

Supplementary Information

Vertical profiles of the aerosol mass concentration observed by
unmanned airborne in-situ and remote sensing instruments
during a dust event

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18 **S1. Mass Concentration Calculation from OPC_a**

19 The mass concentration profiles of the coarse and fine particles were calculated from the size
20 distribution measurements recorded by OPC_a. Before converting the aerosol number concen-
21 trations to mass concentrations, the OPC measurements were averaged over 30 s. This was
22 found to be optimal for suppressing a high frequency noise of the OPC raw data and at the
23 same time for maintaining a relatively high spatial resolution of ≈ 80 m in the vertical direc-
24 tion. The number concentration (dN) of each size bin was converted to volume concentration
25 according to $dV(r) = dN(r)\frac{4}{3}\pi r^3$, where r is the mean radius of the size bin. The volume
26 concentration of particles with diameters larger than $1\ \mu\text{m}$ were summed and multiplied by
27 ρ_d yielding the coarse mode mass concentration (the same procedure was followed for the fine
28 mode particles). The variability in the number size distributions averaged every 30 s propa-
29 gated an uncertainty¹ of the order of 10% in the estimated volume size distributions and the
30 mass concentrations of the particles. Similar calculations but using ρ_{nd} were performed for
31 the fine fraction.

32 **S2. POLIPHON Method - Error Calculation**

33 The uncertainties of β_d and β_{nd} in equation (1) of the main manuscript were calculated
34 using the Monte-Carlo method⁵. For each input parameter of equation (1), we generated 100
35 normally distributed random numbers. The values provided in Table S1 were used as the
36 mean parameter and the standard deviation of the normal distributions. Then, 100 β_d and
37 β_{nd} values were calculated for each point in the atmospheric column and from these the mean
38 values and the standard deviations (errors) of β_d and β_{nd} were estimated. For equations (2)
39 and (3) the uncertainties were calculated analytically using the error propagation law¹.

40 **S3. Measurements of the Dust lidar Ratio**

41 One of the important input parameters of the POLIPHON algorithm is the S_d value, which for
42 the analysis described here was measured by the LIDAR. Actual measurements of S_d were only
43 possible during night-time when the Raman channels were operating. We assumed constant
44 S_d values for all the events analyzed here. This was supported by the backtrajectory analysis
45 showing that the air masses arriving at Cyprus on 15-04-2016 at 01:00 UTC and 07:00 UTC
46 had the same origin (North Sahara) and the same aerosol composition (dust; as depicted by
47 the LIDAR).

48 The vertical profiles (recorded between 00:00-01:40 UTC on 15 April and retrieved by the

49 Raman method) of α (at 355 nm, 532 nm), β (at 355 nm, 532 nm, 1064 nm), S (at 355
50 nm, 532 nm), extinction and backscatter related \hat{A} ($\hat{A}_{355/532}^\beta$, which was calculated from α
51 and β at 355 nm and 532 nm) and δ^p (at 355 and 532 nm) are plotted in Figure S2. The
52 particle depolarization ratio at 532 nm between 3-6 km, ranged from 27 to 31%, which are
53 typical values for pure Saharan dust²⁻⁴. From 2 to 3 km altitude, δ^p was ranged between
54 0.1 and 0.3 as a result of the entrainment of anthropogenic particles into the dust layer. This
55 vertical distribution of the aerosol particles is also confirmed from the $\hat{A}_{355/532}^\beta$ which decreases
56 gradually up to 3 km, reflecting that the coarse particles started to dominate over the fine
57 particles with increasing height.

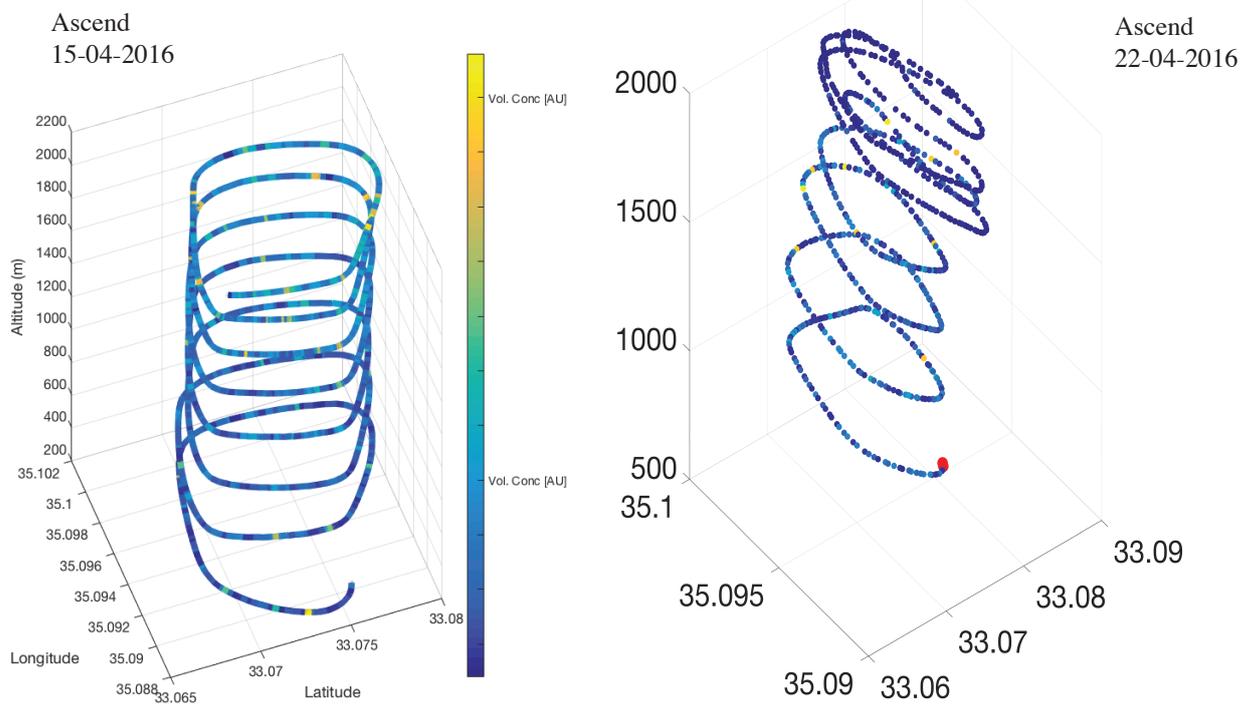


Figure S1: UAV flight patterns.

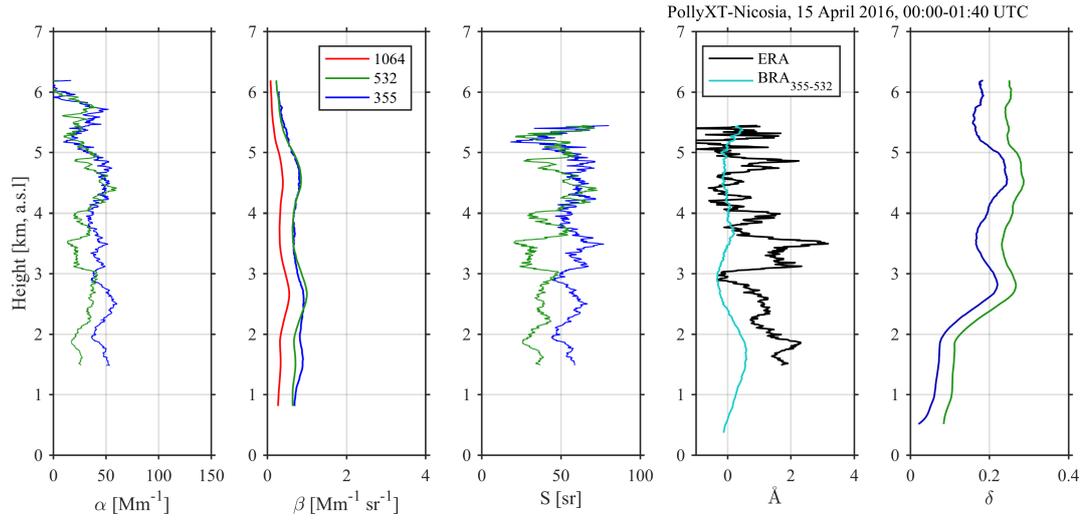


Figure S2: Night-time vertical profiles of extinction coefficient (355 nm; 532 nm), backscatter coefficient (355 nm; 532 nm; 1064 nm), LIDAR ratio (355 nm; 532 nm), Ångström exponent (extinction and backscatter related (355 nm; 532 nm)) and particle depolarization ratio (355 nm; 532 nm).

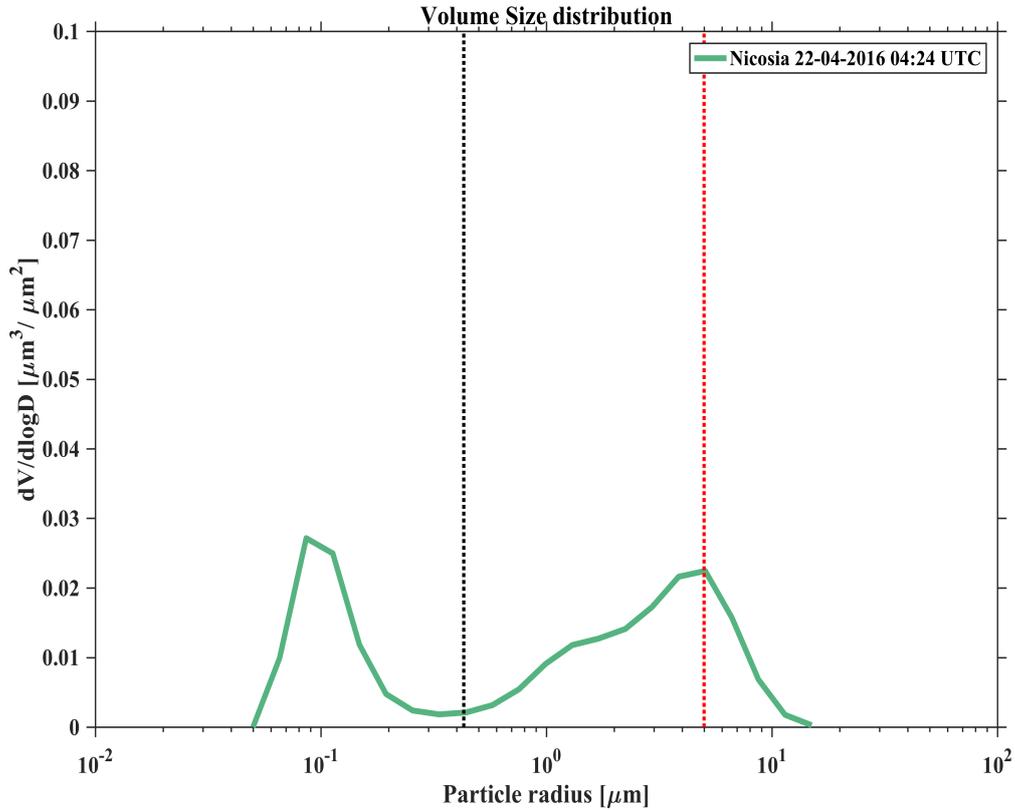


Figure S3: Column-integrated volume size distribution measured with the sunphotometer over Nicosia at 04:24 UTC on 22 April 2016. The black and red vertical lines indicate the inflection point and the upper limit of particle size measured by the OPC_a , respectively.

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