. The design is essentially tubular, 180 mm long and 64 mm in diameter. An elliptical cross-section hole, 30 mm × 22 mm in size, runs the full length of the UCASS body and offers a low impedance path for the sampled aerosol. The main unit (a) is 3D printed from nylon using Selective Laser Sintering (SLS). This process is chosen for the main unit as it produces mechanically rigid and thermally stable chassis, which is crucial for the instrument to endure the temperature cycle associated with a routine sounding. The process uses a layer thickness of  
10 0.12 mm, which is adequate for the placement of mechanical components such as batteries and circuitry. However, to improve the precision for optical components, a modular design is used, thus allowing each optical element to be aligned and secured individually. The beam forming optics are mounted onto an insert (b), which is 3D printed from Polyactic Acid (PLA) using Fused Filament Fabrication (FFF). This method is chosen for the increased mechanical rigidity which is required for the thin wall thicknesses. and the elliptical mirror is mounted on aluminium insert (c), which is machined using a 5-axis mill. Once  
15 the beam forming optics are aligned, the insert (b) is secured into place in the main unit. The elliptical mirror, mounted on insert (c) is aligned using three adjustment screws which secure the mirror onto the main unit. One assembled, the UCASS is enclosed by a 3D printed sleeve (d) (e).



**Figure 1.** 3D printed case for the UCASS, consisting of the main housing unit (a), inserts for the beam forming optics (b) and the outer casing (d) (e). The mirror holder (c) is machined out of aluminium.

The optical assembly is shown in Fig. 2. The input beam is a 658 nm continuous-wave diode laser from Oclaro (part number HL6501MG), operating at 10 mW. The laser is fitted with a collimator from the Optoelectronics Company (part number 500-020012). The collimated beam of elliptical cross-section approximately 3.5 mm by 2 mm is focussed along the short axis by a 50mm focal length cylindrical lens (Edmund Optics part number 68-047) and then passes through a 2 mm aperture perpendicular to the beam's long axis. For ease of assembly, these beam-forming components (parts (a), (b), and (c) in Fig. 2) are mounted onto an insert (part (b) in Fig. 1). A plane 9 mm × 9 mm front silvered mirror (Edmund optics part number 31-004) reflects the beam at 45 °, directing it across the flow of air through the UCASS body. The 2 mm aperture has the effect of reforming the beam's Gaussian intensity profile into an approximate 'top hat' profile such that, at the sensing area (coincident with the focal distance of the cylindrical lens), the beam has a relatively uniform intensity cross-section of  $\approx 2$  mm width by  $\approx 40$   $\mu\text{m}$  depth. Particles passing through the beam parallel to its short axis therefore experience similar levels of irradiance, a prerequisite for accurate particle sizing. This is discussed in Sect. 2.2.

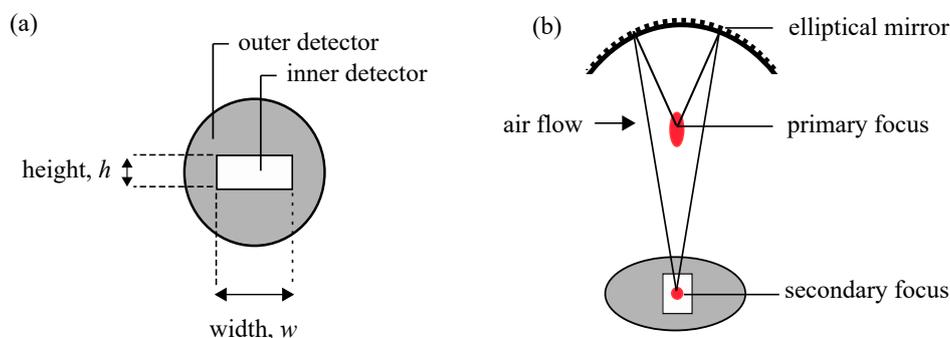
Particles carried in the airflow through the UCASS may pass through the laser at any point in its traverse across the air flow path. However, only those passing through a specific 0.5 mm<sup>2</sup> sensing area within the laser beam are measured. This sensing area is defined optically by a custom designed combination of concave elliptical mirror and a dual-area photodiode detector. Both mirror and detector were designed by the University of Hertfordshire and are now commercially produced by Alphasense Ltd (part number 836-0001-00) and First Sensor GmbH (part number DP12.5-6 SMD) respectively. The mirror collects light scattered by a particle within the sensing area of the laser beam within a solid angle element of 1.69 sr (scattering angles from 16 ° to 104 °) and directs it towards the photodiode detector, where the total intensity is measured.



**Figure 2.** Optical assembly of the UCASS. The top panel shows a cross-sectional view from the side of the instrument, and the bottom panel shows a view from the inlet. The beam forming optics (shown in blue) consist of a laser with collimator (a), a cylindrical lens (b) and a 2 mm aperture (c). The beam is directed into the instrument via a front silvered mirror (d), angled at  $45^\circ$  to the beam. Particles that cross the laser beam will scatter light. The elliptical mirror (e) collects light scattered between angles  $16^\circ$  and  $104^\circ$  and focuses this onto the detector (f), where the pulse height and duration is recorded.

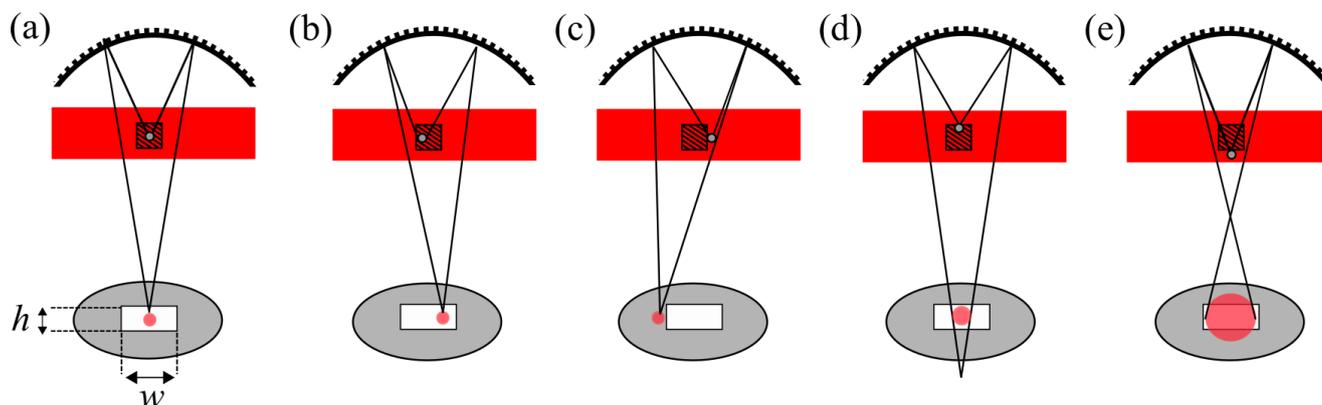
## 2.2 Particle detection and defining the sensing area

The photodiode detector, shown in Fig. 3(a), is a dual area design comprising two electrically isolated photosensitive regions - a central rectangular region measuring  $1.9 \text{ mm} \times 1.0 \text{ mm}$  surrounded by a circular region of diameter 4 mm. The design is such that, particles which are contained centrally within the sensing area will cause scattered light to be focused onto the inner detector only, whereas particles passing immediately outside of the sensing area will cause the scattered light to be focused onto the outer detector - therefore allowing the sample area to be defined optically.



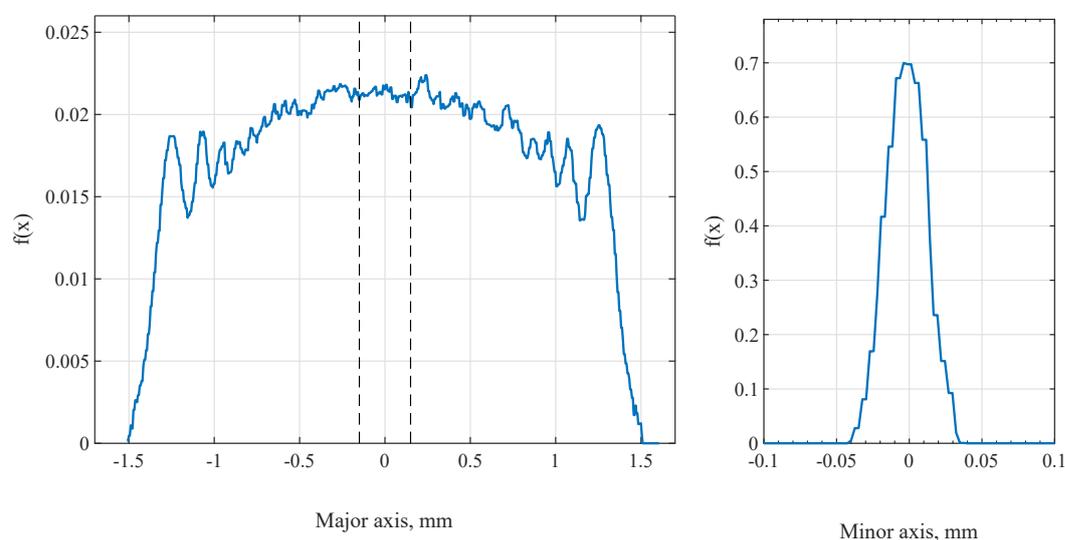
**Figure 3.** (a) The dual element detector consists of an inner rectangular detector of width  $w$  and height  $h$ . (b) A view of the optical system orthogonal to the air flow: the laser beam is centred on the primary focus of the elliptical mirror, and the the detector is centred at the secondary focus. The width,  $w$ , of the inner detector dictates the length of the laser beam (along the beam axis, orthogonal to the air flow) in which a particle can be detected. The height,  $h$ , of the inner detector is wider than the path of a particle crossing the laser beam, and therefore allows the entire pulse caused by a particle transit to be recorded.

The geometry of the detector, mirror and laser are used to define the sensing area. As shown in Fig. 3(b), the focal point of the laser beam is centred at the primary focus of the elliptical mirror, and the detector is centred at the secondary focus. The inner detection area is a rectangle, where the width,  $w$ , directly governs the length of the laser beam (along the beam axis) in which a particle can be detected - and therefore defines one dimension of the sensing area. Figure 4 shows a rotated view of Fig. 3, viewed in the direction of the air flow. Figures 4(a), (b) & (c), show how the width,  $w$ , of the inner detector governs the length of the sensing area in the  $z$  axis. In this direction, particles passing through the primary focus of the elliptical mirror will cause scattered light to be focused at the secondary focus. However, only those passing within  $\pm \frac{1}{2}w$  of the centre of the sensing area will be focused onto the inner detector, whereas particles passing to the left or right will cause light to be focused onto the outer detector (as shown in Fig. 4(c) and will not be counted. Figure 4(d) & (e) shows how the focus of the elliptical mirror is used to define the dimension of the sensing area along the  $y$  axis. Particles passing above or below the primary focus of the mirror will cause scattered light to be focused above or below the detector - creating an enlarged image to form on the detector. In the current arrangement, particles passing 0.15 mm above/below the primary focus will create an image large enough to illuminate both the inner and outer detector, as shown in Fig. 4(e), at which point the particle is considered to be outside of the sensing area. This distance is governed by the geometry of the mirror and the size of the inner detector. For particles straddling the boundary of the sensing volume, light will be focused onto both the inner and outer detectors and therefore a criteria is defined to determine when the particle should or should not be counted. Based upon the intensity measured by the inner detector ( $I_1$ ), and the intensity measured by the outer detector ( $I_2$ ): if  $I_2 \leq \frac{1}{4}I_1$  then the particle is considered to be inside the sensing area.



**Figure 4.** Optically defining the sensing area. The sensing area is shown as a shaded square within the laser beam, particles are represented by grey dots. The airflow direction is into the page. (a) particles passing through the primary focus in the centre of the sensing area cause scattered light to be focussed at the centre of the detector, (b) particles passing to the left/right of the primary focus, but still within the sensing area, will cause scattered light to be focussed to the right/left of the centre of the inner detector, (c) particles passing more than  $\frac{1}{2}w$  to the left/right of the centre of the sensing area will cause scattered light to be focussed onto the outer detector, (d) particles passing above/below the primary focus will cause the scattered light to be focussed below/above the secondary focus and therefore below/above the detector, resulting in a larger image on the detector, (e) particles passing more than 0.15 mm above/below the primary focus will cause scattered light to fall on both the inner and outer detector; the ratio of the inner and outer detector signals is used to determine whether particles are considered to be in the sensing area.

Particles passing through the sensing area will pass through the short axis of the focussed laser beam, causing a pulse of light to fall on the detector. The width of this pulse is equal to the Time-of-Flight (ToF) of the particle across the beam, and the height of this pulse will correspond to the maximum intensity incident upon the detector during this transition. The pulse height depends upon the amount of light scattered by the particle: a function of particle size, optical properties and the intensity of the incident light. Therefore, the pulse height is used to derive particle size, whilst ToF data is stored for data quality assurance. To ensure that the laser beam intensity is consistent across the sample area, the beam is shaped with the use of a 2 mm aperture, thus minimising sizing errors due to beam non-uniformity. Figure 5 shows the relative intensity distributions across the major (left) and minor (right) axes. As discussed above, particles passing at the edges of the major axis will form unfocussed images on the detector and therefore will not be counted, only the central 0.3 mm section of the major axis falls within the sensing area.



**Figure 5.** Probability density functions  $f(x)$  of the laser intensity across the major (left) and minor (right) axes of the focussed laser beam. Only particles passing within the central 0.3 mm of the major axis (shown by dashed lines) will be considered within the sensing area, and thus only these will be counted.

### 2.3 Electronics

A PIC microcontroller (PIC18F27J53) based electronic circuit governs the operation of the UCASS and how measurements using the dual-photodiode system are made and recorded. Light scattered from a particle well centred in the sensing area will be focussed predominantly on the central element of the dual-photodiode. This element is connected to a circuit that converts the photo current to a voltage (0–3.3 VDC) and activates a trigger if the voltage amplitude detected goes above a pre-defined threshold ( $\approx 5$  mV) above the background signal. In this event, the trigger electronics notifies the core microcontroller, which in turn activates peak-detector electronics that will sample and hold the peak of the light signal on both photo-sensing



elements as the particle passes through the laser beam. The microcontroller then digitises both signals with an on-board 12bit Analog-to-Digital Converter (ADC) and categorises the combined peak signal into one of 4095 bins dependent upon the value of the amplitude. This instrument response is known as Amplitude Displacement (AD). All 4095 available bins are used for calibration purposes (discussed in section 2.4.2). However during routine measurements, the data are compressed into fewer size bins, with thresholds defined by the user. Measurements are rejected if the particle is deemed off centre (as discussed in Sect. 2.2), measurements are also rejected if the ToF is deemed too short or too long. A short ToF would occur if an event such as electrical noise spikes triggered the system. A long ToF would be the case if a large body or agglomeration of particles passed through the beam. The default ToF minimum/maximum limits are 1 and 100  $\mu\text{s}$ , respectively.

In this manner, the microcontroller continuously assembles histograms of particle size/count data. Means of ToF data are also recorded for a subset of the size bins to allow some use of this parameter to monitor flow rate through the device. At every one-second interval, the histogram data-set is either saved to an on-board micro SD card or transmitted via a serial link (XDATA protocol) to a radiosonde device for Radio Frequency (RF) transmission of the data. The histogram data are cleared following each data save/transmit event. The XDATA protocol is limiting in terms of bandwidth such that there is only space to utilise 10 of the available 16 size bins. The configuration of the bin boundaries can be altered to accommodate this and make best use of the bin quantity available. The device can measure up to  $10^4$  particles per second and can operate in air flow speeds between 2 and  $15 \text{ ms}^{-1}$  with the standard firmware. The firmware can be modified to change the ToF window, thus allowing shorter or longer particle transitions to be recognised if the measurements platform requires it (this is discussed further in Sect. 3.2).

The size histogram bins utilised by the microcontroller are essentially levels of peak amplitude light signals scattered from measured particles. The bin boundaries are interpreted into particle diameters based on calibration using both theoretical and experimental data.

## 2.4 Calibration

### 2.4.1 Theoretical instrument response

An OPC directly measures the scattering cross-section of a particle. The scattering cross-section is a function of the particle properties (size, shape, refractive index), the collection angle of the instrument, and the wavelength of the beam (Bohren and Huffman, 1998; Rosenberg et al., 2012). This is defined as:

$$\sigma_{sca} = \frac{1}{k^2} \int_0^{2\pi} \int_0^{\pi} \frac{1}{2} [ |S_1(\theta, kD_p, n)|^2 + |S_2(\theta, kD_p, n)|^2 ] \sin \theta \omega(\theta, \phi) d\theta d\phi \quad (1)$$

30 where:



$\sigma_{sca}$  = scattering cross-section,  $m^2$

$k$  = wavenumber in medium,  $m^{-1}$

$S_1$  = amplitude scattering matrix for parallel polarised light

$S_2$  = amplitude scattering matrix representing perpendicularly polarised light

5  $\theta$  = scattering angle measured from the incident beam direction, rad

$\phi$  = angle of the scattered light azimuthally around the incident beam, rad

$D_p$  = diameter of the scattering particle, m

$n$  = refractive index of the particle

$\omega$  = weighting function based on the mirror geometry

10

The weighting function can have values between 0 and 1 and describes the azimuthal extent of the collection optics in a finite element  $\theta \rightarrow d\theta$  (Rosenberg et al., 2012). Optical particle counters often use annular reflectors, such as those in the Cloud Droplet Probe (CDP) (Lance et al., 2010) and the Passive Cavity Aerosol Spectrometer Probe (PCASP) (Cai et al., 2013), both by Droplet Measurement Technologies. The weighting function for these instruments can have values of 0 and 1 only. The

15 UCASS utilises a none-annular reflector in the form of an elliptical mirror centred on a scattering angle of  $\theta = 60^\circ$  with a half angle of  $43.8^\circ$ . The mirror also has a central hole with half angle  $10.7^\circ$ . Using this geometry, the weighting function,  $\omega(\theta, \phi)$ , is given by:

$$\omega(\theta, \phi) = \frac{1}{\pi}(\phi_m(\theta) - \phi_h(\theta)) \quad (2)$$

where  $\phi_m$  and  $\phi_h$  express the angular extent of the mirror and the hole, respectively, azimuthally around the laser beam as a

20 function of scattering angle,  $\theta$ . For a primary reflector not centred on  $\theta = 0^\circ$ ,  $\phi$  is given by:

$$\phi_{m,h} = \begin{cases} \cos^{-1}[(\cos H_{m,h} - \cos L_{m,h} \cos \theta)/(\sin L_{m,h} \sin \theta)] & \text{if } -1 < (\cos H_{m,h} - \cos L_{m,h} \cos \theta)/(\sin L_{m,h} \sin \theta) < 1 \\ 0 & \text{if } (\cos H_{m,h} - \cos L_{m,h} \cos \theta)/(\sin L_{m,h} \sin \theta) \geq 1 \\ \pi & \text{if } (\cos H_{m,h} - \cos L_{m,h} \cos \theta)/(\sin L_{m,h} \sin \theta) \leq -1 \end{cases} \quad (3)$$

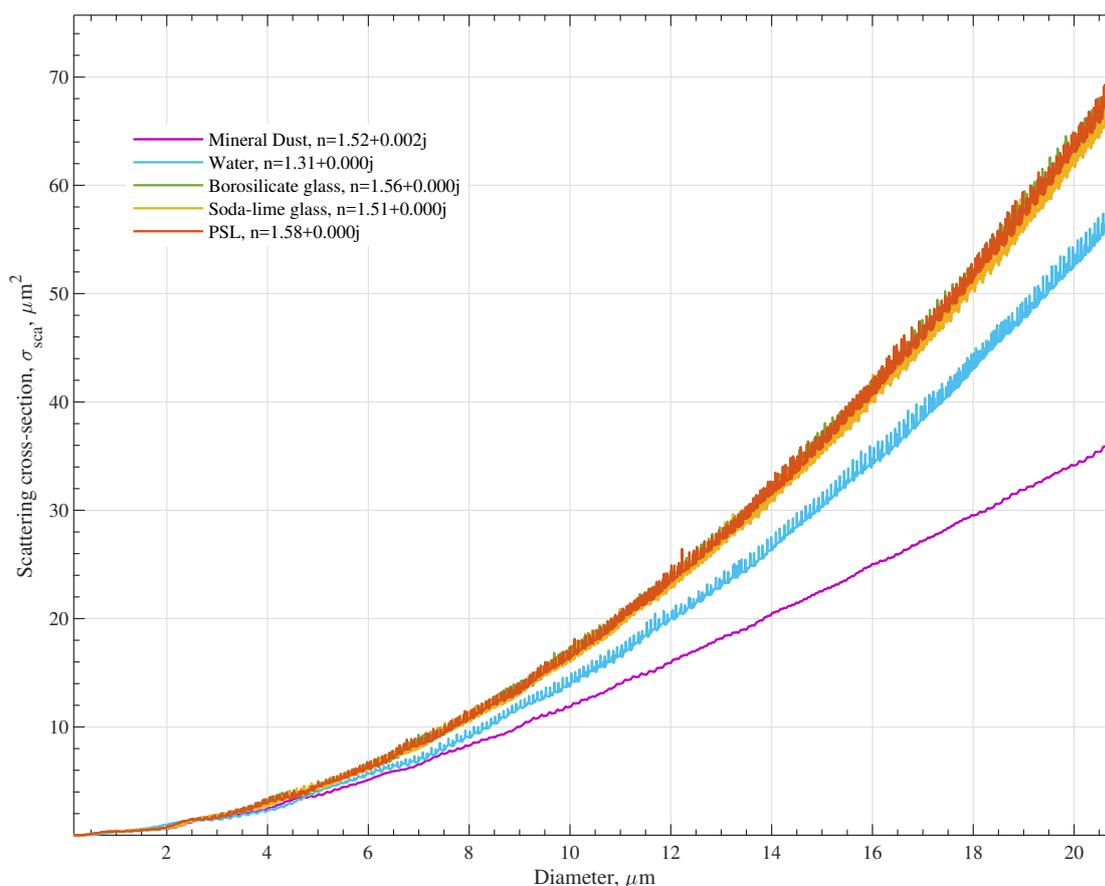
where:

25  $L_{m,h}$  = the lens angle of the mirror and hole, respectively

$H_{m,h}$  = the half angle of the mirror and hole, respectively



Using this weighting function in Eq. 1, we obtain the scattering cross-sections as a function of diameter. The matrices  $S1$  and  $S2$  are computed using Jan Schäfer's mie scattering code 'Matscat' (Schäfer, 2011; Schäfer et al., 2012; Bohren and Huffman, 1998; Lee, 1990; Kerker, 1969). Figure 6 shows the scattering cross-sections for the materials used in calibration: PolyStyrene-Latex (PSL), borosilicate glass and soda-lime glass, and two typical materials measured with the UCASS: water and mineral dust. Typically for OPCs, the relationship between scattering cross-section and geometric diameter is highly none-monotonic due to the presence of mie oscillations. These mie oscillations are most prevalent in the forward scattering region, and so the wide collection angle of the UCASS elliptical mirror results in a near-monotonic relationship, as evidenced in Fig. 6. For each calibration standard used, the sizing error associated with mie oscillations falls well within the manufacturer stated standard deviation of the particle size, and so uncertainties associated with mie oscillations are not further considered.



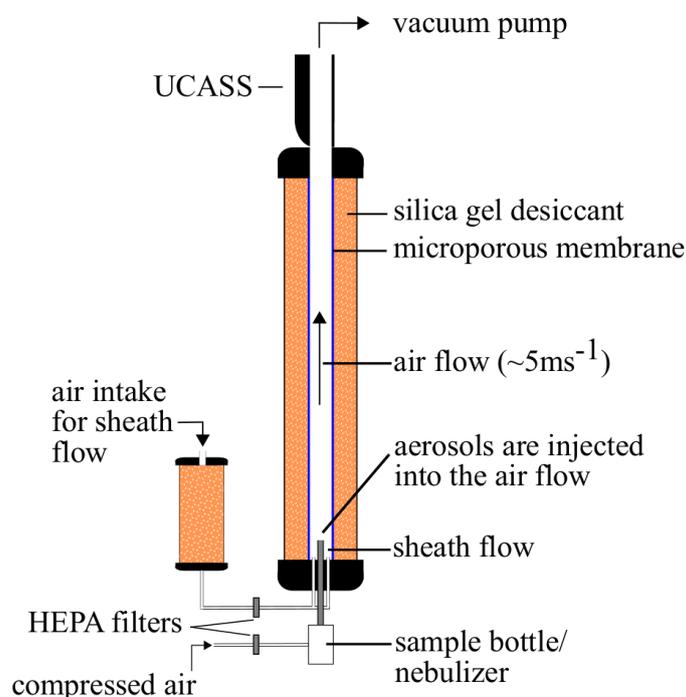
**Figure 6.** Scattering cross-sections,  $\sigma_{sca}$ , as a function of geometric diameter for mineral dust, water, borosilicate glass, soda-lime glass and PSL.



The scattering cross-section directly describes the amount of light collected by the mirror in the instrument. The instrument response is then a function of scattering cross-section, amplifier gain, detector sensitivity and laser power. Variations in the measured output between different units can occur due to manufacturer tolerances in detector sensitivity or laser power, thus causing offsets from the theoretical instrument response. Therefore, to constrain the calibration curve, we combine measured  
5 instrument responses with the theoretical results from mie theory.

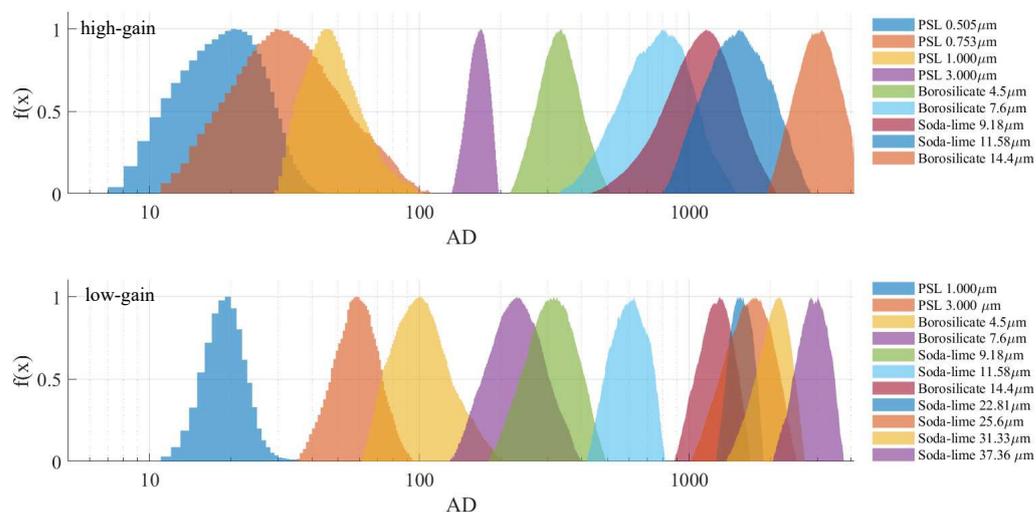
#### 2.4.2 Calibration Measurements

To constrain the calibration curve, we take measurements of the instrument response for a series of NIST-traceable particle standards. These include: PSL from Polysciences; borosilicate glass from Duke scientific; and soda-lime glass from Whitehouse Scientific. A drying column is used to ensure the proper drying and dispersion of the calibration standards, whilst also  
10 establishing a stable, none-turbulent air flow through the UCASS. The experimental set-up is shown in Fig. 7. The air intake consists of an initial drying chamber filled with silica gel desiccant, which dries the ambient air before passing it through a HEPA filter. This air is then directed through a series of pipes at the bottom of the drying column, which form a concentric ring around the aerosol input, thus establishing a clean, dry sheath flow for the aerosol to be injected into. The main drying column  
15 consists of a 1 m tall tube with an inner tube for the airflow. The outer tube is packed with silica gel desiccant and the inner tube is lined with a microporous membrane. This combination ensures that any aerosol injected into the sheath flow will be fully dried before reaching the UCASS at the top of the column. A vacuum pump is used in line with the UCASS to establish an air flow of 3–5 ms<sup>-1</sup> (the intended operational velocity of the UCASS). Two different aerosol input techniques are used for wet and dry dispersion. The PSL microspheres are suspended in an aqueous solution and are aerosolised using a TSI Tri-Jet 3460  
20 which is fed directly into the sheath flow. To disperse the dry calibration standards, a small amount of beads are placed into a small dispersion bottle consisting of an inlet and outlet pipe. A burst of dried, filtered air in the inlet pipe creates a turbulent environment inside the bottle which aids the breakup of any agglomerates. The outlet pipe is then directed into the sheath flow and the dispersed particles are carried in the airflow through the UCASS.



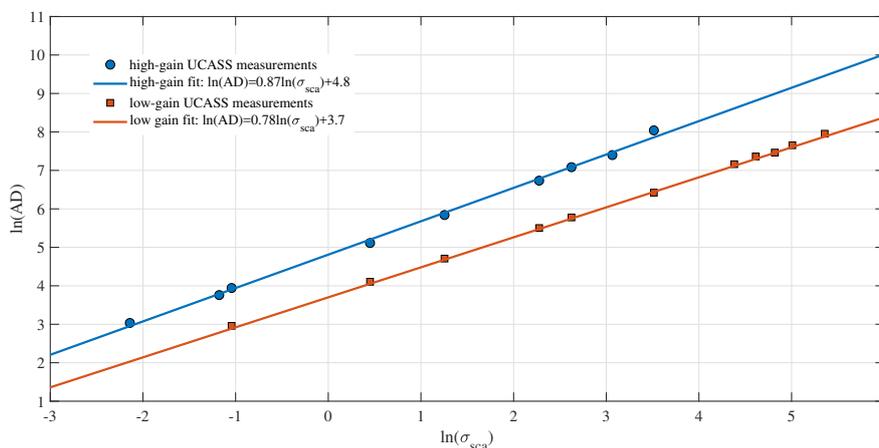
**Figure 7.** Experimental set up used to take calibration measurements. A vacuum pump is used to draw air through the drying column. The air intake is taken from ambient air which is first passed through desiccant and then a HEPA filter. This clean dry air is then directed through a concentric ring of pipes at the bottom of the drying column to create a clean, dry sheath flow. The aerosol is then injected into the sheath flow via the TSI tri-jet aerosol generator (for wet dispersion), or via a sample bottle for dry particles. The sheath flow constrains the aerosol in the centre of the air flow, which is carried directly through the sensing area of the UCASS.

Although PSL and glass beads are typically considered monodisperse for calibration purposes, the full size range can be broad. Therefore the full size distribution of the aerosol is measured, and the mean, median and modal value of the whole distribution are found. Manufacturers may describe the average size of their test particles using different averages, and therefore it is necessary to select the correct average for each sample type. For these calibrations, the mean values are used for PSL and soda-lime glass, whereas the median is used for borosilicate glass. In cases where the full distribution is not captured, the modal value is used. To avoid sizing uncertainties associated with bin-width during calibration, a calibration-specific firmware is installed which allows Particle-by-Particle (PbP) pulse heights to be recorded. Measurements of various calibration standards are shown in Fig. 8. The top panel shows measurements from the ‘high-gain’ version of the UCASS, which has a nominal sizing range of 0.4–17  $\mu\text{m}$ , and the bottom panel shows measurements from the ‘low-gain’ version of the UCASS which has a nominal sizing range of 1–40  $\mu\text{m}$ .



**Figure 8.** Measured instrument responses for various NIST-traceable calibration standards including PSL, Borosilicate glass, and soda-lime glass. For presentation purposes, each distribution is normalised to a peak height of one. The top panel shows measurements conducted using a ‘high-gain’ version of the UCASS, and the bottom panel shows measurements conducted using a ‘low-gain’ version.

Theoretically, the instrument response is directly proportional to the scattering cross-section. However, the instrument response may differ to the theoretical response due to manufacturer tolerances in the optical and electronic components. Some variations such as laser power and amplifier gain will cause constant offsets over the measurement range. However, some offsets may be non-linear due to the varying Signal to Noise Ratio (SNR) across the measurement range and non-linearity in the detector sensitivity. To account for all offsets, the log of instrument response (Amplitude Displacement,  $AD$ ) is plotted against the log of the scattering cross-section, and a line of best fit is applied as shown in Fig. 9.



**Figure 9.** The log of the measured instrument response ( $AD$ ) is plotted against the log of the scattering cross-section ( $\sigma_{sca}$ ) for the calibration standards used. For PSL and Duke Scientific borosilicate glass beads, the mean instrument response is used as this is how the manufacturer defines the particle size. For the Whitehouse Scientific soda-lime glass beads, the sample size is defined by the median and so the median instrument response is plotted.

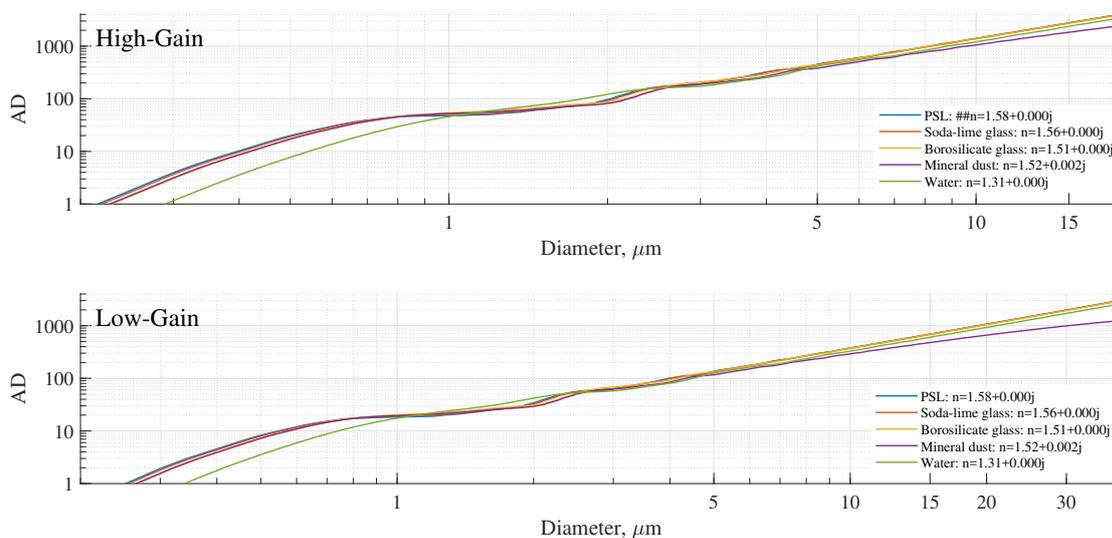
A linear fit is applied to link the scattering cross-section to the instrument response. For the examples shown, the relationship for the high-gain UCASS is given by:

$$\ln(AD) = 0.87\ln(\sigma_{sca}) + 4.8 \quad (4)$$

and the relationship for the low-gain UCASS is given by:

$$5 \quad \ln(AD) = 0.78\ln(\sigma_{sca}) + 3.7 \quad (5)$$

These equations can then be applied to any particulate (e.g. mineral dust, water) by computing the scattering cross-sections for discretized diameters across the measurement range and relating the instrument response to the particle diameter. Using this technique, finalised calibration curves are produced for five materials: PSL, soda-lime glass, borosilicate glass, water and mineral dust, as shown in Fig. 10. If a-priori knowledge of the aerosol type is known, then the calibration curve can be chosen to best suit the application. The calibration curve can then be used to select the boundaries of up to 16 size bins (this is limited to 10 if interfacing the UCASS via XDATA protocol). Since the relationship between  $\ln(AD)$  and  $\ln(\sigma_{sca})$  remains valid for any material/shape of aerosol, if post-factum information on the particle shape or refractive index becomes available (i.e. via analysis of co-located measurements), then a post calibration can be applied.

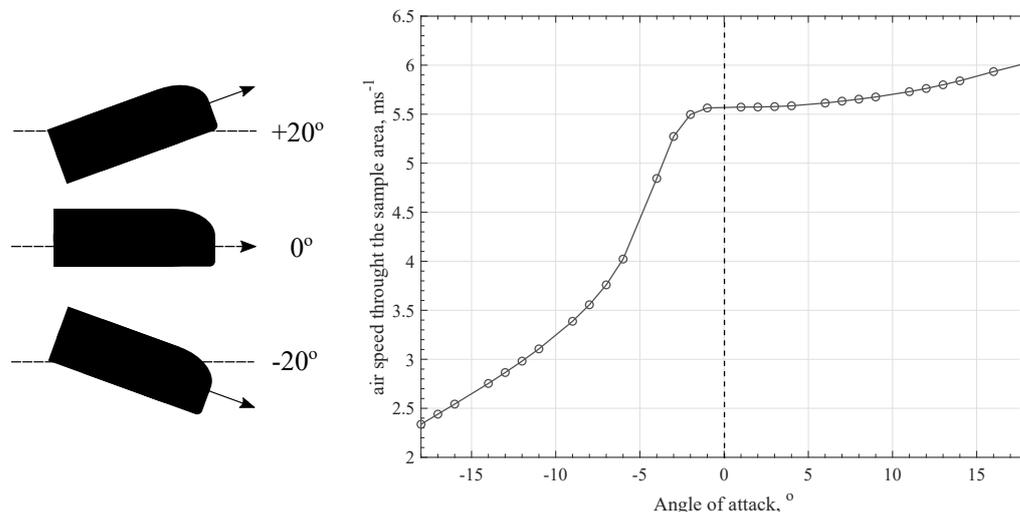


**Figure 10.** Final calibration curves for a high-gain (top) and low-gain (bottom) UCASS. The Geometric diameter is plotted against instrument response for varying materials. If a-priori knowledge of the refractive index is known, the calibration curve can be selected to best suit the material being measured. If post-factum information about the material becomes available, a post calibration can be applied.

## 2.5 Air flow Management

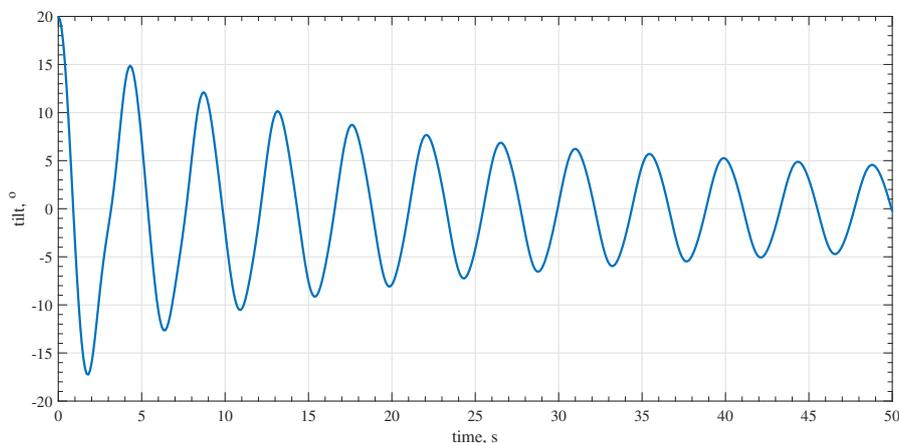
The UCASS is a naturally-aspirated system, by which an external air flow (relative to the UCASS) is required to transport particles through the sensing area. The main advantage of this setup is the reduction of particle loss mechanisms induced by complex aerodynamic systems and non-isokinetic sampling, which has been evaluated empirically for artificially-aspirated systems (Spiegel et al., 2012; Hangal and Willeke, 1990; Von Der Weiden et al., 2009). However, the sampling efficiency is still dependent on the axial characteristics of the external airflow, hence a range of accepted (aerodynamic) operating conditions must be defined. To understand the airflow through the sample area, Computational Fluid Dynamics (CFD) using the ‘Star CCM+’ commercial code is used to model the air flow through the instrument at varying angles of attack. The CFD simulation is a 2-dimensional model on the symmetry plane of the UCASS, this was chosen because the velocity of an ‘air parcel’ in a subsequent plane to this is likely to be similar, meaning the viscous stress between planes is negligible. The fundamental equations behind this CFD are the Reynolds-Averaged Navier-Stokes (RANS) equations, which were chosen for computational efficiency when compared to direct numerical simulation.

Figure 11 shows the air flow through the UCASS for an external velocity of  $5 \text{ ms}^{-1}$ . For an axial air flow, the air velocity through the sample area is  $5.6 \text{ ms}^{-1}$ , 12 % higher than the external air flow. As the UCASS is tilted upwards, the air flow through the sample area increases, however as it is tilted downwards, the air speed decreases. This relationship between tilt and air speed is asymmetric, with negative angles of attack having a greater impact on the air flow. Therefore, to retrieve reliable concentration measurements from the UCASS, the airflow or angle of attack must be known.



**Figure 11.** Model results showing the air velocity in the sensing area. The left panel shows how the tilt direction is described in relation to the instrument geometry. The right hand panel shows the simulated flow velocity in the sample for tilt angles between  $-20^\circ$  and  $20^\circ$  in an external air flow of  $5 \text{ ms}^{-1}$ .

When the UCASS is launched on a meteorological balloon, the balloon-payload configuration acts as a pendulum system. To constrict the movement of the UCASS, the payload is configured as a double pendulum, whereby the UCASS is secured by a line below the balloon, and the meteorological sonde is secured below the UCASS. Although the payload still oscillates, the use of a double pendulum system inhibits the movement of the UCASS, and model results show that the tilt is constrained to within  $\pm 5^\circ$  as shown in Fig. 12.



**Figure 12.** Model results showing the angular tilt of the UCASS with respect to the direction of ascent. Initial oscillations can reach  $\pm 20^\circ$  although these are damped after  $\approx 30 \text{ s}$ , and constrained within  $\pm 5^\circ$ .



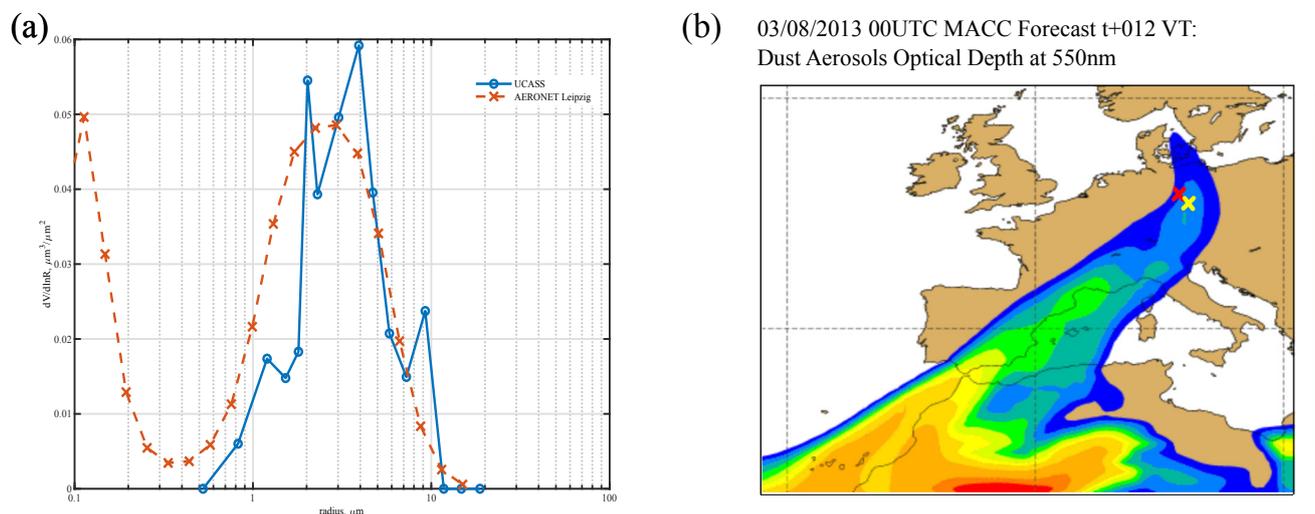
By constraining the tilt of the UCASS within these limits, model results show that the air speed through the sample area is equal to  $5.4 \pm 0.3 \text{ ms}^{-1}$  which is used to compute the number concentrations. For dropsonde systems, a similar configuration is used, whereby the sonde is tethered between the parachute and the meteorological sonde, and the angle of tilt is constrained to within the same boundaries.

## 5 3 Results

### 3.1 In-situ inter-comparisons

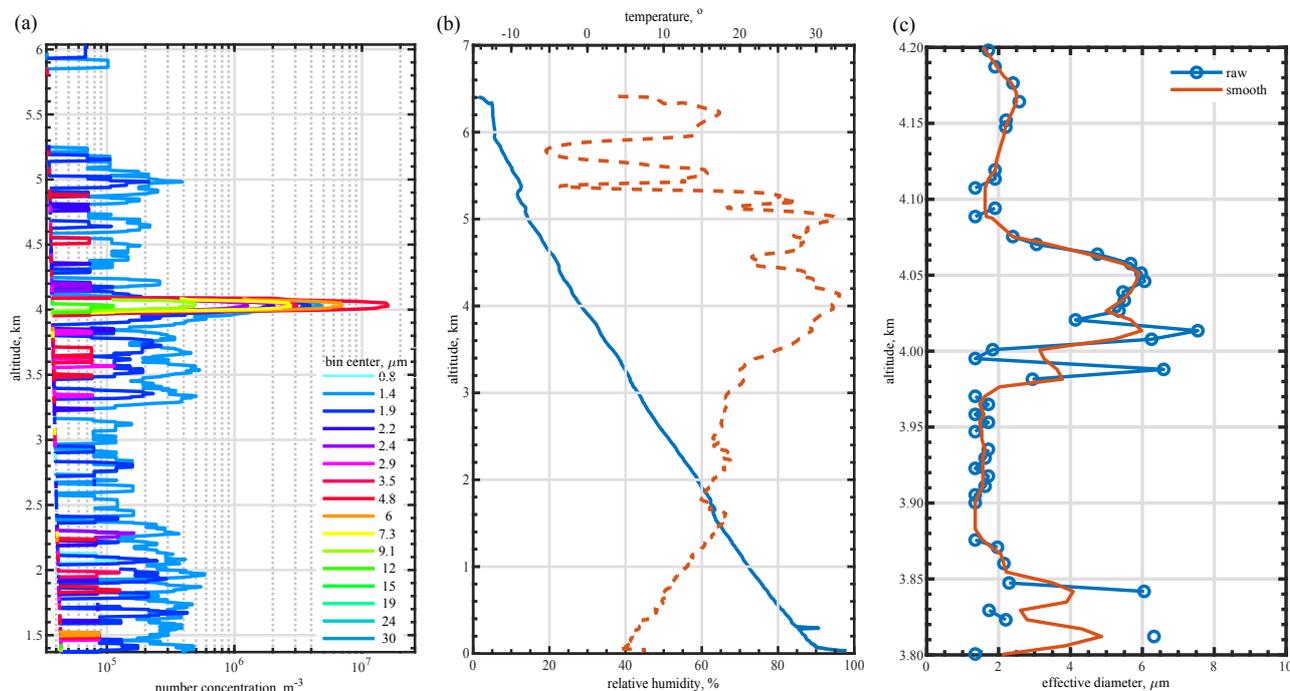
#### 3.1.1 Dropsonde system

Seven dropsondes were launched from the Dornier 128 aircraft north of Magdeburg, Germany, on the 1st and 3rd of August 2013. The packages contained the UCASS in tandem with KITsondes (Wieser et al., 2014), for co-located particle and meteorological data. The size distribution profiles measured by the UCASS indicated mineral dust present in the free troposphere during the 3<sup>rd</sup> August for a drop conducted at 1310UTC. At 1534UTC, AERONET sun photometry retrievals (Dubovik and King, 2000) from IfT Leipzig (located 100 km South-South-East of Magdeburg) confirmed the presence of dust. The integrated size distribution measured by the UCASS is compared to the AERONET inversion from Leipzig in Fig. 13(a). The size distribution is integrated for the column between 3 and 5 km, to exclude boundary layer aerosol over the Colbitz-Letzlinger Heide area (the EDR74 exclusion zone). Figure 13(b) shows the dust dispersion model for 1200UTC using the MACC (Monitoring Atmospheric Composition & Climate) forecast, showing the transport of dust over Germany from the Sahara. The locations of Magdeburg and Leipzig are marked by red and yellow crosses, respectively.



**Figure 13.** (a) Integrated aerosol size distribution measured by the UCASS compared with the AERONET inversion (as retrieved from the sun photometer data) from Leipzig. (b) Dust dispersion model showing Saharan dust transport over Leipzig (marked by a yellow cross) and Magdeburg (marked by a red cross).

In a further drop during this dust event, a further UCASS dropsonde system measured a thin, embedded cloud layer as shown by high number concentrations at larger sizes at 4 km. This is illustrated in Fig. 14(a), where the number concentrations for each of the 16 size bins are shown. Over the 6 km profile, the majority of particulates counted measured less than 3  $\mu\text{m}$  in diameter, except for a thin layer at 4 km with high concentrations of particles in larger size bins. This change in concentration and size suggested the presence of a thin cloud layer, which was confirmed by the humidity profile measured by the attached KITsonde shown in Fig. 14(b), which measured high humidities at 4 km. Figure 14(c) shows the effective diameter with respect to altitude, it can be seen that the saharan dust layers above/below the cloud have an effective diameter of  $\approx 2 \mu\text{m}$ . At 4 km when the UCASS enters the cloud layer, this effective diameter steeply increases to  $\approx 6 \mu\text{m}$ . As discussed in Sect. 1, dust can cause the formation of clouds by acting as a cloud condensation nuclei, and so embedded clouds during dust events are not uncommon. However, the presence of cloud presents difficulties for remote sensing instrumentation, and therefore retrievals are not possible for cloud-contaminated datasets. So long as the UCASS is launched in tandem with a meteorological sonde, the measured humidity profiles can be used to differentiate between cloudy and none cloudy conditions, and therefore the UCASS can be used to measure both cloud and aerosol size distributions even when both are present.



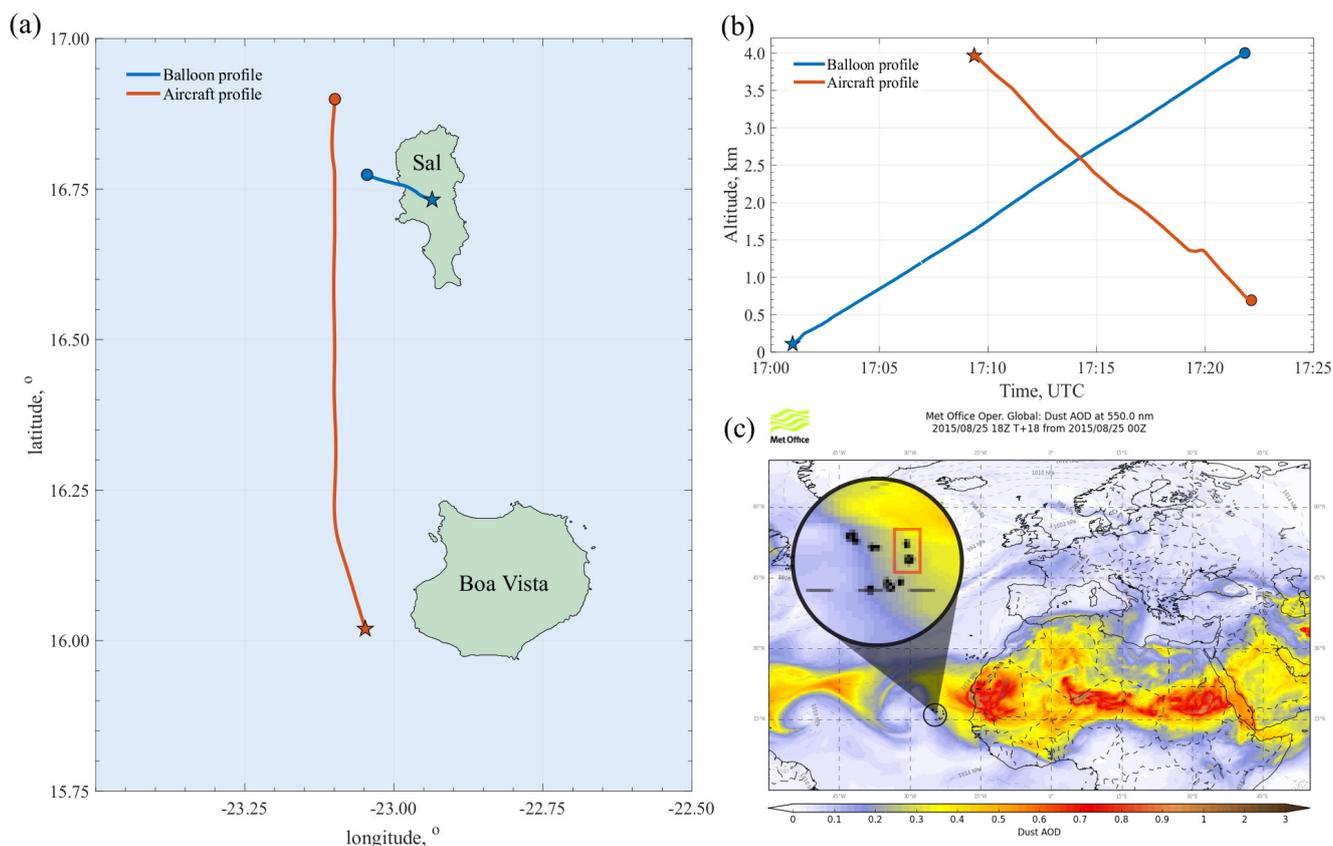
**Figure 14.** (a) Number concentration of particulates in the 16 UCASS size bins with respect to altitude, suggesting an embedded cloud layer at 4 km. (b) Temperature and relative humidity profiles from the attached KITsonde, also suggesting an embedded cloud layer at 4 km with a spike in humidity. (c) Effective diameter with respect to altitude, the embedded cloud layer at 4 km is evident due to the steep increase in effective diameter. The blue trace shows raw data, whilst the orange trace shows smoothed data using a 10 point moving average.

### 3.1.2 Upsonde system

The UCASS was used during the AERosol properties – Dust (AER-D) & Sunphotometer Airborne Validation EXperiment - Dust (SAVEX-D) campaigns over Cape Verde during August 2015 (Marenco et al., 2018). A UCASS was launched as part of a balloon-based system in tandem with a GRAW DFM-09 during a Saharan dust event on the 25<sup>th</sup> August, from Instituto Nacional De Meteorologia E Geofisica (INMG), Espargos, Sal island. The launch was conducted at 1701UTC from ground level (70 m above sea level), this launch was timed to coincide with a deep profile (profile 6) conducted by SAVEX-D flight number B934. The research aircraft started the profile from an altitude of 4 km at 1710UTC off the West coast of Boa Vista island ( $\approx 0.75^\circ\text{S}$ ,  $0.15^\circ\text{W}$  of the balloon launch site). The aircraft then flew due North, passing Sal island until reaching an altitude of 690 m. Figure 15(a) shows the flight paths of the aircraft and balloon (up to 4 km), whilst Fig. 15(b) shows the altitude of both profiles with respect to time. Figure 15(c) shows the near real time dust forecast from the Met Office Global Atmosphere model (Martin et al., 2018) for 1800UTC, where it can be seen that dust is advected westwards from the Sahara. The circular section shows a zoomed in image over the Cape Verde islands, and the islands of Sal and Boa Vista are highlighted by a red rectangle. It can be see from the forecast that the aircraft and balloon profiles were conducted at the leading



edge of a dust layer and therefore, even though the spatial and temporal differences between the two profiles are minimal, some variability in the dust layer may be expected.

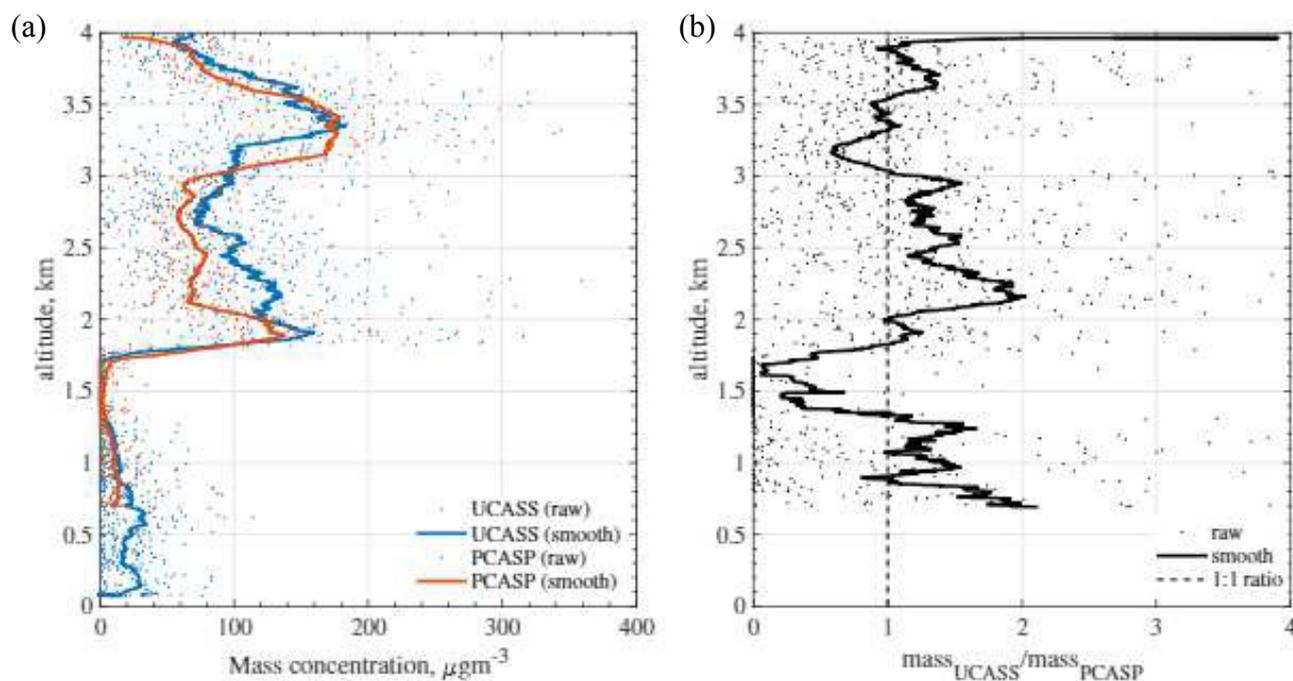


**Figure 15.** A balloon launch was conducted to coincide with profile 6 of SAVEX-D flight B934. (a) shows the flights paths of the balloon (blue) and aircraft (orange) during this profile, where the start and end points are marked by stars and circles, respectively. The altitudes are shown in (b), whereby the aircraft descends from an altitude of 4 km to 690 m. For comparative purposes, only the balloon data up to an altitude of 4 km are shown here. (c) shows the MET office global atmosphere model Aerosol Optical Depths (AODs) for 1800UTC. The cape verde islands are zoomed in and the islands of Sal and Boa Vista are highlighted in a red rectangle. It can be seen from this that the measurements were located to the edge of a dust layer, with anticipated AOD of  $\approx 0.3$ .

Figure 16(a) shows the mass concentration profile measured by the UCASS, compared with measurements from the aircraft mounted PCASP. For comparative purposes, the mass concentrations shown only apply to the overlapping measurable size range of the UCASS and PCASP, in this case the size range is from 0.4–3  $\mu\text{m}$ , corresponding to the first 7 bins of the UCASS. Figure 16(b) shows the ratio of the mass concentrations measured by the UCASS and PCASP, respectively. Over the 4 km profile, the average ratio is 1.2, with the greatest differences between the two measurements occur at the base and top of the dust layer ( $\approx 1.6$  and 4 km respectively). The dust layer was advected from the East as shown in Fig. 15, thus the spatial



separation between the two profiles explains the discrepancies. The more easterly UCASS measurement captures higher concentrations and a higher layer height, whilst the more westerly PCASP measurement measures a lower layer height, likely due to gravitational settling.



**Figure 16.** (a) shows the measured mass concentrations with respect to altitude for the balloon-borne UCASS (blue) and aircraft mounted PCASP (orange). The raw data for each instrument is shown via individual data points, and the smoothed data is shown by a solid line, using a 10 point moving average). The mass concentrations shown pertain to particles within the 0.4–3  $\mu\text{m}$  size range for comparative purposes. (b) shows the ratio of the measured UCASS mass concentration to the PCASP measured mass concentration. It can be seen that the largest differences occur at the top and bottom of the dust layer, which may be due to spatial variation and/or gravitational settling.

### 3.2 Laboratory inter-comparisons

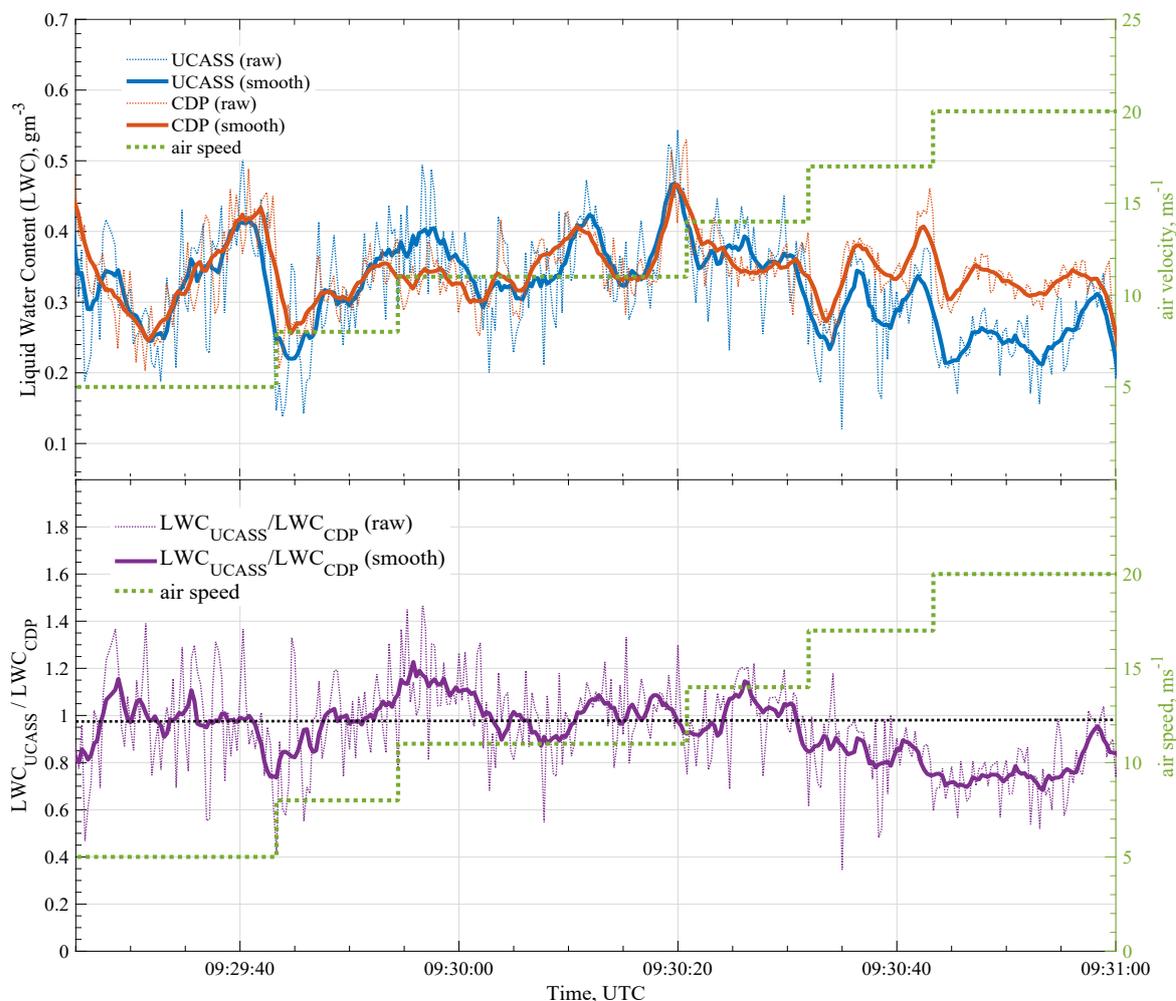
5 Inter-comparison tests were conducted in laboratory settings to assess the sizing and counting ability of the UCASS in comparison to reference instrumentation. A low-gain configuration of the UCASS (calibrated for water droplets) was tested at the Observatoire de Physique du Globe de Clermont-ferrand (OPGC) observatory, located atop the Puy De Dôme volcano at an altitude of 1465 m. The observatory consists of a rooftop measurement platform and an inbuilt wind tunnel capable of pumping through the passing clouds. As the UCASS requires an external air flow, the UCASS was tested in the wind tunnel, co-located  
10 with a Cloud Droplet Probe (CDP-2) from Droplet Measurement Technologies (DMT) (Lance et al., 2010). The CDP is an open geometry system which provides size and concentration measurements in the range 2–50  $\mu\text{m}$ , and is intended for air speeds of 10–250  $\text{ms}^{-1}$ . During a stratocumulus event, the UCASS and CDP were co-located in the wind tunnel, and the cloud



was drawn through at varying speeds from  $5 \text{ ms}^{-1}$  (the lowest operational velocity of the wind tunnel) to  $20 \text{ ms}^{-1}$ . Figure ?? shows the measured Liquid Water Content (LWC) by the UCASS (blue) and the CDP (orange). Both instruments measured at a rate of 1 Hz (shown by dotted lines), and the solid lines represent the smoothed data, using a 10-point moving average. The air speed is shown via a green dotted line corresponding to the right-hand y axis. It can be seen that, throughout the experiment,

5 both the UCASS and CDP capture the same temporal variations in the liquid water content, although the magnitudes start to diverge at air velocities above  $17 \text{ ms}^{-1}$ . This under-counting by the UCASS is expected due to the short ToF (discussed in Sect. 2.3 which causes the particles to be rejected. As discussed in Sect. 2.3, this ToF cut-off is intended to prevent miscounting of overly-long pulses caused by agglomerates. However, the ToF limits can be altered to suit the measurement platform. In the overlapping measurement range (air speeds between  $10$  and  $15 \text{ ms}^{-1}$ ), the ratio of UCASS measured LWC to CDP measured

10 LWC is equal to 1.02.



**Figure 17.** Top: measured Liquid Water Content (LWC) by co-located UCASS and CDP instruments at varying air speeds. The two instruments were installed side-by-side in the wind tunnel at Puy de Dôme during a stratocumulus event and the air speed was increased from  $5 \text{ ms}^{-1}$  to  $20 \text{ ms}^{-1}$ . Bottom: Ratio of the LWC measured by the UCASS to the LWC measured by the CDP. The black dotted line shows the 1:1 ratio.

#### 4 Conclusions

The Universal Cloud and Aerosol Sounding System (UCASS) is a lightweight, open-path Optical Particle Counter (OPC), designed for the measurement of micron-scale particles. The open-geometry design bypasses issues associated with narrow inlets such as clogging and counting uncertainties associated with complex air-flow systems. Furthermore, by removing the need for heavy or expensive pumps, the size, weight and cost of the instrument are kept to a minimum. The UCASS measures 180 mm long, 64 mm in diameter and weighs 280 g. The UCASS can then be used as a stand alone instrument whereby



data is logged autonomously via an on board SD card, or the UCASS can be interfaced via XDATA protocol with several commercially available meteorological sondes. Therefore, the UCASS is suitable to create low-weight payloads for dropsonde systems, balloon-borne systems or Unmanned Aerial Vehicles (UAVs).

The UCASS can be configured as a high-gain or low-gain mode, giving nominal measurable ranges of  $\approx 0.4$ – $17 \mu\text{m}$  and  $\approx$  5 1– $40 \mu\text{m}$ , respectively, with up to 16 configurable size bins. As the UCASS measures predominantly side-scattered light, the sizing ability is not hindered by the presence of mie oscillations that can cause uncertainties with forward scattering probes. However, as side-scattered light is more dependent upon the refractive index of the scattering particle, the exact detection limits are dependent upon the refractive index of the material being measured. To ensure accurate sizing, an 8+ point calibration is applied using a variety of particle standards including PSL, soda-lime glass and borosilicate glass. To account for differences 10 in refractive index, the geometric size is converted to scattering cross-section using mie-theory (Rosenberg et al., 2012), and the measured instrument response is plotted against scattering cross-section. For each unit, a relationship is found between the scattering cross-section ( $\sigma_{sca}$ ) and instrument response ( $AD$ ), in the form  $\ln(AD) = a \ln(\sigma_{sca}) + b$ , where  $a$  &  $b$  are constants found through calibration measurements. By using scattering cross section, rather than geometric diameter, this equation can be applied to particles of any refractive index or shape. Typically, the bin boundaries are chosen using a-priori information on 15 the particles being measured (i.e. mineral dust or water), however with this relationship established, post calibrations can be applied easily if post factum information of the aerosol properties becomes available. Due to this, the UCASS can be used to measure in mixtures of cloud and aerosol (as discussed in Sect. 3.1.1) where humidity profiles can be used to determine dry and cloudy air masses. The bin boundaries for different layers can then be modified for water or mineral dust accordingly.

The UCASS has been tested in a number of field and lab-based studies for the measurement of both aerosol and cloud 20 droplets. During a Saharan dust in Magdeburg, Germany, the UCASS was used along the KITsonde as a dropsonde system, and dropped from the Dornier 128 research aircraft. The volume size distribution measured by the UCASS shows good agreement with the AERONET inversion conducted at Leipzig, 100 km South-South-East of Magdeburg. The UCASS has also been tested as an upsonde system during the AER-D/SAVEX-D campaign, conducted over Cape Verde in 2015. The UCASS was interfaced with a GRAW DFM-09 and launched from Sal island, coincident with a SAVEX-D research flight conducted 80 25 km South-South-West of the launch site. Comparisons of the mass concentration measurements from the UCASS and the aircraft-mounted PCASP show both instruments agreement within 20%. The largest discrepancies are associated with the top and bottom of the dust layer, as anticipated from model results which show the profiles to be conducted on the edge of a dust layer. For more controlled tests, the UCASS was tested in a wind tunnel at Observatoire de Physique du Globe de Clermont-ferrand (OPGC) observatory, Puy De Dôme, during a stratocumulus event. The UCASS was co-located with a CDP-2 with a spatial 30 separation of  $<0.5 \text{ m}$ . The two instruments are designed for measurement in different ambient air speeds, but in the overlapping range ( $10$ – $15 \text{ ms}^{-1}$ ), the liquid water content measurements from the UCASS and CDP agree within 2%.

*Code and data availability.* The AERONET data used in Fig. 13 is available online via the AERONET site: <https://aeronet.gsfc.nasa.gov/>. PCASP data used in Fig. 16 are taken from SAVEX-D flight B934, access to this dataset can be requested through the Centre for Environ-



mental Data Analysis (CEDA) and is archived at : CEDA/Data Server/badc/ice-d/data/bae-146/b934-2015-aug-25. CDP data used in Fig. 17 was provided by OPGC and is not publicly available. All UCASS data, plot data and analysis code can be made available upon request by contacting the lead author.

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*Competing interests.* The authors declare that they have no conflict of interest.

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