



32 1 Introduction

33 Atmospheric aerosol particles are known to have an important effect on climate through directly
34 scattering or absorbing solar and terrestrial radiation as well as through indirect effects such as acting
35 as Cloud Condensation Nuclei (CCN) or Ice-Nucleating Particles (INPs) (Albrecht, 1989; DeMott et al.,
36 2010; Haywood and Boucher, 2000; Hoose and Mohler, 2012; Lohmann and Diehl, 2006; Lohmann
37 and Gasparini, 2017). Aerosol particles across the fine (diameter $< 2 \mu\text{m}$) and coarse ($> 2 \mu\text{m}$) modes
38 are important for these atmospheric processes. For example, aerosol in the accumulation mode are
39 important CCN (Seinfeld and Pandis, 2006), whereas supermicron particles are thought to contribute
40 substantially to the INP population (Mason et al., 2016; Pruppacher and Klett, 1997) Hence, being
41 able to sample across the fine and coarse modes is required to understand the role aerosol play in our
42 atmosphere. However, sampling aerosol particles without biases can be challenging, this being
43 especially so on a fast moving aircraft (Baumgardner et al., 2011; Baumgardner and Huebert, 1993;
44 McMurry, 2000; Wendisch and Brenguier, 2013).

45 It is necessary to sample aerosol from aircraft because in many cases aircraft offer the only
46 opportunity to study aerosol and aerosol-cloud interactions at cloud relevant altitudes (Wendisch and
47 Brenguier, 2013). However, the relatively high speeds involved present a set of unique challenges for
48 sampling aerosol particles. This is especially so for coarse mode aerosol which are prone to both losses
49 as well as enhancements because their high inertia inhibits their ability to follow the air stream lines
50 when they are distorted by the aircraft fuselage and the inlet (Brockmann, 2011; McMurry, 2000; von
51 der Weiden et al., 2009). Therefore, inlet design and characterisation becomes extremely important
52 when sampling aerosol particles.

53 In this study we characterise the inlet system used for collecting filter samples (known as the Filters
54 system) on board the UK's BAe 146 Facility for Airborne Atmospheric Measurements (FAAM) research
55 aircraft. This system has been used for many years, but its characterisation has been limited. Our goal
56 in this characterisation work was to define recommendations for the use of the inlet system to
57 minimise sampling biases and define the size limitation and the biases that exist. While the filter
58 samples could be used for a variety of offline analyses, we have done this characterisation with two
59 specific goals in mind: firstly, we want to use this inlet system for quantification of INP (the technique
60 for this analysis has been described previously (Price et al., 2018) and will not be further discussed
61 here); secondly, we have adapted and developed a technique for quantification of size distributions
62 and size resolved composition using Scanning Electron Microscopy (SEM). We use the SEM derived
63 size distributions in comparison with the size distributions obtained from the underwing probes to
64 experimentally test the inlet efficiency. These experiments are underpinned by calculations which
65 elucidate how the biases are impacted by variables such as flow speed, angle of attack and use of the
66 bypass system. Finally we present an example of the use of the inlet for determining the size resolved
67 composition of an aerosol sample collected from the FAAM aircraft.

68 2. Description and theoretical sampling characteristics of the filter inlet system on the Facility for 69 Airborne Atmospheric Measurements (FAAM) aircraft

70 Ideally, aerosol particles would be sampled through inlets without enhancement or losses. However,
71 this is typically not the case when sampling from aircraft, hence it is important to know how the size
72 distribution of the aerosol particles is affected by the sampling. Generally, an aircraft moves at high
73 velocities with respect to the air mass that it is being sampled. During sampling on the FAAM aircraft
74 the indicated airspeed is 100 m s^{-1} , which yields to a true airspeed that fluctuates between 100 and
75 120 m s^{-1} . The air mass has to decelerate when passing through the inlet (Baumgardner and Huebert,
76 1993) and this tends to result in inertial enhancement of coarse mode aerosol. There are also losses



77 through the inlet system, for example, through inertial impaction at bends or gravitational settling in
78 horizontal sections of pipework. These inlet characteristics need to be considered if the subsequent
79 analysis of the aerosol samples is to be quantitative. In this section we first describe the existing inlet
80 system and then present theoretical calculations for the size dependent losses and enhancements.

81 2.1 Description of the Filters system

82 The UK's FAAM BAe-146 research aircraft has two identical inlets for sampling aerosol onto filters for
83 offline analysis. This inlet system was used to sample aerosol particles on board of the C-130 aircraft
84 before being installed on the FAAM Bae-146 (Andreae et al., 1988; Andreae et al., 2000; Talbot et al.,
85 1990), and it has been used to sample aerosol particles on the FAAM Bae-146 e.g. (Chou et al., 2008;
86 Hand et al., 2010; Price et al., 2018; Young et al., 2016). A diagram of the inlet system can be seen in
87 Fig. 1. The two parallel inlet and filter holder systems, which each have a nozzle whose leading edge
88 profile follows the criteria for aircraft engine intakes at low Mach numbers (low speeds when
89 compared with the speed of sound; for FAAM during sampling this is ~ 0.3), and it is designed to avoid
90 the distortion of the pressure field at the end of the nozzle, flow separation and turbulence (Andreae
91 et al., 1988). The inlet has a bypass to remove water droplets or ice crystals through inertial separation
92 and also enhance the flow rate at the inlet nozzle (Talbot et al., 1990). The flow through the bypass
93 (bypass flow) can be regulated using a valve and it is driven passively by the pressure differential
94 between the ram pressure inlet and the Venturi effect on the exhaust. After turning inside the aircraft,
95 the airstream containing the aerosol particles continue through the filter stack after passing a valve.
96 The air flow through the filter (filter flow) is measured by a mass flow meter, which reports the
97 standard litres sampled (273.15 k, 1013.529 hPa). The signal is integrated by an electronics unit to give
98 the total volume of air sampled for any given time period. There is also a valve between the pump and
99 the flow meter. The valve allows the inlet and pump to be isolated from the filter holder when
100 changing the filter. The system uses a double-flow side channel vacuum pump model SAH55 made by
101 Rietschie, aided by the ram effect of the aircraft. The flow rate at the inlet nozzle (total flow) is the
102 sum of the bypass flow and the filter flow. The inlet nozzle is located at 19.5 cm of the aircraft fuselage,
103 so the sampling is carried out in the free stream, outside the boundary layer.

104

105 2.2 Sampling efficiency

106 We present theoretical estimates of the losses and enhancements due to aspiration, inlet inertial
107 deposition, turbulent inertial deposition, inertial deposition in bends and gravitational effects in Fig
108 2a. We used the term 'efficiency' to define the ratio between the number concentrations of particles
109 after they were perturbed relative to the unperturbed value. If the efficiency is above one, the number
110 of particles is enhanced whereas if it is below 1, particles are lost.

111 The sampling efficiency of any inlet depends strongly on the flow rates and the flow regime (laminar
112 vs turbulent). Filter flow rates for 0.4 μm polycarbonate filters normally vary between 10 and 50 L min^{-1}
113 depending on altitude (see section 2.3 for a discussion of flow rates). The bypass flow rate (when it
114 is fully open) can go up to 35 L min^{-1} at 30 m and 22 L min^{-1} at 6 km, but it is not measured routinely.
115 In the 2.5 cm diameter section of the inlet, the Reynolds number (Re) is below the turbulent regime
116 threshold ($\text{Re} > 4000$) for flow rates below 65 L min^{-1} . For larger values of Re, the flow starts becoming
117 turbulent. At the inlet nozzle, where the diameter 0.7 cm, Re is above 4000 for flow rates above 20 L min^{-1} ,
118 so the flow is briefly in the turbulent regime at the inlet for most sampling conditions. Fully
119 characterising the losses and enhancements of aerosol particles passing through the inlet is very



120 challenging since there are several aerosol size dependent mechanisms than can enhance or diminish
121 the amount of aerosol particles that arrive at the filter.

122 Here we have considered the most important of these mechanisms in order to estimate the inlet
123 efficiency (see Fig. 2a) for a total flow rate of 50 L min^{-1} . These loss mechanisms and their importance
124 in this inlet system are defined as follows (a discussion on the choice of equations and how they have
125 been applied can be found in Appendix A):

126 *Aspiration efficiency* accounts for the fact that the speed of the sampled air mass (U_0) and the speed
127 of the air through the beginning of the inlet nozzle (U) are different. In the case of the filter inlet
128 system on the FAAM aircraft, the speed of the air mass is greater than the speed through the inlet
129 (sub-isokinetic conditions), which leads to an enhancement of larger aerosol particles. Here, we have
130 used the empirical equation as develop in Belyaev and Levin (1972) and Belyaev and Levin (1974). As
131 one can see in Fig. 2a this mechanism enhances aerosol particles, tending to 1 for small diameters and
132 to the ratio U/U_0 for large ones.

133 *Inlet inertial deposition* is the inertial loss of aerosol particles within the nozzle because the flow
134 expands inside the nozzle and the trajectory is therefore bent towards the wall. It has been
135 characterised using the equation given in Liu et al. (1989) which quantifies this effect. In Fig. 2a one
136 can see that it produces some losses, with a minimum efficiency of down to 50% for sizes about $6 \mu\text{m}$,
137 without affecting the lower and upper limit of the aerosol size.

138 *Turbulent inertial deposition* happens when some particles are collected by the wall due to
139 turbulences in the system. In our case, this occurs throughout the whole inlet system for flow rates
140 above 65 L min^{-1} and only occurs in the inlet nozzle for flow rates below this threshold. We have used
141 the equation given by Brockmann (2011) in order to account for this mechanism, which can be seen
142 in Fig. 2a. This mechanism gradually decreases the efficiency for aerosol particles above $5 \mu\text{m}$.

143 *Bending inertial deposition* of aerosol particles is important in this inlet system at the 45° bend, where
144 some particles are not able to follow air streamlines at bend. We have characterised these losses using
145 the equation given in (Brockmann, 2011). This efficiency mechanism, which can be seen in Fig. 2a,
146 adds a size cut off with a D50 value at $\sim 25 \mu\text{m}$.

147 *Gravitational settling* of aerosol particles was considered using the equations developed in Heyder
148 and Gebhart (1977) and Thomas (1958), as stated in Brockmann (2011). This efficiency mechanism
149 adds another size cut off with a D50 value of $35 \mu\text{m}$, as one can see in Fig. 2a.

150 *Diffusional efficiency and filter collection efficiency* have not been included in Fig. 2. The first
151 mechanism has been calculated using the analytical equation given by Gormley and Kennedy (1948),
152 but it is not shown since it is very close to 1 for all the considered size range. For the filter types and
153 pore sizes we used, filter collection efficiency is also close to a 100% across the relevant size range
154 (Lindsley, 2016; Soo et al., 2016).

155 *Anisoaxial losses* are losses produced by the fact that the inlet is not aligned with the velocity of the
156 air mass, being offset by an angle, θ (related to the angle of attack). The anisoaxial sampling can affect
157 the sub-isokinetic efficiency, but using the equations given by Hangal and Willeke (1990a), we
158 calculated that this effect is minimal for our conditions. In addition, anisoaxial sampling can lead to
159 inertial losses when particles impact the inner walls of the inlet. This phenomena has been quantified
160 using the equations in Hangal and Willeke (1990b) and the results can be seen in Fig. 3. As one can
161 see, this efficiency mechanism adds an additional cut off for large aerosol particles (with values of D50
162 down to $\sim 20 \mu\text{m}$), depending on the value of the sampling angle.



163 One can see all the efficiency mechanisms combined for four different flow rates in Fig. 2b. These have
164 been derived by multiplying all the efficiencies for the individual mechanisms. This overall efficiency
165 is the ratio between the particles that reach the filter and the particles in the ambient air. The
166 sampling efficiency for the submicron aerosol is close to 1. At sizes above $1\ \mu\text{m}$, the different loss
167 mechanisms become increasingly significant. For the range of flow rates considered, the efficiency
168 approaches zero between 20 and $50\ \mu\text{m}$, with D50 values in between ~ 13 and $\sim 33\ \mu\text{m}$ (although these
169 values could be lower under certain values of angles of attack if considering the anisoaxial losses of
170 from Fig. 3, which haven't been included). For the $80\ \text{L min}^{-1}$ case, the flow is turbulent through all the
171 pipe, leading to enhanced losses of coarse aerosol particles which partially compensate the sub-
172 isokinetic enhancement of the system.

173 One can also see that the sub-isokinetic enhancement of large aerosol particles increases when
174 decreasing the flow rate of the system. This effect is about a factor 3.5 for $10\ \mu\text{m}$ particles when
175 sampling at $15\ \text{L min}^{-1}$, but only a factor of two at $50\ \text{L min}^{-1}$. The sub-isokinetic enhancement can be
176 mitigated using the bypass, which enhances the flow through the nozzle. This can be seen in Fig. 2c
177 where one can see a comparison between the total efficiency of a $20\ \text{L min}^{-1}$ flow rate through the
178 filters with no bypass flow and the same case when the bypass is open. Since the considered bypass
179 flow is comparable to the flow rate through the filters, the difference between the total flows for the
180 two cases is approximately a factor 2. As a consequence, the maximum sub-isokinetic enhancement
181 of large aerosol particles is almost a factor 2 larger when sampling with the bypass closed. Hence, the
182 sub-isokinetic enhancement can be reduced by keeping the bypass fully open.

183

184 2.3 Sampling flow rate

185 Here we show flow rate data from four field campaigns in order to examine how the flow rate of the
186 filter inlet system varied based on different factors. We have used the data collected during the ICE-D
187 campaign, in Cape Verde during August 2015 (Price et al., 2018). The rest of the data is from some
188 flight test carried out during 2017 and 2018, and three field campaigns. The first one was EMERGE,
189 based in south east England, in July 2017. The second one was VANAHAEIM, based in Iceland in
190 October 2017. The last campaign was MACSSIMIZE, based in Alaska in 2018. The flow rate of the inlet
191 system is known to vary with altitude, with a lower flow rate at high altitudes because of the reduced
192 pressure differential across the filter and the fact that the pump efficiency decreases at low pressure.
193 In addition, it changes depending on the filter type and the pore size.

194 In Fig. 4, where all the flow rate data has been presented, one can see that the flow rate tends to
195 decrease with altitude and change with filter type as expected, but the flow rates are not consistent
196 for each altitude and filter type, varying up to a factor two for each filter type/line/altitude/campaign.
197 The filter type effect on flow rate can be seen in Fig. 4, where the average flow rate for $0.4\ \mu\text{m}$
198 polycarbonate filters is about twice the flow rate of the $0.45\ \mu\text{m}$ PTFE filters. In order to investigate
199 the inconsistency in the flow rate at each altitude, we analysed the flow rate data by comparing it with
200 different parameters (ambient air and cabin temperature, ambient air and cabin pressure, wind
201 direction and speed with respect to the aircraft movement), but there was no correlation with any of
202 these variables. Different mesh supports were used, but this does not affect the flow rate significantly
203 according to some ground based tests. We checked the flow rate through each sampling period and
204 found it did not change over time on a particular filter set (even after stopping the sampling and
205 starting it again). In addition, we performed some tests on the ground and during flights to study the
206 effect of potential leaks by inserting paper disks of the same dimension as the filters in the filter
207 holders and found no evidence of leaks in the system.



208 We conclude that this variability in the flow rate comes from variability in the pump performance in
209 combination with subtle differences in individual filter pairs. The side displacement pump is not the
210 ideal pump for this system and operates at its maximum capacity. Hence, we suggest that to improve
211 the performance of the system that flow rates are actively controlled and also the side displacement
212 pump is replaced with a more appropriate design. This would also have the advantage that flow rates
213 would be maintained at smaller pressure drops and allow sampling at higher altitudes.

214

215 **3. FAAM underwing optical particle counters**

216 Later in the paper we compare results from the underwing optical particle counters with our electron
217 microspore derived size distributions, hence we describe the optical instruments here. The BAe-146
218 FAAM research aircraft operates underwing optical particle counters to measure aerosol size
219 distributions. These include the Passive Cavity Aerosol Spectrometer Probe 100-X (PCASP) and the
220 Cloud Droplet Probe (CDP). The PCASP measures particles with diameters in the approximate range
221 0.1-3 μm and the CDP measures the particles with diameters in the range of 2-50 μm . These
222 instruments are placed outside the aircraft fuselage, below the wings. These instruments and the
223 methods for calibration are described in (Rosenberg et al., 2012).

224 The instruments were calibrated and had optical property corrections applied as per Rosenberg et al.
225 (2012). We used a refractive index of $1.56 + 0i$ and a spherical approximation (Mie theory) in the
226 optical property corrections. In Fig. 5, one can see a sensitivity test on the refractive index value we
227 used in order to examine how variability in refractive index affect the bin centres position, their width,
228 and therefore the size distribution obtained from the PCASP and CDP. As one can see in Fig. 5a,
229 modification of the real part of the refractive index from 1.5 to 1.7 can change the position of the
230 PCASP bin centres up to a factor 1.5, but its effect on the CDP is not significant. When varying the
231 imaginary part of the refractive index from 0 to 0.01, the bin centre positions of the first half of the
232 range of the PCASP and CDP do not change but it can change the position of the bins of the end of the
233 range of both instruments (less than a factor 1.5). However, for the purposes of this work, the
234 differences produced by the variation in the refractive index are not large enough to modify the
235 conclusions of the analysis, therefore we use a value of $1.56 + 0i$.

236 The chosen refractive index range for this sensitivity analysis can be justified on the basis that the SEM
237 compositional analysis showed that the composition of the aerosol samples used in this study was
238 very heterogeneous, dominated by carbonaceous particles (biogenic, organic and black carbon) and
239 with some contributions of mineral dust and other particle types. Values of the real part of the
240 refractive index in the 1.5 to 1.6 range are compatible with sodium chloride and ammonium sulphate
241 (Seinfeld and Pandis, 2006), as well as most mineral dusts (McConnell et al., 2010). The range is very
242 close to values for the real part of the refractive index of organic carbon but below the values for black
243 carbon (Kim et al., 2015). As a consequence, the refractive index choice might not be accurate for a
244 black carbon dominated sample. However, black carbon is not likely to dominate in the size range
245 where a value of the real part of the refractive index of 1.7 dramatically changes the size distribution
246 (diameters above 0.5 μm) (Seinfeld and Pandis, 2006), so our refractive index choice is valid. In Fig.
247 5b one can see that changing the imaginary part of the refractive index from 0 to 0.01 only produces
248 small changes in the distribution. The imaginary part of the refractive index of many aerosol types as
249 sodium chloride, sulphates and mineral dust falls within the shown range (Seinfeld and Pandis, 2006),
250 (McConnell et al., 2010). For values of the imaginary part of the refractive index above 0.01 (not shown
251 in the image), the size distribution dramatically changes for sizes above 1 μm (but not for smaller
252 values of it), overlapping and disagreeing with the CDP. However, values above 0.01 in the imaginary



253 part of the refractive index can only happen in black carbon, which will dominate only in the submicron
254 sizes (Seinfeld and Pandis, 2006). The submicron part of the size distribution doesn't change for values
255 of the imaginary part of the refractive index above 0.01, so our refractive index choice is still
256 acceptable even for samples with significant contributions from black carbon in submicron sizes.

257 For the PCASP-CDP, we have considered two uncertainty sources. The first one is the Poisson counting
258 uncertainty in the number of particles in each bin and the second one is the uncertainty in the bin
259 width that is given by the applied optical property corrections. Both sources have been propagated in
260 order to obtain the errors of $dN/d\log D_p$ and $dA/d\log D_p$. Other uncertainties such as the refractive
261 index assumption or particle shape effect, as well as the uncertainty in the bin position haven't been
262 shown in this study. Sampling biases haven't been quantified or corrected yet so they haven't been
263 included. The size distributions produced by the PCASP-CDP have been taken as a reference value for
264 the purposes of this study.

265

266 **4. Scanning Electron Microscopy technique for aerosol characterization**

267 Scanning Electron Microscopy is used in order to study composition and morphology of aerosol
268 particles, in a similar way to previous works such as Chou et al. (2008), Young et al. (2016) and Price
269 et al. (2018). We use a Tescan VEGA3 XM scanning electron microscope (SEM) fitted with an X-max
270 150 SDD Energy-Dispersive X-ray Spectroscopy (EDS) system controlled by an Aztec 3.3 software by
271 Oxford Instruments, at the Leeds Electron Microscopy and Spectroscopy Centre (LEMAS) at the
272 University of Leeds. In order to get data from thousands of particles in an efficient way, data collection
273 was controlled by the AztecFeature software expansion.

274 Aerosol particles were collected with the filter inlet of the FAAM aircraft on polycarbonate track
275 etched filters with 0.4 μm pores (Whatman, Nucleopore). Samples for SEM are usually coated with
276 conductive materials in order to prevent the accumulation of charging on the sample surface (Egerton,
277 2005). For aerosol studies, materials like gold (Hand et al., 2010), platinum (Chou et al., 2008), or
278 evaporated carbon (Krejci et al., 2005; Reid et al., 2003; Young et al., 2016) have been used. When it
279 comes to choosing which signal to detect, some previous studies used backscattered electrons (Gao
280 et al., 2007; Price et al., 2018; Reid et al., 2003; Young et al., 2016) and some others choose secondary
281 electrons (Kandler et al., 2007; Kandler et al., 2011; Krejci et al., 2005). We started the development
282 of this analysis using a carbon coating and the backscattered electron detector. This technique
283 produced reproducible images and almost no artefacts from the pore edges, consistent with Gao et
284 al. (2007). However, we noticed that we were undercounting a significant fraction of the small carbon
285 based particles, which looked transparent under the backscattered electron imaging but not under
286 the secondary electron detector, as seen in Fig. 6. This likely happened because the contrast in the
287 secondary electron images mainly depends on the topography of the sample whereas the contrast in
288 the backscattered electron images depends on the mean atomic number of each sample phase
289 (Egerton, 2005). Since the polycarbonate filters are made of C and O, particles containing only these
290 elements in a similar proportion to the background did not exhibit a high contrast under the
291 backscattered electron detector (Laskin and Cowin, 2001). However, when using secondary electron
292 imaging with carbon coatings, images were less reproducible and contained artefacts from the pore
293 edges, probably resulting from charging or topographical effects. We found that coating the samples
294 with 30 nm of iridium helps to improve the secondary electron image reproducibility and reduced the
295 pore edge artefacts as well as allowing us to locate small organic particles. An increase in the size of
296 the particle as a consequence of the coating may introduce an uncertainty in the size of the smallest



297 particles. An additional advantage of using Ir is that the energy dispersive X-ray spectrum of Ir does
298 not overlap greatly with the elements of interest.

299 In the SEM the sample was positioned at a working distance of 15 mm. The SEM's electron beam had
300 an accelerating voltage of 20 KeV and a spot size chosen to produce the optimum number of input
301 counts in the EDS detector. Images are taken at two different magnifications with a pixel dwell time
302 of 10 μ s and a resolution of 1024 x 960 pixels per image. High magnification images (x 5000 or similar)
303 were used to identify particles down to 0.3 or 0.2 μ m depending on the sample, and medium
304 magnification images (x1500 or similar) are used to identify particles down to 1 μ m. A brightness
305 threshold with upper and lower limits that correspond to pixels of certain shades of grey was manually
306 adjusted for each image by the operator to discriminate particles from the background. Based on the
307 manually set brightness threshold, AztecFeature identified the pixels that fall within the limits as
308 aerosol particles and calculated several morphological properties of the particle as cross sectional
309 area, length, perimeter, aspect ratio, shape factor or equivalent circular diameter. The equivalent
310 circular diameter is defined as $\sqrt{4 A \pi^{-1}}$, where A the cross sectional area of the aerosol particle.

311 For this analysis we placed a section of the 47 mm filter on a 25 mm stub. In order to collect
312 morphological and chemical information from a few thousand particles, we only scanned a fraction of
313 the filter. We collected information from 5 to 20 different areas, and each area consisted of a montage
314 of several SEM images. The areas were chosen by the user from all over the surface of the selected
315 fraction of the filter, since aerosol particles were evenly distributed all over the central ~30mm of the
316 filter (the area which exposed to the air) as one can see in Fig. 7. Each area was selected in the
317 software, manually adjusting the particle detection threshold. The Z position of the stage was also
318 adjusted manually for each image in order to produce properly focused images. After doing this, the
319 image scanning and EDS acquisition was performed in an automated way. Morphological information
320 was recorded for all particles with an ECD greater than the specified size threshold (typically 0.2 or 0.3
321 μ m).

322 EDS analysis was restricted to the first 12 or 15 particles detected in each image. This reduces the
323 likelihood of charging problems caused by exposure to the electron beam. The software performed
324 EDS in the centre of the particles, obtaining around 50,000 counts per particle. The raw data for any
325 given particle were matrix corrected and normalised by the AZtec software to produce element weight
326 percent values with a sum total of 100%, using a value of the confidence interval of 2 (a further
327 discussion on the confidence interval can be seen in Fig. S1). Then particles were categorised based
328 on their chemical composition using a classification scheme which can be created and modified within
329 the AztecFeature software. The characteristic X-rays taken at one point are emitted by a certain
330 interaction volume which is bigger than some of the analysed particles (typically $< 2\mu\text{m}^3$, decreasing
331 with atomic number and increasing with incident electron energy). As a consequence, a part of the X-
332 ray counts attributed to each particle come from the background (C and O from the polycarbonate
333 filter and Ir from the coating) and the weight percentages obtained from the X-ray spectra do not
334 match the actual weight percentages of the particle itself. As a consequence, when categorising the
335 particles based on their composition, we only use the presence or absence of certain elements, and
336 the ratio between the weight percentages of non-background elements. The classification scheme
337 works by checking if the composition of each particle falls within a range of values which are manually
338 defined by the user. Particles not matching the first set of rules are tested again for a second set of
339 rules, and so on, until reaching the last set of rules. A few sets of rules can be merged into a category.
340 In the supplementary information (Fig. S3), we give the details of the 32 sets of rules used, which are
341 then summarised into 10 composition categories. A description of the most abundant elements in
342 each category and an interpretation of these categories is included in Sect. 5.



343 The detection of particles has certain limitations. The edges of the pores can look brighter than the
344 rest of the filter in the SE images (probably because they consist of a larger surface area from which
345 secondary electrons can be generated, hence a larger signal). As a consequence, they can look like
346 $\sim 0.2 \mu\text{m}$ particles, which is the main reason why particles below $0.3\text{--}0.2 \mu\text{m}$ (depending on the sample)
347 are not included in this analysis. These artefacts had a chemical composition similar to the filter, so
348 they were labelled as “Carbonaceous” by the classification scheme, falling at the same category as
349 most biogenic and black carbon particles. However, these artefacts were only around 1 to 10 percent
350 depending on the sample. If they appear in larger quantities, they can be removed manually after or
351 during the analysis. Another limitation arises from the fact that some aerosol particles did not have
352 sufficient brightness in the SE image and were not detected as a particle. This happens more
353 frequently for smaller particles, but it can also happen with some larger particles, particularly if they
354 are only composed of Na and Cl or S. This issue can be addressed if necessary by setting a very low
355 limit of detection, which adds lots of artefacts as well as the low brightness particles, and then
356 removing the artefacts manually (the artefacts can be easily identified by the user). In other infrequent
357 instances, only a fraction of the particle had a brightness above the threshold, so they were detected
358 as a smaller particle or multiple smaller particles, or if two particles are close enough, they can be
359 detected as a single larger particle. However, we feel that in the vast majority of the cases a
360 representative cross sectional area of the particle was picked by the software.

361 Blank polycarbonate filters can contain some particles on them from manufacturing or transport
362 before being exposed to the air. In addition, handling and preparing the filters can introduce additional
363 particles to it. In order to assess these artefacts, we scanned a few clean blank filters. We also
364 examined a filter that had been brought to the flight, loaded in the inlet system (but not exposed to a
365 flow of air), and then stored at $-18 \text{ }^\circ\text{C}$ for a few months (like most of the aerosol samples on filters).
366 The results of both the handling blank and the blank can be seen in Fig. S2. The number of particles is
367 very low, typically about the order of magnitude of one particle per $100 \text{ by } 100 \mu\text{m}$ square, which is
368 well below the typical particle loading on a filter exposed to the atmosphere. In Fig. S2 one can see
369 that the vast majority of particles found in both blank filters and the handling blank belong to the
370 metallic rich category. However, further examination of the composition of these particles revealed
371 that almost all of them were Cr rich particles (about 97 % in the case of the blank filters and about
372 96% in the case of the handling blank). As a consequence, we excluded all the Cr rich particles from
373 the analysis of atmospheric aerosol (it was only ever a very minor component). By doing this, we make
374 sure that we excluded more than half of the artefacts of the analysis. The composition of the particles
375 present in the blank filters and in the handling blank was very similar, suggesting that most of these
376 artefacts are not produced by the loading, manipulation and storage of the filter. However, there was
377 a very small but significant contribution of mineral dust origin particles (Al-Si rich, Si rich and Si only)
378 for sizes in between 0.7 and $5 \mu\text{m}$ in the handling blank, which should be taken into consideration.

379

380 **5. Inlet characterisation and sampling efficiency using Scanning Electron Microscopy**

381 In order to experimentally test the inlet efficiency, to complement the efficiency calculations
382 presented in Section 2.2, we have used SEM to quantify the size distribution of particles collected on
383 filters (Sect. 4) and compare this with the measurements from the under-wing optical probes (Sect.
384 3). The calculations in Sect. 2.2 suggest that there is an enhancement of the coarse mode aerosol
385 particles, which is larger when sampling with the bypass closed. To test this we have collected aerosol
386 onto $0.4 \mu\text{m}$ pore size polycarbonate filter in both lines in parallel. In one of the lines, the bypass was
387 kept open, and in other line the bypass was kept closed. Using our SEM approach described in the
388 Sect. 4, we calculated the size distribution of the aerosol particles on top of each filter. We compared



389 these size distributions with the ones measured by the underwing optical probes (PCASP-CDP), as
390 described in Sect. 3. We performed the test twice in two different test flights based in the UK.

391 The results of these comparisons can be seen in Fig. 8 and Fig. 9 for both number and surface area size
392 distribution. There are some discrepancies between the optical probes and the SEM size distributions
393 from the filters, which has also been reported in previous works (Chou et al., 2008), (Price et al., 2018;
394 Young et al., 2016). There is significant disagreement between the submicron particles detected on
395 top of the filter and the submicron particles detected in situ by the PCASP-CDP in Fig. 9a in a very
396 similar way to Young et al. (2016). Since the 0.4 μm polycarbonate filters have a high collection
397 efficiency at these length scales (Lindsley, 2016; Soo et al., 2016), the disagreement at the submicron
398 regime could be produced by several effects. Some small particles that may not have sufficient
399 brightness to be detected might produce some undercounting, despite the fact we made efforts to
400 mitigate this problem. Also, volatilization of certain type of aerosol particles (which are more abundant
401 in the submicron fraction (Seinfeld and Pandis, 2006)) can happen during heating or sampling (Bergin
402 et al., 1997; Hyuk Kim, 2015; Nessler et al., 2003) and this effect could be enhanced by the fact that
403 samples are exposed to high vacuum during the SEM analysis. The presence of nitrates could be tested
404 using an aerosol mass spectrometer. In addition, the SEM techniques measure the dry diameter and
405 the optical probes measure the aerosol diameter at ambient humidity. This hygroscopic effect is
406 known to shift the dry size distributions to smaller sizes, which could also explain part of the
407 disagreement (Nessler et al., 2003; Young et al., 2016). Disagreement in the measurements can be
408 also produced by the fact that the techniques are measuring different diameters; optical diameter in
409 the case of the PCASP-CDP and circular equivalent geometric diameter in the case of the SEM.

410 One can see that the concentration of aerosol particles measured by the SEM on the filters was higher
411 than the particles detected by the optical probes for sizes above $\sim 8 \mu\text{m}$ in Fig. 8 and Fig. 9. These
412 results are consistent with Price et al. (2018), where they observed a similar enhancement of large
413 aerosol particles in two mineral dust dominated samples collected close to Cape Verde. In addition,
414 the enhancement was larger when sampling under closed bypass conditions. The results of these
415 comparisons are in agreement with the theoretical calculations in Sect. 2.2.

416 In Fig. 10 we have presented some other SEM size distributions compared with the PCASP-CDP data
417 from three different aerosol samples in contrasting locations. Since these data were taken during the
418 scientific field campaigns and not test flights, we only collected one polycarbonate filter for SEM since
419 the other line was used for INPs analysis on Teflon filters. All the sampling was done with the bypass
420 open. The agreement between the optical and SEM obtained size distributions in Fig. 10 is similar to
421 the one in Fig. 8 and Fig. 9 (for the open bypass case). One can observe in Fig. 10c that there was a
422 loss of particles smaller than $\sim 0.5 \mu\text{m}$ and also that in Fig. 10a and Fig. 10b there was an enhancement
423 of the coarse mode.

424 The data shown in this section have been obtained in very different locations (South England in the
425 case of Fig. 8, Fig. 9 and Fig. 10a, Iceland in the case of Fig. 10b and north Alaska in the case of Fig. 10c.
426 As a consequence, the studied aerosol samples are very different in both morphology and chemistry.
427 From the comparisons, we can state that the sampling carried out by the filter system has certain
428 biases, but it captures particles with a size distribution with similar features to the ones measured by
429 the underwing optical probes.

430



431 6. Recommendations for aerosol sampling with the Filters system on the FAAM aircraft

432 Based on the calculations in Sect. 2.1, we suggest keeping the total flow rate (including the flow
433 through the filters measured by the electronics box plus the bypass flow, which can be between 20
434 and 35 L min⁻¹) above 50 L min⁻¹. Below this range, the sub-isokinetic enhancement of large aerosol
435 particles is above a factor 2, according to the calculations in Sect. 2.2 that can be seen in Fig. 2b. For
436 total flow rates above 65 L min⁻¹, the flow becomes turbulent throughout the line, which associated
437 losses. However, the calculations shown in Fig. 2c indicate that the combination of the isokinetic
438 enhancements and turbulent losses at 80 L min⁻¹ lead to a reasonably representative sampling, but
439 when it reaches 150 L min⁻¹, the position of the D50 drops to 6.5 µm (not shown in the graph) so such
440 a high flow rate would not be ideal if the user wants to sample coarse aerosol particles. Hence, we
441 recommend an operational upper limit of 80 L min⁻¹. For 0.45 µm PTFE filters and the 0.4 µm
442 polycarbonate filters presented in Fig. 4, sampling close to this flow rate range is often achievable by
443 keeping the bypass open, since this increases the total flow rate and brings it closer to the suggested
444 range, as one can see in Fig. 2c. If other filter types are used, these results should be taken in
445 consideration when choosing the pore size or equivalent pore size in order to avoid dramatic sampling
446 biases.

447 We already mentioned in Sect. 2.3 that we recommend replacing the side displacement pump with a
448 design that would provide a greater pressure drop. In addition, we also recommend that the bypass
449 flow rate is also routinely measured and controlled in order that the flow at the inlet nozzle can be
450 optimised while sampling.

451 7. SEM compositional categories

452 Here we describe the 10 categories we have used in our compositional analysis, which are a summary
453 of the 32 rules described in the supplementary information. The approach has some similarities with
454 the ones in previous studies (Chou et al., 2008; Hand et al., 2010; Kandler et al., 2011; Krejci et al.,
455 2005; Young et al., 2016), but it is distinct. Because of the fact that the filter is made of C and O,
456 background elements (C and O) were detected in all the particles. Particles in each category can
457 contain smaller amounts of other elements apart from the specified ones. This classification scheme
458 has been designed a posteriori to categorise the vast majority of the aerosol particles in the three field
459 campaigns previously described and some ground collected samples in the UK and Barbados. The main
460 limitation of the classification scheme is the difficulty to categorise internally mixed particles. The
461 algorithm has been built in a way it can identify mixtures of mineral dust and sodium chloride (they
462 appear as mineral dust but they could be split into a different category if necessary) and sulphate or
463 nitrate ageing on sodium chloride (they appear as Na rich but it could also be split into a different
464 category). However, other mixtures of aerosol wouldn't be identified, and they would be categorised
465 by the main component in the internal mixture in most cases.

466 7.1. Carbonaceous

467 The particles in this category contained only background elements (C and O). The components of the
468 carbonaceous particles consist in either black carbon from combustion processes or organic material,
469 which can be either directly emitted from sources or produced by atmospheric reactions (Seinfeld and
470 Pandis, 2006). Particles containing certain amount of K and P in addition to the background elements
471 were also accepted in these category. These elements are consistent with biogenic origin aerosol
472 particles (Artaxo and Hansson, 1995). Distinction between organic and black carbon aerosol
473 unfortunately could not reliably be done. Since N is not analysed in our SEM set up, any nitrate aerosol
474 particle would fall into this category if it is on the filter. However, since these particles are semi-



475 volatile, some of these aerosol particles would not resist the low pressure of the SEM chamber. This
476 could be further investigated in the future.

477 7.2. S rich

478 Aerosol particles in this category contained a substantial amount of S. These EDS signals are
479 compatible with sulphate aerosol particles, which are solid or liquid sulphuric acid particles (Kumar
480 and Francisco, 2017). In the same way as the nitrates, this particles are semi-volatile and some of them
481 might not resist the low pressure of the SEM chamber.

482 7.3 Metal rich

483 The composition of particles in this category is dominated by one of the following metals: Fe, Cu, Pb,
484 Al, Ti, Zn or Mn. These EDS signatures are compatible with metallic oxides or other metal rich particles.
485 These metal containing particles can originate from both natural sources and anthropogenic sources.
486 Some metallic oxides are common crustal materials that could go into the atmosphere but are also
487 produced during some combustion processes (Seinfeld and Pandis, 2006). In addition, many types of
488 metal and metallic derivatives particles are produced as component of industrial emissions and other
489 anthropogenic activities (Buckle et al., 1986), (Fomba et al., 2015).

490 7.4. Na rich

491 Sodium chloride particles are the main component of the sea spray aerosol particles which are emitted
492 through wave breaking processes (Cochran et al., 2017). These particles can age in the atmosphere by
493 reacting with atmospheric components such as sulphuric or nitric acid (Graedel and Keene, 1995),
494 (Seinfeld and Pandis, 2006). As a consequence of this reaction, a part of their Cl content will end up in
495 the gaseous phase (as HCl), leading to an apparent chlorine deficit in the aged sea spray aerosol
496 particles. Particles in this category have an EDS signature compatible with sea spray aerosol particles
497 since they are identified by the presence of Na, containing in most cases Cl and/or S (N is not included
498 in our SEM analysis).

499 7.5 Cl rich

500 Particles in this category contained mainly Cl and sometimes also K but never Na, so they are not
501 sodium chlorine particles. Significant concentrations of Cl and metals in aerosol particles have been
502 linked to industrial activities and automobile emissions (Paciga et al., 1975), whereas Cl and K in
503 aerosol particles could be originated by the use of fertilisers (Angyal et al., 2010), or emitted during
504 pyrotechnic events (Crespo et al., 2012).

505 7.6 Ca rich

506 The composition of the particles in this category is dominated by Ca. In this category, particles
507 containing only Ca (plus C and O, the background elements) are consistent with calcium carbonate, a
508 major component of mineral dust (Gibson et al., 2006). If other elements such as Mg and S are present,
509 the signature of the particles compatible with some mineral origin elements as gypsum and dolomite
510 respectively. In addition, presence of minor amounts of Si, Al and other elements could indicate mixing
511 of these Ca rich particles with some other soil components as silicates. However, since Ca is a biogenic
512 element, we cannot discard the biogenic origin of some of the Ca-rich particles (Krejci et al., 2005).

513 7.7 Al-Si rich

514 Particles in the Al-Si rich category were detected by the presence of Al and Si as major elements. Very
515 often, this particles also contained smaller amounts of Na, Mg, K, Ca, Ti, Mn and Fe. Particles in this



516 category are very likely to have mineral origin and are commonly described as aluminosilicates which
517 include a range of silicates such as feldspars and clays (Chou et al., 2008; Hand et al., 2010). Mixed
518 mineral origin particles containing both Al and Si can also fall into this category. Strong presence of Na
519 and Cl could indicate internal mixing with some sea spray aerosol, whereas a strong S presence could
520 indicate atmospheric acid ageing.

521 7.8 Si only

522 The particles in this category contained only Si apart from the background elements. Particles in this
523 category are very likely to be a silica polymorph (mainly quartz), one of the major components of the
524 earth's crust. Since we cannot determine if the C signal in the EDS of these particles is produced from
525 the background or from the particle itself, a particle containing only C, Si and O would fall into this
526 category, however, mineral phases containing these elements are extremely rare.

527 7.9 Si rich

528 The composition of these particles was dominated by Si, and other elements Na, Mg, K, Ca, Ti, Mn and
529 Fe. The main difference with the particles in Sect. 7.7 is that the ones described here didn't contain Al
530 above the limit of detection. The EDS signal of particles in this category is compatible with any silicate
531 that does not contain Al as a major component in its phase such as talc or olivine. The only exception
532 is quartz, which falls in the 'Si only' category described above. They could also be internal mixtures of
533 silica or silicates without aluminium as a major component in its phase. Because of the high limit of
534 detection of the Al (See the SI), some particles in this category could contain small amounts of Al, and
535 should belong to Al-Si rich category. As in the Al-Si rich particles case, strong presence of Na and Cl
536 could indicate internal mixing with some sea spray aerosol, whereas a strong S presence could indicate
537 atmospheric acid ageing.

538 Some of these categories could be further grouped. For example, the particles in the Ca rich, Al-Si
539 rich, Si only and Si rich categories could be considered as "mineral dust". However, if the sample
540 contains ash from combustion processes or volcanic origin, it will also appear in these last categories
541 since its composition is similar to mineral dust (Chen et al., 2012; Nakagawa and Ohba, 2003).

542 **8. Application to a sample collected from the atmosphere above S.E. England**

543 The SEM technique described in Sect. 4 has been applied to samples collected from the FAAM aircraft
544 in various locations. Here we show an example of some of the capabilities of the developed technique
545 applied to one of the samples collected in S. E. England; the resulting size resolved composition is
546 shown in Fig. 11. The fraction of particles corresponding to each compositional category described in
547 Sect. 7 for each size can be seen in Fig. 11a. The SEM size distribution of each composition category
548 can be seen in Fig. 11b. By looking at this analysis, one can see that the sample was clearly dominated
549 by carbonaceous aerosol particles in all the sizes, but there was a clear mineral dust mode (Si only, Si
550 rich Al-Si rich and Ca rich) and some smaller contributions of other aerosol types (metal rich and S
551 rich). A potentially useful application of the size resolved composition is calculating the surface area
552 or mass of an individual component of a heterogeneous aerosol. As an example, we have grouped the
553 mineral dust categories Si only, Si rich Al-Si rich and Ca rich to produce the surface area size
554 distribution of mineral dust (and potentially ash) in Fig. 11c.

555 There are very few ways to obtain the size-resolved composition of an aerosol sample. Single particle
556 laser based mass spectroscopy has been used in order to obtain the size-resolved composition of
557 aerosol samples, both on the ground and in an aircraft (Pratt and Prather, 2012), (Wendisch and
558 Brenguier, 2013). Examples of this are instruments like the Particle Analysis by Laser Mass



559 Spectrometer (PALMS) (Cziczo et al., 2006; Thomson et al., 2000), the Aerosol Time-Of-Flight Mass
560 Spectrometer (ATOFMS) (Pratt et al., 2009), or the Aircraft-Based Laser Ablation Aerosol Mass
561 Spectrometer (ALABAMA) (Brands et al., 2011). However, a limitation of these techniques is that they
562 focus on the fine mode, with limited information about the coarse mode. A size-resolved
563 compositional analysis as the one described here is able to obtain the accumulation and coarse mode
564 size distribution of mineral dust in most aerosol samples, even in the ones dominated by other species,
565 as shown in Fig. 11. In addition, this approach can be used to get size-dependent information about
566 the shape factor and other morphological properties of the mineral dust, as well as ratios in between
567 the element concentrations.

568 9. Conclusions

569 In this work we have characterised the filter inlet system on board the BAe-146 UK FAAM research
570 aircraft which is used for the collection of atmospheric aerosol particles for off line analysis. Our
571 primary goal is to use this inlet system for quantification of INP concentrations and size resolved
572 composition measurements, but it could also be used to derive other quantities with other analytical
573 techniques.

574 In order to characterise the inlet system we made use of an electron microscope technique to study
575 the inlet efficiency, by comparing the SEM size distributions with the in situ size distributions
576 measured with underwing optical probes (PCASP-CDP). The sub-isokinetic enhancement of large
577 aerosol particles predicted by the calculations in Sect. 2.2 was observed in these comparisons. We also
578 experimentally verify that this enhancement is minimised by operating the inlet with the bypass open
579 which maximised the flow rate through the inlet nozzle. In addition, we note that we performed tests
580 with three very different aerosol distributions and the size distribution of the particles on the filters
581 was comparable to those measured by the underwing optical probes. Overall, the inlet tends to
582 enhance the concentration of aerosol in the coarse mode with a peak enhancement at $\sim 10 \mu\text{m}$. The
583 inlet efficiency decreases rapidly for sizes above about $20 \mu\text{m}$ and becomes highly dependent upon
584 the specifics of the sampling such as flow rates and angle of attack. Based on these tests we
585 recommend that the total flow rates at the nozzle are maintained at between 50 and 80 l min^{-1} , and
586 also that improvements are made to the pump and bypass flow control (see Sect. 2.3).

587 We also established an SEM technique to determine the size resolved composition of the aerosol
588 sample. Each particle can be categorized based on its chemical composition using a custom made
589 classification scheme. Using this technique we showed that the filter system on board of the BAe-146
590 spreads the particles evenly across the filter surface, which is necessary for the SEM size distribution
591 analysis.

592 Having a well characterised inlet allows us to sample aerosol particles up to around $20 \mu\text{m}$ with
593 knowledge of the likely biases from the aircraft. Hence, we can use this inlet system to collect aerosol
594 for offline analysis at altitudes which are relevant for clouds. For example, this may allow us to use the
595 size resolved aerosol composition to quantify the size distribution of individual aerosol components
596 at a particular location and combine this information with INP measurements to quantify the surface
597 area normalised ice nucleating ability of a specific class of aerosol.

598



599 **Appendix A: discussion of the inlet efficiency calculations**

600 Here we include a further description of the efficiency mechanisms used in the inlet model described
601 in Fig. 2 and discuss the choice of the equations and their limits of validity:

602 *Aspiration efficiency* accounts for the fact that the speed of the sampled air mass (U_0), and the speed
603 of the air through the beginning of the nozzle (U) are different. When these two speeds are equal, the
604 sampling is called “isokinetic”, whereas when the speeds don’t match, the sampling is called super
605 isokinetic or sub-isokinetic depending on if U_0 is smaller or larger than U respectively. In our case, the
606 air mass moves at the flying speed, which varies with the altitude (110 m s^{-1} is a typical value for
607 sampling altitudes), and the speed at the start of the inlet is almost always below 35 m s^{-1} (sub-
608 isokinetic conditions). As a consequence, some air streamlines will be forced around the inlet, while
609 high inertia particles won’t, which will lead to an aspiration efficiency above 1 for coarse mode aerosol
610 particles. This enhancement is greater for large particles due to their large inertia which makes difficult
611 their ability to follow the air streamlines. The enhancement reaches a maximum value of U_0/U in its
612 high diameter limit (when none of the particles in the sampled air mass follow the streamlines that
613 escape from the inlet and all of them are sampled). The aspiration efficiency tends to 1 (no
614 enhancement) for small diameters.

615 This behaviour has been characterised by several studies (we will only look at the sub-isokinetic range
616 of the equations since it is impossible to reach the super isokinetic range during flight). An empirical
617 equation was developed based on laboratory experiment by Belyaev and Levin (1972) and Belyaev
618 and Levin (1974) (referred as B&L) for certain range of U/U_0 ratio and Stokes number. However, for
619 ratios below its experimental range ($U/U_0 > 0.2$), the B&L function doesn’t make physical sense since
620 it converges to values above 1 for small particle sizes. The aircraft inlet system works at smaller U/U_0
621 ratios sometimes, so this function is not very accurate to describe the behaviour of the system in such
622 conditions. Liu et al. (1989) developed another function (referred as LZK) by means of a numerical
623 simulation based on computational fluid mechanics. The U/U_0 ratio and Stokes number valid range is
624 wider than the B&L expression (down to 0.1). It agrees with the B&L expression in the U/U_0 ratio the
625 latter was developed for. For smaller values of the ratio, the LZK function are believed to be more
626 accurate, since it predicts the known physical behaviour (no sub-isokinetic enhancement for small
627 particle sizes). It reaches U/U_0 ratios down to 0.2, which is enough to cover most of the total flow
628 rates achieved in the inlet system. (Krämer and Afchine, 2004) developed another expression (referred
629 as K&A) for $0.007 < U/U_0 < 0.2$ based on computational fluid dynamics. However, for low particle sizes,
630 the efficiency doesn’t converge to 1. As a consequence, we have used the LZK (Liu et al., 1989) function
631 since it covers most of the U/U_0 ratios we get in the inlet system, it agrees with the experimental data
632 in Belyaev and Levin (1972) and Belyaev and Levin (1974) and it converges to U_0/U for large particles
633 sizes and 1 for small particle sizes. Outside its valid range ($U/U_0 < 0.1$), the LZK function agrees with
634 the K&A function for large radius and converges to 1 for small particle sizes. The equation is valid for
635 $0.01 < \text{Stks} < 100$, which is enough to cover the range in between 1 and $100 \mu\text{m}$. As already stated, it
636 tends to 1 for small particles sizes and to U_0/U for large particles sizes (At 50 L min^{-1} , the ratio U/U_0 is
637 0.2). All the calculations were done under standard conditions (0°C and 1 bar).

638 *Inlet inertial deposition* is defined as the inertial loss of aerosol particles when they enter nozzle. It is
639 produced by the fact that the streamlines bend towards the walls at the moment they enter the
640 nozzle, some large inertia particles can impact the walls and get deposited. Here, we have used the
641 equation given in Liu et al. (1989) which quantifies this effect. It is also valid for $0.01 < \text{Stks} < 100$, which
642 is enough to cover the range in between 1 and $100 \mu\text{m}$.



643 *Turbulent inertial deposition* happens when some particles are collected by the wall when travelling
644 in a pipe in the turbulent regime because some of the particles cannot follow the eddies of the
645 turbulent flow. In order to include this mechanism, we used the equation given in Brockmann (2011),
646 using the relation in between the deposition velocity and dimensionless particle relaxation time given
647 by Liu and Ilori (1974). These calculations are valid for a cylindrical pipe, whereas the turbulent section
648 of the inlet considered here is the nozzle, which has a conical shape. In order to account for this, we
649 divided the conical nozzle into 90 conical sections with an increasing diameter and a length of 1mm,
650 and combined the effect of all the sections. As already mentioned, above 65 L min^{-1} , turbulent flow
651 occurs in the whole inlet tube. This has been taken into account in the 80 L min^{-1} case in Fig. 2b. The
652 equation used here has been tested for size ranges in between 1.4 and $20 \mu\text{m}$, and doesn't depend on
the Reynolds number values it was tested for (10000 and 50000) (Liu and Ilori, 1974).

654 *Bending inertial deposition* was also considered, since the line curves with an angle of 45° in order to
655 bring the airstream into the cabin. The inertia of some particles may keep them in their original track
656 and they are not able to follow the air streamlines that are bending towards the cabin, following the
657 inlet tubes. In order to account for these losses, we have used the empirical equation given in
658 Brockmann (2011) based on the data from Pui et al. (1987) for laminar flow. This equation was
659 developed for Reynold numbers of 1000, and we have used it for higher values. However, in
660 Brockmann (2011), one can see that the data from Pui et al. (1987) for $\text{Re}=6000$ (beginning of the
661 turbulent flow regime) doesn't differ that much from the fit we have used (valid for $\text{Re}=1000$). Since
662 our Re numbers for the thick section of the tube almost never go above 5000, we can still use the
663 laminar flow fit. This model has been tested for $0.08 < \text{Stks} < 1.2$, which is enough to cover most of the
664 range where the inertial deposition efficiency drops from 1 to 0. The main caveat of this calculation is
665 that the model considers that the flow rate before and after the bending is the same, while in the inlet
666 system, if the bypass flow is on, the flow rate before and after the bending is different (before it,
667 it would be equal to the total flow rate, whereas after the bending, it would be equal to the filter flow
668 rate). As a consequence we assumed that the flow rate after the bending is equal to the total flow
669 rate.

670 *Gravitational settling* was also considered. We used the analytical equation given by Thomas (1958),
671 as stated in Brockmann (2011). We applied this equation for the section of the pipe from the nozzle
672 to the bend (15 cm long). We used the modification (also analytical) of the previous equation given in
673 Heyder and Gebhart (1977) in order to account for the losses in the second section of the tube which
674 is 40 cm long and it is bended 45° . The gravitational losses in the nozzle were neglected since the
675 settling distance is much shorter and the time the air takes to pass it is smaller since it travels quicker.
676 As stated previously, the lower part of the turbulent regime can be reached for high flow rates through
677 all the tube. For these cases, we still use this equation which is only valid for the laminar regime, since
678 the gravitational settling efficiencies for the turbulent regime are very close to the laminar regime
679 ones (Brockmann, 2011) and wouldn't make a significant difference in our calculations.

680 *Diffusional efficiency* accounts for the fact that small aerosol particles could diffuse to the walls of the
681 pipe via Brownian motion. In order to account for this phenomenon, we have used the analytic
682 equation by Gormley and Kennedy (1948) as stated in Brockmann (2011). We have assumed that
683 diffusion happens only in the tube (before and after the bend) and excluding the diffusion in the nozzle
684 since it is negligible because these losses are a function of the residence time and the residence time
685 of the aerosol particles in the nozzle is much smaller than the rest of the tube. For this calculation,
686 we have assumed 0°C and 1 atm. We didn't show the efficiency associated to diffusion in Fig. 2a
687 because it was very close to 1 for all considered sizes. It only becomes slightly smaller than 1 for sizes
688 below 20 nm at 50 L min^{-1} . As a consequence, the inlet could be potentially used to sample nucleation



689 mode aerosol particles, even though for this study we will only focus on the particles larger than 0.1
690 μm .

691 *Filter collection efficiency* accounts for the fact that some particles can pass through the pores of the
692 filter, if they are smaller than the pores. However, filter pore size (in the case of polycarbonate
693 capillarity filters) and filter equivalent pore size (in the case of PTFE porous filters) is sometimes
694 misunderstood as a size cut off at which smaller particles are lost and larger particles are captured.
695 However, particle collection on filters happens through several mechanisms including interception,
696 impaction, diffusion, gravitational settling or by electrostatic attraction under certain conditions
697 (Flagan and Seinfeld, 1988; Lee and Ramamurthi, 1993). As a consequence, particles with diameters
698 below the pore size are normally collected (Lindsley, 2016; Soo et al., 2016). 99.48% of the generated
699 sodium chloride particles with sizes in between 10.4 and 412 nm were collected by a 0.4 μm
700 polycarbonate filter at flow rates below 11.2 L min^{-1} (smaller than most of the flow rates at which the
701 air passes through the same filters in the FAAM filter inlet system) (Soo et al., 2016). As a consequence,
702 we assumed a filter collection efficiency of 100% across the whole considered size range (0.1 to 100
703 μm).

704 *Anisoaxial losses* have not been considered in the analysis shown in Fig. 2, after estimating that they
705 would only affect particles significantly larger than 10 μm and the fact that the alignment of the inlet
706 is difficult to quantify and the angle of attack changes during the flight. Using the equations explained
707 in Hangal and Willeke (1990a), we calculated that the modification of the sub-isokinetic behaviour of
708 the inlet produced by small values of θ is negligible. The equation was used beyond its experimental
709 limit, but this extrapolation was justified by the fact that the equation for $\theta = 0$ made asymptotic
710 physical sense at the low and high Stokes number limits and produced very similar results to the ones
711 showed in Fig. 2a. Anisoaxial sampling can also produce inertial losses when particles impact the walls
712 of the inlet. These ones have been quantified using the expression given by Hangal and Willeke
713 (1990b) for different values of θ and they can be seen in Fig. 3. This mechanism looks very similar to
714 the gravitational and bend deposition efficiency shown in Fig. 2a. Anisoaxial inertial losses add a cut
715 off that prevents large particles to be sampled. As one can see in Fig. 3, the effect is very dependent
716 on the angle and only affects particles significantly larger than 10 μm in most cases, so it hasn't been
717 included in the total analysis shown in the Fig 2. One can see in Fig. 3 that the position of the D50 of
718 the anisoaxial cut off decreases when increasing values of θ up to 2°. For values of θ between 2° and
719 6°, it increases when increasing θ .

720

721 Acknowledgements

722 We are grateful for helpful discussions with Paola Formenti (Centre National de la Recherche
723 Scientifique), Hannah Price (FAAM) and Eduardo Morgado (School of Earth and Environment,
724 University of Leeds). Airborne data were obtained using the BAe-146-301 Atmospheric Research
725 Aircraft operated by Directflight Ltd (now Airtask) and managed by FAAM, which is a joint entity of the
726 Natural Environment Research Council (NERC) and the UK Met Office. We acknowledge the dedicated
727 work of FAAM, Directflight and Avalon. We thank all the people involved in the EMERGE, VANAHEIM
728 and MACSSIMIZE campaigns. We acknowledge the Leeds Electron Microscopy and Spectroscopy
729 Centre for the use of their microscopy facilities and the Centre for Environmental Data Analysis for
730 providing us with the FAAM datasets used here. We are grateful for funding from the European
731 Research Council (ERC) (648661 MarineIce).

732

733 **References**

- 734 Albrecht, B. A.: Aerosols, cloud microphysics, and fractional cloudiness, *Science*, 245, 1227-1230, 1989.
- 735 Andreae, M. O., Berresheim, H., Andreae, T. W., Kritz, M. A., Bates, T. S., and Merrill, J. T.: Vertical-
736 Distribution of Dimethylsulfide, Sulfur-Dioxide, Aerosol Ions, and Radon over the Northeast Pacific-
737 Ocean, *J. Atmos. Chem.*, 6, 149-173, 1988.
- 738 Andreae, M. O., Elbert, W., Gabriel, R., Johnson, D. W., Osborne, S., and Wood, R.: Soluble ion
739 chemistry of the atmospheric aerosol and SO₂ concentrations over the eastern North Atlantic during
740 ACE-2, *Tellus B*, 52, 1066-1087, 2000.
- 741 Angyal, A., Kertész, Z., Szikszai, Z., and Szoboszlai, Z.: Study of Cl-containing urban aerosol particles by
742 ion beam analytical methods, *Nucl. Instrum. Methods Phys. Res. B*, 268, 2211-2215, 2010.
- 743 Artaxo, P. and Hansson, H. C.: Size Distribution of Biogenic Aerosol-Particles from the Amazon Basin,
744 *Atmos. Environ.*, 29, 393-402, 1995.
- 745 Baumgardner, D., Brenguier, J. L., Bucholtz, A., Coe, H., DeMott, P., Garrett, T. J., Gayet, J. F., Hermann,
746 M., Heymsfield, A., Korolev, A., Kramer, M., Petzold, A., Strapp, W., Pilewskie, P., Taylor, J., Twohy, C.,
747 Wendisch, M., Bachalo, W., and Chuang, P.: Airborne instruments to measure atmospheric aerosol
748 particles, clouds and radiation: A cook's tour of mature and emerging technology, *AtmRe*, 102, 10-29,
749 2011.
- 750 Baumgardner, D. and Huebert, B.: The airborne aerosol inlet workshop: Meeting report, *J. Aerosol Sci*,
751 24, 835-846, 1993.
- 752 Belyaev, S. P. and Levin, L. M.: Investigation of aerosol aspiration by photographing particle tracks
753 under flash illumination, *J. Aerosol Sci*, 3, 127-140, 1972.
- 754 Belyaev, S. P. and Levin, L. M.: Techniques for collection of representative aerosol samples, *J. Aerosol*
755 *Sci*, 5, 325-338, 1974.
- 756 Bergin, M. H., Ogren, J. A., Schwartz, S. E., and McInnes, L. M.: Evaporation of ammonium nitrate
757 aerosol in a heated nephelometer: Implications for field measurements, *Environ. Sci. Technol.*, 31,
758 2878-2883, 1997.
- 759 Brands, M., Kamphus, M., Böttger, T., Schneider, J., Drewnick, F., Roth, A., Curtius, J., Voigt, C., Borbon,
760 A., Beekmann, M., Bourdon, A., Perrin, T., and Borrmann, S.: Characterization of a Newly Developed
761 Aircraft-Based Laser Ablation Aerosol Mass Spectrometer (ALABAMA) and First Field Deployment in
762 Urban Pollution Plumes over Paris During MEGAPOLI 2009, *Aerosol Sci. Technol.*, 45, 46-64, 2011.
- 763 Brockmann, J. E.: Aerosol Transport in Sampling Lines and Inlets. In: *Aerosol Measurement*, John Wiley
764 & Sons, Inc., 2011.
- 765 Buckle, R., Tsakiroopoulos, P., and C. Pointon, K.: Preparation and properties of metallic aerosol
766 particles, *Int. Mater. Rev.*, 31, 258-288, 1986.
- 767 Chen, H., Laskin, A., Baltrusaitis, J., Gorski, C. A., Scherer, M. M., and Grassian, V. H.: Coal fly ash as a
768 source of iron in atmospheric dust, *Environ. Sci. Technol.*, 46, 2112-2120, 2012.
- 769 Chou, C., Formenti, P., Maille, M., Ausset, P., Helas, G., Harrison, M., and Osborne, S.: Size distribution,
770 shape, and composition of mineral dust aerosols collected during the African Monsoon
771 Multidisciplinary Analysis Special Observation Period 0: Dust and Biomass-Burning Experiment field
772 campaign in Niger, January 2006, *J. Geophys. Res. Atmos.*, 113, 1-17, 2008.
- 773 Cochran, R. E., Ryder, O. S., Grassian, V. H., and Prather, K. A.: Sea Spray Aerosol: The Chemical Link
774 between the Oceans, Atmosphere, and Climate, *Acc. Chem. Res.*, 50, 599-604, 2017.



- 775 Crespo, J., Yubero, E., Nicolas, J. F., Lucarelli, F., Nava, S., Chiari, M., and Calzolari, G.: High-time
776 resolution and size-segregated elemental composition in high-intensity pyrotechnic exposures, *J.*
777 *Hazard. Mater.*, 241-242, 82-91, 2012.
- 778 Cziczo, D. J., Thomson, D. S., Thompson, T. L., DeMott, P. J., and Murphy, D. M.: Particle analysis by
779 laser mass spectrometry (PALMS) studies of ice nuclei and other low number density particles, *Int. J.*
780 *Mass spectrom.*, 258, 21-29, 2006.
- 781 DeMott, P. J., Prenni, A. J., Liu, X., Kreidenweis, S. M., Petters, M. D., Twohy, C. H., Richardson, M. S.,
782 Eidhammer, T., and Rogers, D. C.: Predicting global atmospheric ice nuclei distributions and their
783 impacts on climate, *Proc Natl Acad Sci U S A*, 107, 11217-11222, 2010.
- 784 Egerton, R. F.: *Physical principles of electron microscopy*, Springer US, 2005.
- 785 Flagan, R. C. and Seinfeld, J. H.: Removal of particles from gas streams. In: *Fundamentals of Air*
786 *Pollution Engineering*, Prentice-Hall, Inc., Englewood Cliffs, NJ, 1988.
- 787 Fomba, K. W., van Pinxteren, D., Muller, K., Iinuma, Y., Lee, T., Collett, J. L., and Herrmann, H.: Trace
788 metal characterization of aerosol particles and cloud water during HCCT 2010, *Atmos. Chem. Phys.*, 15,
789 8751-8765, 2015.
- 790 Gao, Y., Anderson, J. R., and Hua, X.: Dust characteristics over the North Pacific observed through
791 shipboard measurements during the ACE-Asia experiment, *Atmos. Environ.*, 41, 7907-7922, 2007.
- 792 Gibson, E. R., Hudson, P. K., and Grassian, V. H.: Aerosol chemistry and climate: Laboratory studies of
793 the carbonate component of mineral dust and its reaction products, *Geophys Res Lett*, 33, 1-5, 2006.
- 794 Gormley, P. G. and Kennedy, M.: Diffusion from a Stream Flowing through a Cylindrical Tube, *Math.*
795 *Proc. Royal Ir. Acad.*, 52, 163-169, 1948.
- 796 Graedel, T. E. and Keene, W. C.: Tropospheric budget of reactive chlorine, *Global Biogeochem. Cycles*,
797 9, 47-77, 1995.
- 798 Hand, V. L., Capes, G., Vaughan, D. J., Formenti, P., Haywood, J. M., and Coe, H.: Evidence of internal
799 mixing of African dust and biomass burning particles by individual particle analysis using electron
800 beam techniques, *J. Geophys. Res.*, 115, 1-11, 2010.
- 801 Hangal, S. and Willeke, K.: Aspiration efficiency: unified model for all forward sampling angles, *Environ.*
802 *Sci. Technol.*, 24, 688-691, 1990a.
- 803 Hangal, S. and Willeke, K.: Overall Efficiency of Tubular Inlets Sampling at 0-90 Degrees from Horizontal
804 Aerosol Flows, *Atmos Environ a-Gen*, 24, 2379-2386, 1990b.
- 805 Haywood, J. and Boucher, O.: Estimates of the direct and indirect radiative forcing due to tropospheric
806 aerosols: A review, *Rev Geophys*, 38, 513-543, 2000.
- 807 Heyder, J. and Gebhart, J.: Gravitational deposition of particles from laminar aerosol flow through
808 inclined circular tubes, *J. Aerosol Sci*, 8, 289-295, 1977.
- 809 Hoose, C. and Mohler, O.: Heterogeneous ice nucleation on atmospheric aerosols: a review of results
810 from laboratory experiments, *Atmos. Chem. Phys.*, 12, 9817-9854, 2012.
- 811 Hyuk Kim, C.: Characterization of Volatilization of Filter-Sampled PM_{2.5} Semi-Volatile Inorganic Ions
812 Using a Backup Filter and Denuders, *Aerosol Air Qual. Res.*, 2015, 2015.
- 813 Kandler, K., Benker, N., Bundke, U., Cuevas, E., Ebert, M., Knippertz, P., Rodríguez, S., Schütz, L., and
814 Weinbruch, S.: Chemical composition and complex refractive index of Saharan Mineral Dust at Izaña,
815 Tenerife (Spain) derived by electron microscopy, *Atmos. Environ.*, 41, 8058-8074, 2007.



- 816 Kandler, K., Lieke, K., Benker, N., Emmel, C., Kupper, M., Muller-Ebert, D., Ebert, M., Scheuven, D.,
817 Schladitz, A., Schutz, L., and Weinbruch, S.: Electron microscopy of particles collected at Praia, Cape
818 Verde, during the Saharan Mineral Dust Experiment: particle chemistry, shape, mixing state and
819 complex refractive index, *Tellus B*, 63, 475-496, 2011.
- 820 Kim, J., Bauer, H., Dobovicnik, T., Hitzenberger, R., Lottin, D., Ferry, D., and Petzold, A.: Assessing
821 Optical Properties and Refractive Index of Combustion Aerosol Particles Through Combined
822 Experimental and Modeling Studies, *Aerosol Sci. Technol.*, 49, 340-350, 2015.
- 823 Krämer, M. and Afchine, A.: Sampling characteristics of inlets operated at low U/U₀ ratios: new
824 insights from computational fluid dynamics (CFX) modeling, *J. Aerosol Sci*, 35, 683-694, 2004.
- 825 Krejci, R., Strom, J., de Reus, M., and Sahle, W.: Single particle analysis of the accumulation mode
826 aerosol over the northeast Amazonian tropical rain forest, Surinam, South America, *Atmos. Chem.*
827 *Phys*, 5, 3331-3344, 2005.
- 828 Kumar, M. and Francisco, J. S.: Elemental sulfur aerosol-forming mechanism, *Proc. Natl. Acad. Sci.*
829 *U.S.A.*, 114, 864-869, 2017.
- 830 Laskin, A. and Cowin, J. P.: Automated single-particle SEM/EDX analysis of submicrometer particles
831 down to 0.1 μm , *Anal Chem*, 73, 1023-1029, 2001.
- 832 Lee, K. W. and Ramamurthi, M.: Chapter 10: Filter Collection. In: *Aerosol Measurement: Principles,*
833 *Techniques, and Applications*
- 834 Willeke, K. and Baron, P. (Eds.), *Van Nostrand Reinhold*, New York, 1993.
- 835 Lindsley, W. G.: Filter Pore Size and Aerosol Sample Collection, *NIOSH Manual of Analytical Methods*,
836 2016. 1-14, 2016.
- 837 Liu, B. Y. H. and Ilori, T. A.: Aerosol deposition in turbulent pipe flow, *Environ. Sci. Technol.*, 8, 351-
838 356, 1974.
- 839 Liu, B. Y. H., Zhang, Z. Q., and Kuehn, T. H.: A Numerical Study of Inertial Errors in Anisokinetic
840 Sampling, *J. Aerosol Sci*, 20, 367-380, 1989.
- 841 Lohmann, U. and Diehl, K.: Sensitivity Studies of the Importance of Dust Ice Nuclei for the Indirect
842 Aerosol Effect on Stratiform Mixed-Phase Clouds, *J. Atmospheric Sci.*, 63, 968-982, 2006.
- 843 Lohmann, U. and Gasparini, B.: A cirrus cloud climate dial?, *Science*, 357, 248-249, 2017.
- 844 Mason, R. H., Si, M., Chou, C., Irish, V. E., Dickie, R., Elizondo, P., Wong, R., Brintnell, M., Elsasser, M.,
845 Lassar, W. M., Pierce, K. M., Leaitch, W. R., MacDonald, A. M., Platt, A., Toom-Sauntry, D., Sarda-
846 Esteve, R., Schiller, C. L., Suski, K. J., Hill, T. C. J., Abbatt, J. P. D., Huffman, J. A., DeMott, P. J., and
847 Bertram, A. K.: Size-resolved measurements of ice-nucleating particles at six locations in North
848 America and one in Europe, *Atmos. Chem. Phys*, 16, 1637-1651, 2016.
- 849 McConnell, C. L., Formenti, P., Highwood, E. J., and Harrison, M. A. J.: Using aircraft measurements to
850 determine the refractive index of Saharan dust during the DODO Experiments, *Atmos. Chem. Phys*,
851 10, 3081-3098, 2010.
- 852 McMurry, P.: A review of atmospheric aerosol measurements, *Atmos. Environ.*, 34, 1959-1999, 2000.
- 853 Nakagawa, M. and Ohba, T.: Minerals in Volcanic Ash 1: Primary Minerals and Volcanic Glass, *Global*
854 *Environ. Res.*, 6, 2003.



- 855 Nessler, R., Bukowiecki, N., Henning, S., Weingartner, E., Calpini, B., and Baltensperger, U.:
856 Simultaneous dry and ambient measurements of aerosol size distributions at the Jungfraujoch, Tellus
857 B, 55, 808-819, 2003.
- 858 Paciga, J. J., Roberts, T. M., and Jervis, R. E.: Particle-Size Distributions of Lead, Bromine, and Chlorine
859 in Urban-Industrial Aerosols, Environ. Sci. Technol., 9, 1141-1144, 1975.
- 860 Pratt, K. A., Mayer, J. E., Holecek, J. C., Moffet, R. C., Sanchez, R. O., Rebotier, T. P., Furutani, H., Gonin,
861 M., Fuhrer, K., Su, Y., Guazzotti, S., and Prather, K. A.: Development and characterization of an aircraft
862 aerosol time-of-flight mass spectrometer, Analytical Chemistry, 81, 1792-1800, 2009.
- 863 Pratt, K. A. and Prather, K. A.: Mass spectrometry of atmospheric aerosols--recent developments and
864 applications. Part II: On-line mass spectrometry techniques, Mass Spectrom. Rev., 31, 17-48, 2012.
- 865 Price, H. C., Baustian, K. J., McQuaid, J. B., Blyth, A., Bower, K. N., Choulaton, T., Cotton, R. J., Cui, Z.,
866 Field, P. R., Gallagher, M., Hawker, R., Merrington, A., Miltenberger, A., Neely Iii, R. R., Parker, S. T.,
867 Rosenberg, P. D., Taylor, J. W., Trembath, J., Vergara-Temprado, J., Whale, T. F., Wilson, T. W., Young,
868 G., and Murray, B. J.: Atmospheric Ice-Nucleating Particles in the Dusty Tropical Atlantic, J. Geophys.
869 Res. Atmos., 123, 2175-2193, 2018.
- 870 Pruppacher, H. R. and Klett, J. D.: Microphysics of Clouds and Precipitation Dordrecht: Kluwer
871 Academic Publishers, 1997.
- 872 Pui, D. Y. H., Romay-Novas, F., and Liu, B. Y. H.: Experimental Study of Particle Deposition in Bends of
873 Circular Cross Section, Aerosol Sci. Technol., 7, 301-315, 1987.
- 874 Reid, E. A., Reid, J. S., Meier, M. M., Dunlap, M. R., Cliff, S. S., Broumas, A., Perry, K., and Maring, H.:
875 Characterization of African dust transported to Puerto Rico by individual particle and size segregated
876 bulk analysis, J. Geophys. Res. Atmos., 108, -, 2003.
- 877 Rosenberg, P. D., Dean, A. R., Williams, P. I., Dorsey, J. R., Minikin, A., Pickering, M. A., and Petzold, A.:
878 Particle sizing calibration with refractive index correction for light scattering optical particle counters
879 and impacts upon PCASP and CDP data collected during the Fennec campaign, Atmos. Meas. Tech., 5,
880 1147-1163, 2012.
- 881 Seinfeld, J. H. and Pandis, S. N.: Atmospheric Chemistry and Physics: From Air Pollution to Climate
882 Change, Wiley, 2006.
- 883 Soo, J. C., Monaghan, K., Lee, T., Kashon, M., and Harper, M.: Air sampling filtration media: Collection
884 efficiency for respirable size-selective sampling, Aerosol Sci Technol, 50, 76-87, 2016.
- 885 Talbot, R. W., Andreae, M. O., Berresheim, H., Artaxo, P., Garstang, M., Harriss, R. C., Beecher, K. M.,
886 and Li, S. M.: Aerosol Chemistry during the Wet Season in Central Amazonia - the Influence of Long-
887 Range Transport, J. Geophys. Res. Atmos., 95, 16955-16969, 1990.
- 888 Thomas, J. W.: Gravity Settling of Particles in a Horizontal Tube, J. Air Pollut. Control Assoc., 8, 32-34,
889 1958.
- 890 Thomson, D. S., Schein, M. E., and Murphy, D. M.: Particle Analysis by Laser Mass Spectrometry WB-
891 57F Instrument Overview, Aerosol Sci. Technol., 33, 153-169, 2000.
- 892 von der Weiden, S. L., Drewnick, F., and Borrmann, S.: Particle Loss Calculator – a new software tool
893 for the assessment of the performance of aerosol inlet systems, Atmos. Meas. Tech., 2, 479-494, 2009.
- 894 Wendisch, M. and Brenguier, J. L.: Airborne Measurements for Environmental Research: Methods and
895 Instruments, 2013.

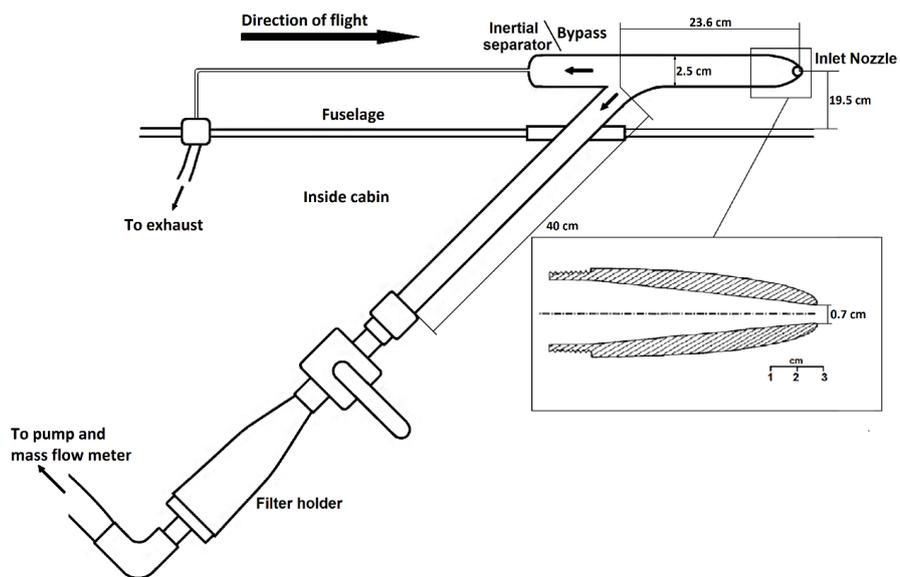


896 Young, G., Jones, H. M., Darbyshire, E., Baustian, K. J., McQuaid, J. B., Bower, K. N., Connolly, P. J.,
897 Gallagher, M. W., and Choularton, T. W.: Size-segregated compositional analysis of aerosol particles
898 collected in the European Arctic during the ACCACIA campaign, Atmos. Chem. Phys, 16, 4063-4079,
899 2016.

900

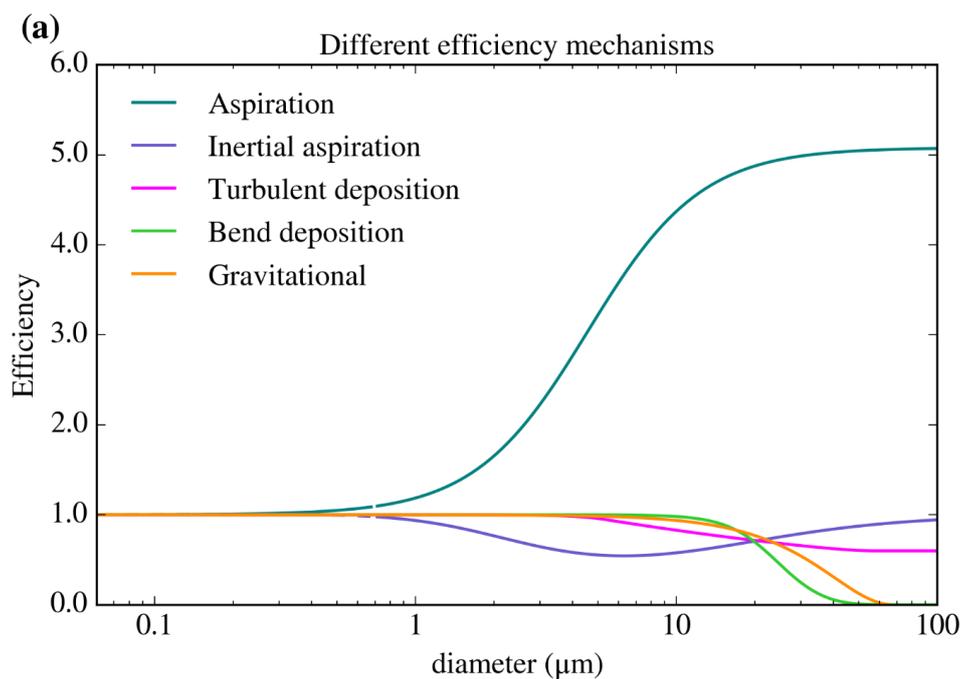


901 **Figures**

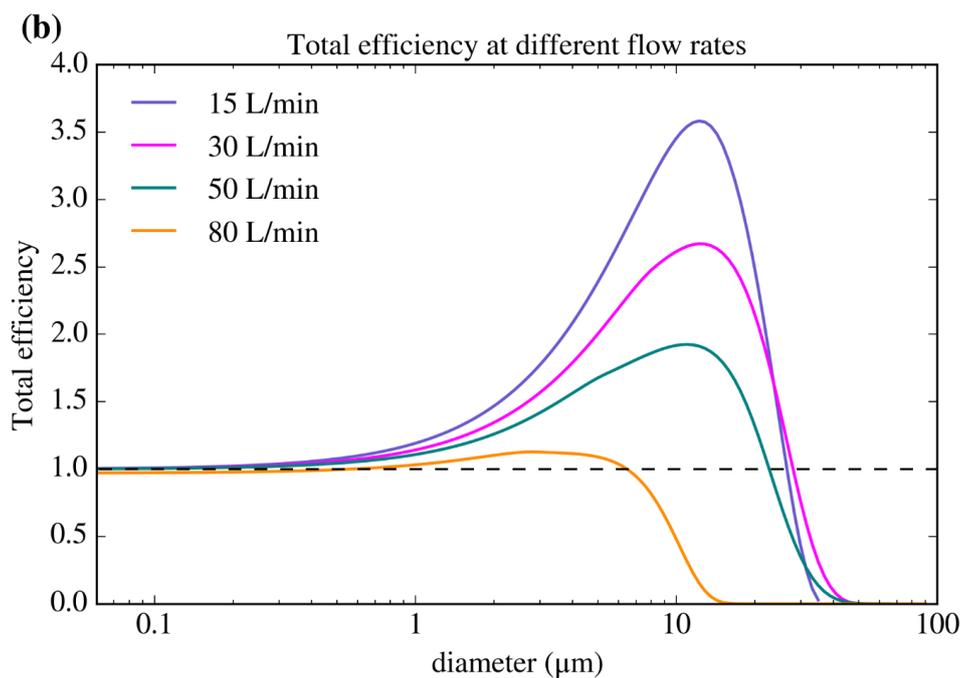


902

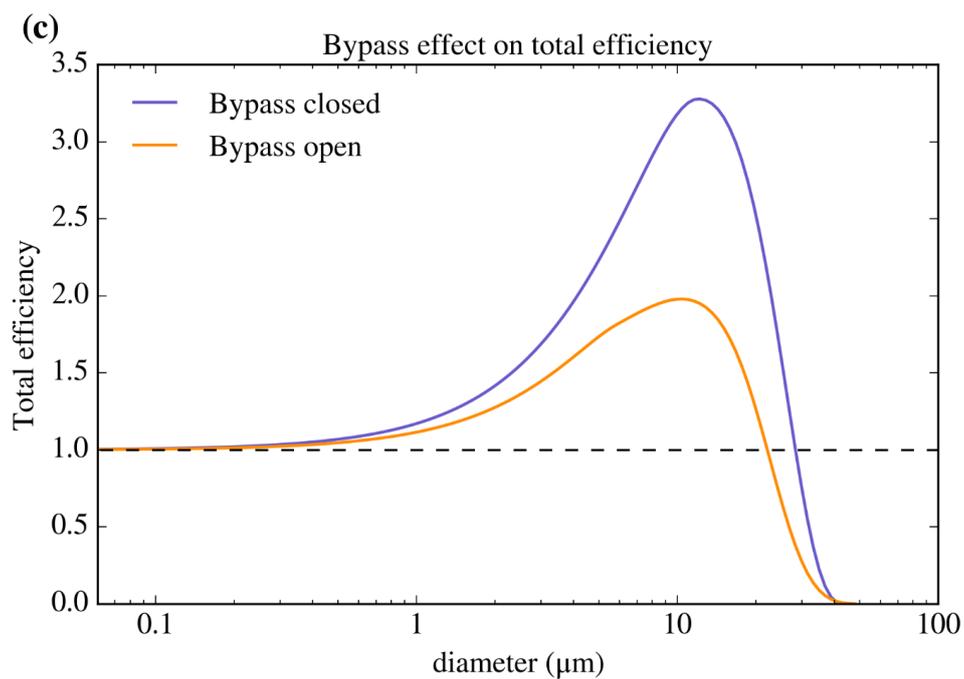
903 Fig. 1. Schematic diagram of one of the two parallel lines of the inlet system.



904



905



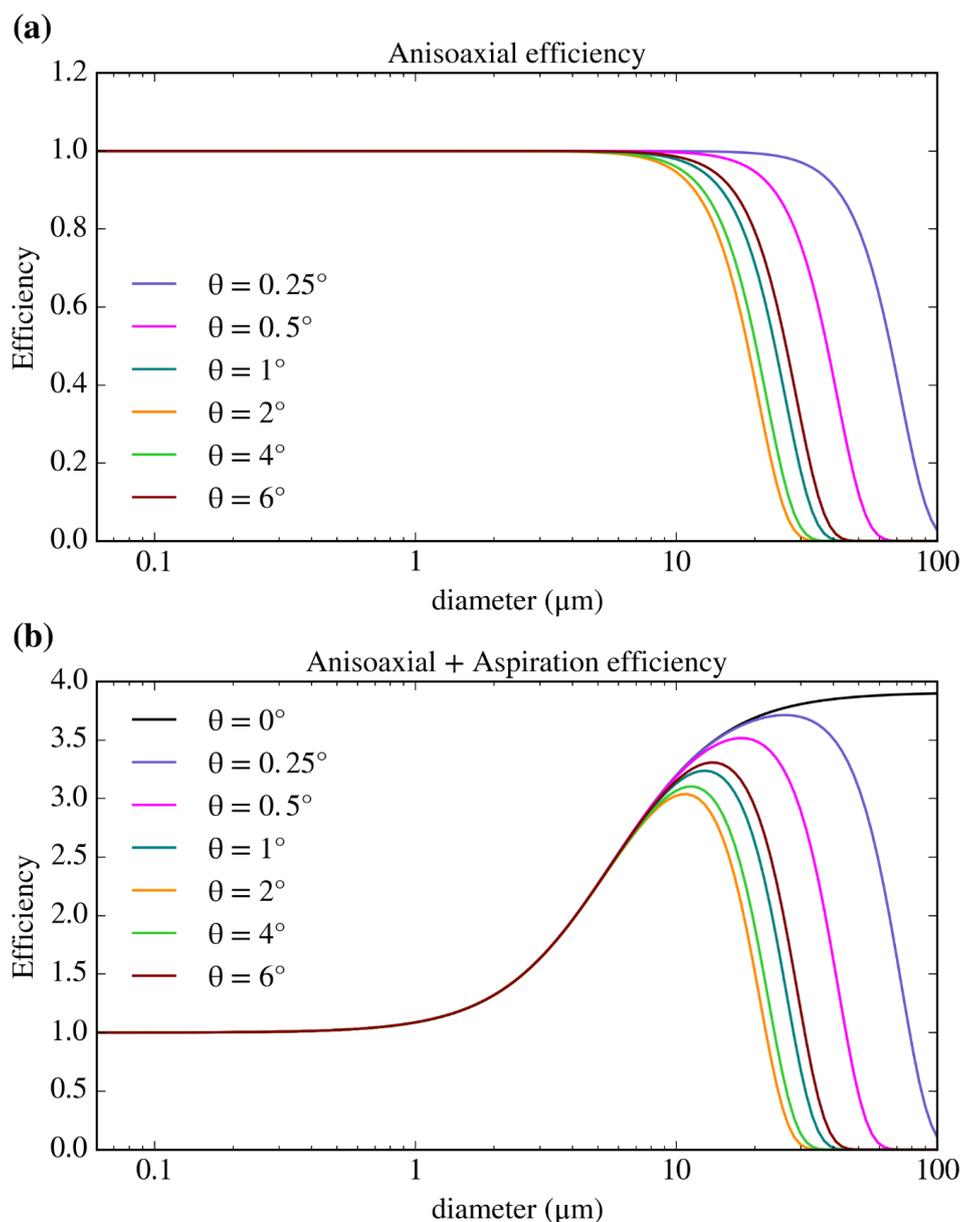
906

907 Fig. 2. *Theoretical efficiencies of the Filter inlet system.* (a) Efficiencies of the four mechanisms considered in this work for a
908 total flow rate of 50 L min^{-1} . We have assumed a dynamic viscosity of $1.82 \times 10^{-5} \text{ kg m}^{-1} \text{ s}^{-1}$ (value for 0°C) and a particle density
909 of 1000 kg m^{-3} . The speed of the air mass (U_0) was 110 m s^{-1} , a typical FAAM flying speed at low altitudes. (b) Total efficiency
910 for four different total flow rates. For the 80 L min^{-1} case, turbulent deposition through the whole line was considered since
911 the flow was turbulent through the whole pipe. (c) Total efficiency considering all the described mechanisms for a 20 L min^{-1}
912 filter flow rate with the bypass closed and a 20 L min^{-1} filter flow rate with the bypass open (considering a bypass flow of 25
913 L min^{-1}).

914



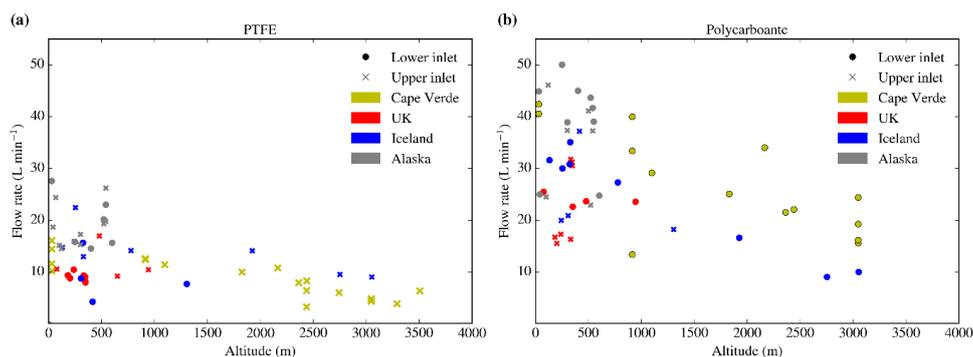
915



916
917 Fig. 3. Anisoaxial inertial losses of the sampling carried out by the Filters inlet system for different values of the angle in
918 between the inlet and the flight direction. The calculations have been presented by themselves (a) and combined with the
919 aspiration efficiency (b), which one can see in Fig. 2a. The anisoaxial calculations have been done using the equations given
920 by (Hangal and Willeke, 1990b), using the same parameters and dimensions than in Fig. 2, apart from the flow rate, which
921 was set to 65 L min^{-1} in order to be within the valid range of U/U_0 that was used to develop the equation. For smaller or
922 larger values of the flow rate (under which most of the sampling is carried out), the differences in the efficiency from the
923 ones show here are minimal.



924

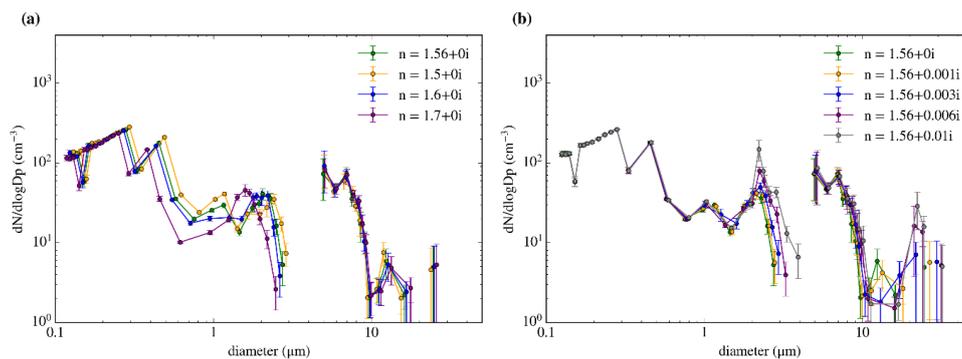


925

926

927 Figure 4. Filter flow rate of different samplings carried out in different campaigns at each altitude using: (a) Sartorius PTFE
928 membrane filters (47mm diameter with a pore size of 0.45 μm) and (b) Whatman nucleopore polycarbonate track etched
929 filters (47mm diameter with a pore size of 0.4 μm). The crosses represent samples taken in the upper line of the inlet system,
930 whereas dots represent the sampling in the bottom line. Different mesh supports were used for the data collection. The data
931 from Cape Verde was extracted from (Price et al., 2018) and the notes of the analysis carried out by the authors whereas the
932 altitude data from the other three was obtained from the pressure altitude measurement carried out by the Reduced Vertical
933 Separation Minimum system on board of the aircraft. The FAAM core datasets used (via the Centre for Environmental Data
934 Analysis) were C019, C022, C024, C025, C058, C059, C060, C061, C062, C063, C085, C086, C087, C088, C089, C090 and C091.
935 The bypass was closed for all the data in Cape Verde whereas it was open for all the data in the other campaigns. Note that
936 the flow rate here corresponds to the filter flow rate (measured with the mass flow meter), not the total one.

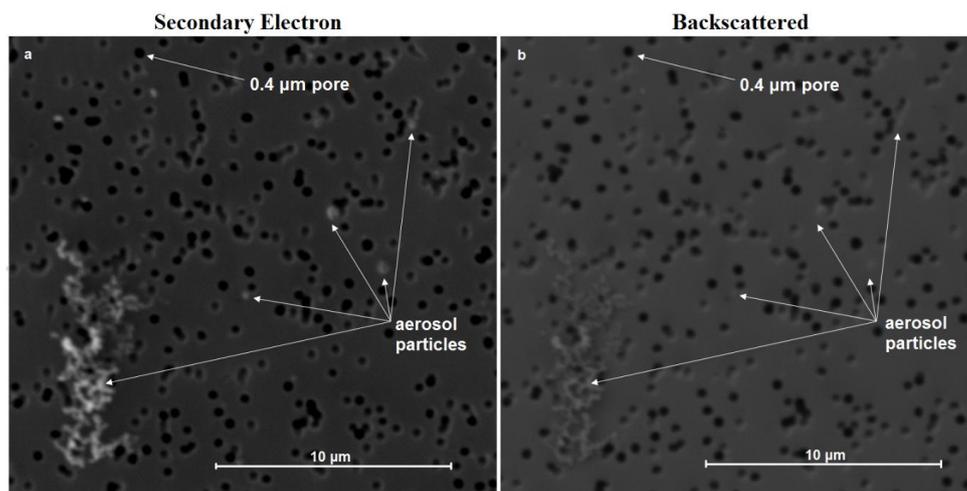
937



938

939 Figure 5. Sensitivity of the size distributions measured by the PCASP-CDP during the C010 flight on the 2017/05/10 from
940 11:24 to 11:38 UTC to small variations in the refractive index. We tested both the real part (a) and imaginary part (b). The
941 errors are calculated according to the methods explained in Sect. 3.

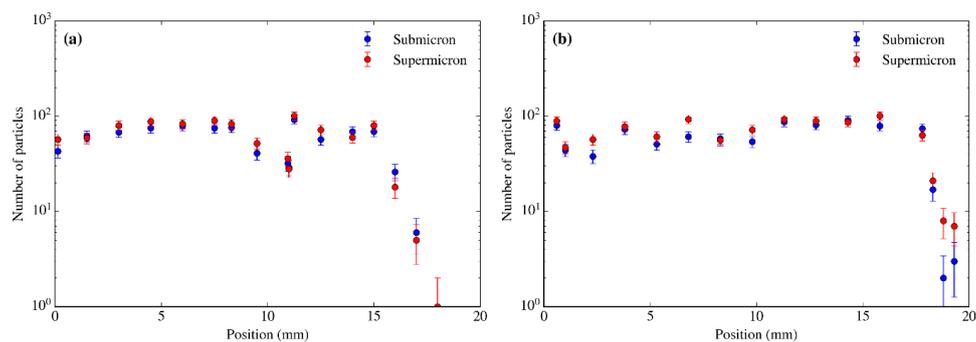
942



943

944 Figure 6. Secondary electron image (a) and Back Scattered Electron image (b) of the same area of the same filter, collected
945 on the 2018/07/05 from 13:32 to 13:47 in the upper line with the bypass open. As one can see, some of the small particles
946 in the SE image appear almost transparent under the BSE image. Even the 10μm soot particle in the bottom left of the
947 image shows a very low contrast in the BSE image.

948

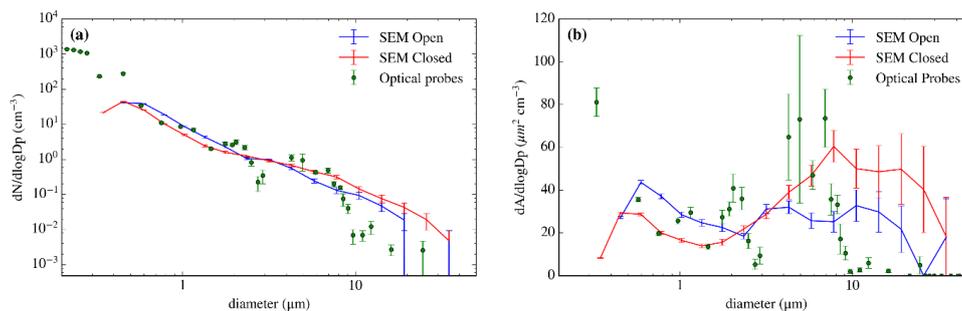


949

950 Figure 7. Radial distribution of particles test on the sample collected on the 2017/10/02 from 16:24 to 16:40 UTC about 320
951 m high in south Iceland, using the lower line and open bypass, sampling 219 L. Number of submicron and supermicron
952 particles in same size areas ($\sim 160 \times 190 \mu\text{m}^2$) radially distributed versus the distance from the approximate centre through a
953 radius of the filter (a) and another trajectory from the centre of the filter deviated 30° from the first radius (b). The analysis
954 was done at 20 KeV and $\times 5000$. The number of both supermicron and submicron particles remains very constant all over the
955 surface of the filter, until reaching the edges of it (which are cover by a rubber O-ring during the sample) and the number of
956 particles drops to the limit of the detection within a few millimetres. The error in the number of particles comes from Poisson
957 counting statistics.

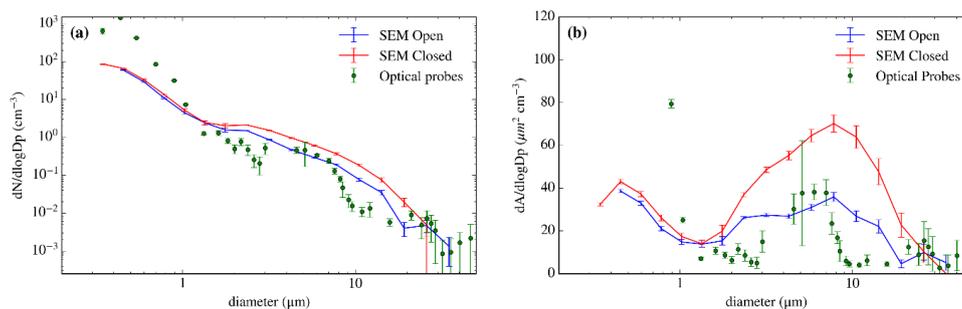
958

959



960
961 Figure 8. Bypass test carried out during the C010 flight on the 2017/05/10 from 11:24 to 11:38 UTC. The lower line sampled
962 226 L with the bypass closed, whereas the upper line sampled 141 L with the bypass open. The flow rates were 16.1 L min^{-1}
963 and 10.6 L min^{-1} respectively. The optical probes are the PCASP-CDP, using the closest calibration to the sampling date and a
964 refractive index of 1.56 as stated in the Sect. 2.3. The data is shown in both number size distribution (a) and surface area size
965 distribution (b). The only error source considered for the SEM size distribution is the Poisson counting error, which has been
966 included in this figure and all the subsequent figures including SEM size distributions.

967

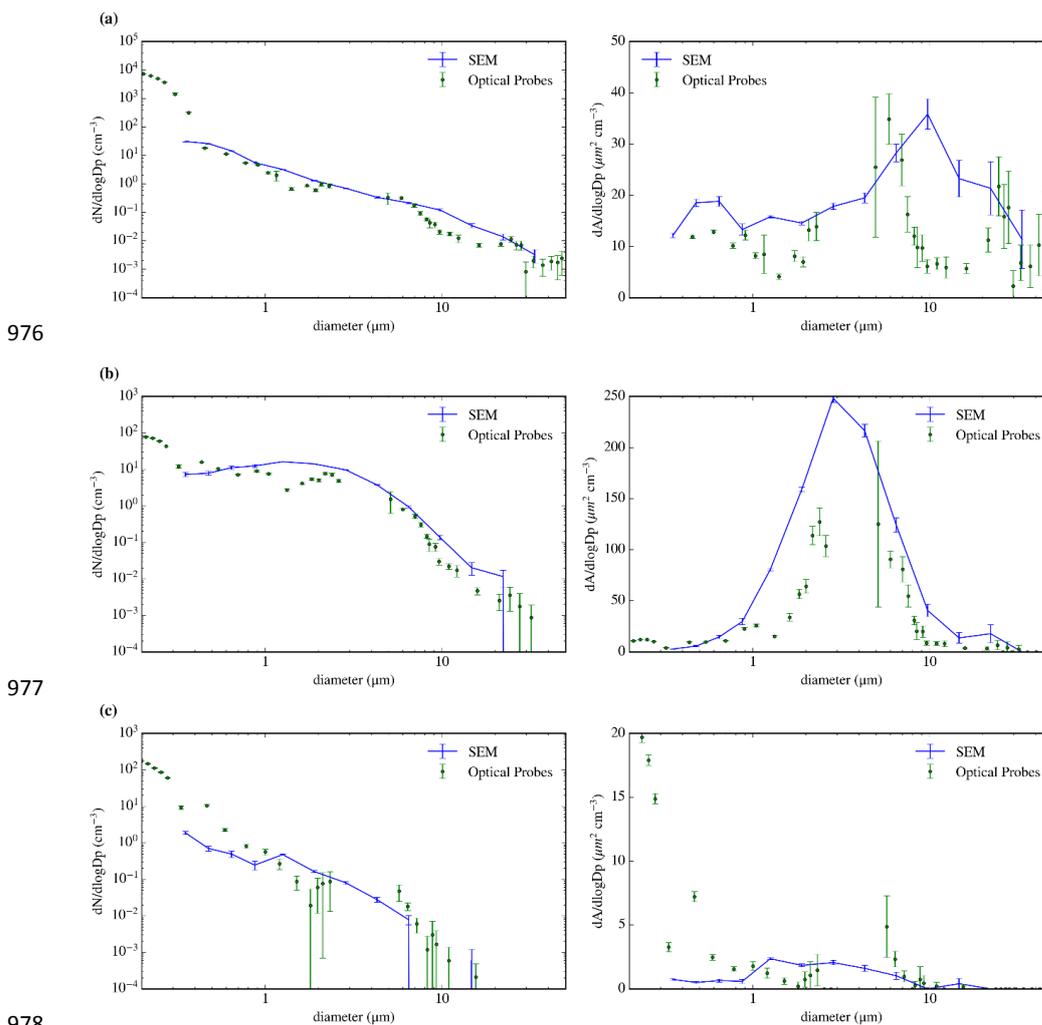


968

969 Figure 9. Bypass test carried out during the C057 flight on the 2017/09/27 from 13:33 to 13:50 UTC. The lower line sampled
970 555 L with the bypass open, whereas the upper line sampled 499 L with the bypass closed. The flow rates were 34.7 L min⁻¹
971 and 31.2 L min⁻¹ respectively. The position of the closed and open line was swapped with respect to the first analysis in Fig.
972 8. The optical probes are the PCASP-CDP, using the closest calibration to the sampling date and a refractive index of 1.56 as
973 stated in the Sect. 2.3. The sampling was interrupted for a minute to avoid a turn.

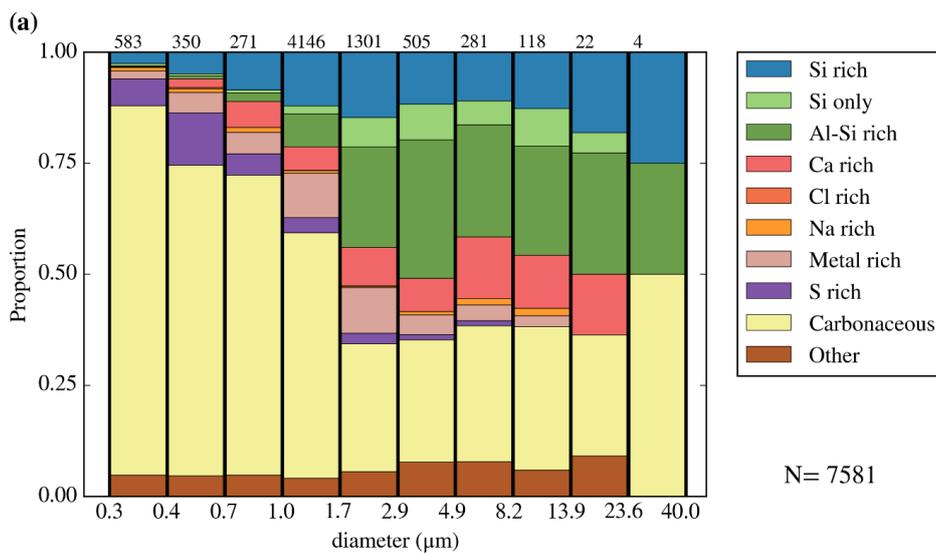
974

975

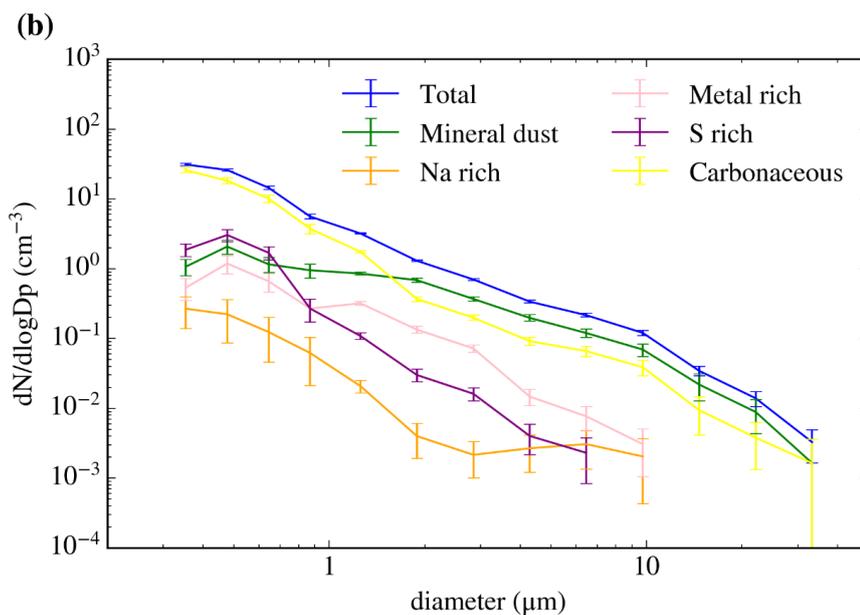


979 Figure 10. SEM obtained size distribution compared with PCASP-CDP online size distribution for three different sampling
980 periods in three different aerosol environments. Close to London, on the 2017/07/19 from 15:20 to 15:51 UTC, sampling 953
981 L (a), south of Iceland on the 2017/10/02 from 16:24 to 16:40 UTC, sampling 432 L (b) and in north Alaska on the 2018/03/20
982 from 20:15 to 20:37, sampling 724 L (c). All the sampling was done in the upper line with the bypass open. The flow rates
983 through the filter holders are 30.9, 30.5 and 42.0 L min⁻¹ respectively. The optical probes are the PCASP-CDP, using the closest
984 calibration to the sampling date and a refractive index of 1.56 as stated in Sect. 2.3.

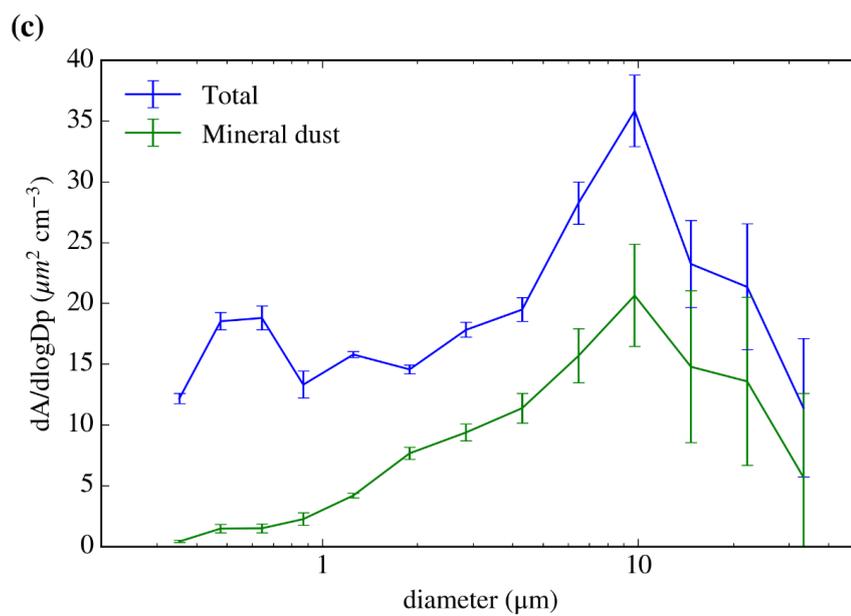
985



986



987



988

989 Fig. 11. Size-segregated compositional and morphological analysis of a sample collected close to London on the 2017/07/19
990 from 15:20 to 15:52 UTC by the lower line with the bypass open, sampling a total of 953L. (a) Fraction of particles
991 corresponding to each compositional category (described in the Sect. 7) for each size. The number of particles per bin can
992 be seen in the top of the figure. (b) Number size distribution for each composition. Cl rich particles were not included since
993 only two particles in this category were found. The errors have been calculated from the Poisson counting statistics (applying
994 it to both the size distribution and the compositional measurements). (c) Surface area of both all the detected aerosol
995 particles and the ones whose composition was consistent with mineral dust. Errors have been calculated in the same way as
996 before. By integrating the green curve in the figure (c) we obtained the total surface area of mineral dust in the sample (19.1
997 $\mu\text{m}^2 \text{cm}^{-3}$).

998