

# 1 A method for stable carbon isotope ratio and concentration 2 measurements of ambient aromatic hydrocarbons

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8  
9 **Abstract.** A technique for compound specific analysis of stable carbon isotope ratios and  
10 concentration of ambient volatile organic compounds (VOC) is presented. It is based on selective  
11 VOC sampling onto adsorbent filled cartridges by passing large volumes of air (up to 80 L)  
12 through the cartridge. The hydrocarbons are recovered by thermal desorption followed by two  
13 step cryogenic trapping and then are separated by gas chromatography in the laboratory. Once  
14 separated, individual VOC are subjected to online oxidation in a combustion interface and  
15 isotope ratio analysis by isotope ratio mass spectrometry. The method allows measurements of  
16 stable carbon isotope ratios of ambient aromatic VOC present in low pptV to ppbV levels with  
17 an accuracy of typically better than 0.5 ‰. The precision of concentration measurements is better  
18 than 10 %. Examples of measurements conducted as part of a joint Environment Canada-York  
19 University (EC-YU) measurement campaign at a semi-rural location demonstrate that the ability  
20 to make accurate measurements in air with low VOC mixing ratios is important to avoid bias  
21 from an over-representation of samples that are strongly impacted by recent emissions.

## 22 23 1 Introduction

24  
25 Volatile ~~Organic-organic Compounds-compounds~~ (VOC) comprise an important class of  
26 atmospheric pollutants emitted in large quantities from anthropogenic and biogenic sources  
27 (Atkinson, 2000; Guenther et al., 2000; Niedojadlo et al., 2008; Piccot et al., 1992; Rudolph,  
28 2002; Sawyer et al., 2000). While their atmospheric mixing ratios are mostly in the ranges of  
29 pptV to ppbV, these compounds play an important role in tropospheric chemical processes  
30 (Atkinson, 2000; Jordan, 2009; Kleinman et al., 2003). Aromatic VOC are important constituents  
31 of urban and rural air masses (Forstener et al., 1997; Lurmann and Main, 1992). These  
32 compounds, mostly originating from fossil fuel use, are found in evaporated gasoline, fuels and  
33 solvents, vehicle exhaust and many other anthropogenic-related emissions (Hurley et al., 2001;  
34 Jang and Kamens, 2001; Reimann and Lewis, 2007). A large fraction of urban VOC are  
35 composed of aromatic compounds, up to 60 % to 75 % of which are benzene, toluene,  
36 ethylbenzene, and 1,2,4-trimethylbenzene (Jang and Kamens, 2001; Smith et al., 1998). The  
37 atmospheric oxidation of these aromatics by reaction with OH and NO<sub>3</sub> can result in formation

1 of ozone as well as oxygenated and nitrated products that can contribute to the formation of  
2 secondary organic aerosol (SOA) (Forstener and Flagan, 1997; Jang and Kamens, 2001).

3 In the atmosphere, VOC undergo various chemical and physical processes that lead to  
4 their transformation, removal, transport and re-distribution (Atkinson, 2000; Helmig et al., 2008;  
5 Jenkin and Clemmshaw, 2000; Parrish et al., 2007). The majority of the presently used methods  
6 use concentration measurements as an indicator of photochemical processing of VOC. However,  
7 since mixing of air parcels of different origin is a complex and dynamic process, use of  
8 concentration alone as a marker for photochemical processing is problematic (Parrish et al.,  
9 2007). The interpretation of the results is quite challenging due to the complexity of atmospheric  
10 processing and mixing, and thus many conclusions are made based on a substantial number of  
11 assumptions (de Gouw et al., 2005; Gelencsér et al., 1997; Jobson et al., 1998). The change in  
12 relative composition of the ambient samples is considered to be a better indicator of chemical  
13 processing, since the concentration ratios are less affected by physical mixing and dilution  
14 (Honrath et al., 2008; Kleinman et al., 2003; McKeen and Liu, 1993; McKeen et al., 1996;  
15 Parrish et al., 2007; Roberts et al., 1984; Rudolph and Johnen, 1990), but still requires several  
16 assumptions about the history of the studied air mass. It was recently shown that the use of the  
17 stable carbon isotope ratios of VOC requires very few and easily tested assumptions and  
18 therefore is more useful in providing insights into photochemical transformation and mixing of  
19 VOC in ambient air (Goldstein and Shaw, 2003; Roberts et al., 1984; Rudolph and Czuba, 2000;  
20 Rudolph et al., 2002; Rudolph et al., 2003; Rudolph, 2007; Stein and Rudolph, 2007).

21 Measuring the isotope composition of VOC in the atmosphere is challenging due to the  
22 high precision and accuracy necessary to derive meaningful information. Rudolph et al. (1997)  
23 published a method for compound specific determination of the stable carbon isotopic  
24 composition for atmospheric VOC at sub-ppbV levels. The uncertainty of measured isotope  
25 ratios was close to 0.5 ‰, and Rudolph et al. (1997) suggested that further improvements in  
26 method may allow a precision close to 0.1 ‰. Within several years different research groups  
27 published results of stable carbon isotope measurements for a variety of atmospheric VOC  
28 (Anderson et al., 2004; Czapiewski et al., 2002; Iannone et al., 2003; Iannone et al., 2005; Irei et  
29 al., 2006; Norman et al., 1999; Rogers and Savard, 1999; Rudolph et al., 2002; Rudolph et al.,  
30 2003; Smallwood et al., 2002; Thompson et al., 2003). Nevertheless, the number of publications  
31 on isotopic composition measurements and their application is still quite limited due to the need  
32 for elaborate and expensive experimental techniques and challenging data interpretation  
33 (Eckstaedt et al., 2011; Fisseha et al., 2009; Giebel et al., 2010; Iannone et al., 2005; Iannone et  
34 al., 2009, 2010; Irei et al., 2006; Li et al., 2010; Moukhtar et al., 2011). An overview of existing  
35 techniques to measure stable carbon isotope ratios of VOC is given in a recent paper (Gensch et  
36 al., 2014).

37 Most of the online techniques used for measurements of the stable carbon isotope  
38 composition include a combination of a combustion interface (CI), a gas chromatograph (GC)  
39 and an isotope ratio mass spectrometer (IRMS) (Matthews and Hayes, 1978). In addition, for

1 gaseous compounds, sample purification, pre-concentration and separation steps are frequently  
2 added (Anderson et al., 2003; Czapiewski et al., 2002; Rudolph et al., 2002; Iannone et al., 2007;  
3 Iannone et al., 2010; Redeker et al., 2007). While VOC analysis by GC-IRMS is well  
4 established, collection of ambient VOC for the isotope analysis is challenging. One of the  
5 requirements for accurate GC-IRMS measurements is sufficient mass, which is usually 3 ng to 5  
6 ng of carbon for each analysed VOC (Goldstein and Shaw, 2003; Rudolph, 2007; Thompson,  
7 2003). Consequently, compounds that are present at low pptV levels need to be extracted from  
8 30 L of air or more. Many of the currently reported stable carbon isotope ratio data for ambient  
9 VOC were obtained using whole air sampling in stainless steel canisters, which is adequate when  
10 measuring VOC at ppbV and high pptV mixing ratios such as for source studies or polluted areas  
11 (Czapiewski et al., 2002; Redeker et al., 2007; Rudolph et al., 1997; Rudolph et al., 2002; Saito  
12 et al., 2002). However, collecting whole air samples with the large volumes required for  
13 measurement of isotope ratios for VOC mixing ratios well below one ppbV is technically  
14 challenging (Wintel et al., 2013).

15 Cryogenic sampling has been used to collect VOC from very large volumes of air (Bill et  
16 al., 2004 Iannone et al., 2007 Zuiderweg et al., 2011) but the technical and logistical challenges  
17 for field sampling are substantial. Selective sampling on adsorbent cartridges is a widely used  
18 inexpensive method for measurement of VOC concentrations and has recently been applied to  
19 collect VOC samples for isotope ratio analysis (Eckstaedt et al., 2011; Eckstaedt et al., 2012).  
20 However, this sampling technique suffered from substantial sample breakthrough problems for  
21 volumes exceeding 3 L. Although it was shown that breakthrough did not significantly impact  
22 isotope ratios, the relatively low breakthrough volume substantially restricts the useable sample  
23 volume. Furthermore, breakthrough would seriously affect the ability to measure concentrations  
24 and isotope ratios in the same sample.

25 In this article a method for the sampling and isotope ratio analysis of ambient VOC from  
26 large air volumes is presented. It is based on selective sampling of atmospheric compounds onto  
27 cartridges filled with an adsorbent (Carboxene 569). Thermal desorption of VOC from the  
28 cartridges is followed by two-step cryogenic trapping, separation by GC and online analysis by  
29 IRMS. Various validation tests were conducted and the results will be discussed. A brief  
30 overview of results from measurements at two locations with different levels of atmospheric  
31 pollution will be given.

32

## 33 **2. Materials and method**

34

### 35 **2.1 Overview**

36

37 Ambient VOC were analysed for concentration and isotope ratios ( $\delta^{13}\text{C}$ ) using the  
38 following steps (1) collection onto adsorbent packed cartridges, (2) thermal desorption of VOC  
39 in a furnace at 553 K in a flow of high purity helium, (3) a two-stage preconcentration, (4)

1 chromatographic separation, (5) combustion to CO<sub>2</sub> and H<sub>2</sub>O in a furnace at 1173 K, (6) analysis  
2 of the CO<sub>2</sub> isotopologues by Isotope Ratio Mass Spectrometry (IRMS).

## 3 **2.2 Sampling**

4  
5 Selective VOC sampling from volumes between 20 L and 80 L of ambient air was done  
6 on adsorbent filled cartridges. The cartridges were made by filling stainless steel tubes of 13 cm  
7 to 15 cm in length and 1/4" OD (approximately 5 mm ID) (Swagelok, Canada), with 1 g (±0.2 g)  
8 of Carboxene 569 (Supelco Inc., Bellefonte, USA). To keep the adsorbent in place both ends  
9 were plugged with 0.3 g (±0.1 g) of quartz wool (Restek, USA). Both ends of the tubes were  
10 equipped with 1/4" stainless steel Swagelok nuts, which were closed with 1/4" stainless steel  
11 caps during storage and transportation.

12 Cartridges were cleaned at temperatures between 523 K and 573 K in a furnace while  
13 continuously purging them with a flow of 160 mL min<sup>-1</sup> to 200 mL min<sup>-1</sup> of pure helium for at  
14 least 24 hours. Used cartridges were cleaned for 12 hours to 24 hours. Cleaned cartridges were  
15 capped and stored at room temperature in closed glass containers. Some of the cleaned cartridges  
16 were analysed without sampling to determine blank values, as well as the influence of storage,  
17 transport and possible material degradation during use.

18 For sampling, one end of the sampling cartridge was connected to a 1/4" OD stainless  
19 steel inlet line (Swagelok, Canada) and the other end to a mass flow controller with a flow  
20 totalizer, and a pump. To prevent physical damage, the sampling equipment was placed inside a  
21 temperature controlled housing. The sampling flow rate was varied between 10 mL min<sup>-1</sup> and 50  
22 mL min<sup>-1</sup>, depending on the required sampling time and volume. Typically the sampled air  
23 volume ranged from 30 L to 50 L. For sampling under conditions of very high humidity, a  
24 stainless steel water trap was added to the sampling line. The water trap was cooled by a Portable  
25 Ice Machine (Polar by Greenway, USA) to approximately 290 K. Condensed water was removed  
26 from the trap at regular intervals, depending on sampling flow rate and ambient humidity, using  
27 a removable plug at the bottom of the trap.

28 Trapping efficiency of Carboxene 569 was tested by sampling ambient air through two  
29 cartridges connected in series. Stability of VOC sampled on cartridges was tested by loading  
30 cartridges with test mixtures and analyzing them after storage at room temperature or in the  
31 freezer.

## 33 **2.3 Sample processing**

### 35 **2.3.1 VOC desorption**

36  
37 VOC were extracted from the cartridges by thermal desorption. The cartridge was  
38 connected to a helium supply line on one end and to a preconcentration system on the other by  
39 1/16" OD heated SS tubing. During desorption the cartridge was placed in the center of a

1 temperature controlled 30 cm long ceramic furnace (Omega, USA). Standard conditions for  
2 cartridge analysis were 555 K for 40 min with a carrier gas flow rate between 60 mL min<sup>-1</sup> and  
3 80 mL min<sup>-1</sup>. A range of conditions were tested to identify optimum conditions. The flow rate of  
4 the carrier gas was varied from 30 mL min<sup>-1</sup> to 100 mL min<sup>-1</sup>, the desorption temperature from  
5 523K to 623 K, and the desorption time from 10 min to 50 min.

### 7 **2.3.2 Two stage preconcentration**

9 Desorbed VOC were cryogenically trapped from the helium gas flow using a custom  
10 build two stage preconcentration system (TSPS) (Fig. 1) similar in design to the method  
11 described by Rudolph et al. (1997). It contained two cryogenic traps: one (Trap 1) is a U-  
12 shaped 1/8" OD stainless steel tube filled with glass beads (60/80 mesh, Chromatographic  
13 Specialties Inc., Canada); the other (Trap 2), consisted of a 15 cm long piece of 1/32" DB-1  
14 capillary column placed in a 1/16" OD stainless steel tube. The GC column was connected with  
15 a Valco zero volume connector (VICI, USA) to a 1/32" stainless steel tube that was attached to  
16 a six-port valve. The heating of transfer lines and the valve were individually temperature  
17 controlled using temperature sensors (Quick disconnect thermocouple assembly with 12" and  
18 18" length and 1/16" diameter 304 stainless steel probes (Omega, USA)) and were constantly  
19 maintained at 473 K. For cooling, Trap 1 was immersed in liquid nitrogen (10 cm to 13 cm  
20 deep), and Trap 2 was cooled by a liquid nitrogen flow that was controlled by a solenoid valve.

21 The two-position six-port valve was used to direct the gas and sample flow either through  
22 Trap 1 (Position A) or bypassing it (Position B). At the start of the analysis sequence the six-  
23 port valve was set to Position B and the cartridge was flushed with pure helium for five minutes.  
24 At the same time Trap 1 was cooled to 97 K. During desorption, the valve was set to Position A  
25 and desorbed VOC were transferred in a flow of helium to Trap 1, where they were adsorbed on  
26 the glass beads. At the end of this stage, Trap 2 was cooled to 97 K and the valve was rotated  
27 back to Position B, directing the GC carrier gas through rapidly heated Trap 1 (423K) to Trap 2.  
28 After seven minutes, Trap 2 was flash-heated to 493 K by applying approximately 2.5 V to the  
29 1/16" OD stainless steel tube for seven minutes, injecting the VOC into the GC column. For  
30 testing purposes, temperatures of the traps were varied between 93 K and 123 K for trapping  
31 and 393 K and 513 K for desorption.

### 33 **2.3.3 Gas chromatographic separation and VOC combustion**

35 The separation was performed in a HP5890 Series II gas chromatograph equipped with a  
36 DB-1 column (100 m, 0.25 mm ID, 0.5 µm film thickness). The VOC were focused on the head  
37 of the column at 298 K for ten min, then the oven was heated at 2 K min<sup>-1</sup> to 323 K followed by  
38 heating at 3K min<sup>-1</sup> to 363 K and then at 10 K min<sup>-1</sup> to 403 K. The column temperature was then  
39 maintained at 403 K for ten minutes. At the end of each analysis the oven temperature was

1 raised to 473 K and kept at this temperature for ten to twenty minutes. Helium was used as  
2 carrier gas with the flow rate controlled by an electronic pressure controller (EPC) at  
3 approximately 2 mL min<sup>-1</sup>. Once separated in the GC column, the column effluent was directed  
4 either to the FID or the combustion interface by opening or closing a pneumatic valve as shown  
5 in Fig. 1.

6 The combustion interface consisted of a 1/4" ceramic tube (0.5 mm ID, 44 cm length)  
7 with copper, nickel and platinum wires inside. Helium containing traces of oxygen was added as  
8 a makeup gas at the inlet of the furnace to prevent peak tailing and to provide oxygen. During  
9 analysis the interface was kept at 1223 K, and CuO and NiO formed on the wire surfaces  
10 converted VOC into CO<sub>2</sub> and H<sub>2</sub>O. The oxide layers on the Cu and Ni wires were regenerated  
11 every night by flushing the furnace and adjacent tubing with the He-O<sub>2</sub> mixture while reducing  
12 the interface temperature to 823 K. A flow restrictor split (Fig. 1) was used in an open split  
13 configuration to direct between 10 % and 20 % of the flow through a Nafion Dryer (25 cm, 0.6  
14 mm ID, 0.8 mm OD) into the IRMS.

#### 16 2.4 Isotope ratio mass spectrometry and δ<sup>13</sup>C determination

17  
18 CO<sub>2</sub> was analysed by IRMS (GV Instruments, Manchester, UK) by detection of m/z 44,  
19 45 and 46. The isotope ratios <sup>45</sup>R and <sup>46</sup>R were determined from the peak areas (A) for the  
20 individual m/z ratios as described by Rudolph (2007):  $^{45}\text{R} = \frac{A[^{45}\text{CO}_2]}{A[^{44}\text{CO}_2]}$  and  $^{46}\text{R} = \frac{A[^{46}\text{CO}_2]}{A[^{44}\text{CO}_2]}$   
21  
22 <sup>46</sup>R was used for correction of the <sup>45</sup>R for the contribution from <sup>12</sup>C<sup>16</sup>O<sup>17</sup>O (Craig, 1957;  
Santrock et al., 1985).

23 Peak integration was done manually using a method similar to that described by Rudolph  
24 et al. (1997). To minimize bias due to the manual definition of peak boundaries, each peak was  
25 integrated ten times with start and end points varying every time by 0.1 s. This allowed verifying  
26 that averaged delta values were not significantly biased by the choice of peak boundaries. All  
27 target VOC, except p- and m-xylene, were completely separated. Due to their partial overlap the  
28 peaks for p- and m-xylene were integrated together.

29 Throughout the run, reference CO<sub>2</sub>, calibrated relative to the V-PDB (Vienna Peedee  
30 Belemnite) standard was introduced several times directly into the IRMS in pulses of 20 s to 40 s  
31 duration. Between three and five reference gas injections were made while the pneumatic valve  
32 was open and the flow from the GC column was directed to the FID at the beginning and at the  
33 end of each run. Between four and seven injections were made between peaks during the  
34 chromatographic separation while the flow from the GC column was directed to the combustion  
35 interface and IRMS.

#### 36 37 2.5 Validation tests

1  
2 Gaseous test mixtures of target VOC were prepared in stainless steel canisters with VOC  
3 mixing ratios in the ppbV to ppmV range by injecting known quantities of pure VOC and  
4 diluting with helium. The VOC concentration in these test mixtures ranged from 0.1 ng cm<sup>-3</sup> to  
5 0.7 µg cm<sup>-3</sup> (15 ppbV to 1.7 ppmV). All mixtures were quantified using a standard mixture of  
6 VOC (ppbV levels) provided by the National Research Council of Canada (NRC). VOC used to  
7 prepare gaseous mixtures, were individually combusted at high temperatures in vacuum sealed  
8 tubes containing CuO. The resulting CO<sub>2</sub> was cryogenically separated, extracted and later  
9 analysed by IRMS, using a dual-inlet system, to determine their δ<sup>13</sup>C values. These offline  
10 isotope ratios were: benzene -28.40 ‰ ± 0.02 ‰, toluene -27.02 ‰ ± 0.07 ‰,  
11 ethylbenzene -26.84 ‰, p-xylene -25.69 ‰ ± 0.05 ‰, m-xylene -26.92 ‰ and o-xylene -28.16  
12 ‰ ± 0.07 ‰. Standard deviations for ethylbenzene and m-xylene cannot be given since repeat  
13 measurements are not available. These mixtures were used for calibration as well as  
14 determination of the accuracy of the δ<sup>13</sup>C measurements.

15 A number of tests were conducted to optimize and evaluate the accuracy and  
16 reproducibility of the method. Operation conditions and instrument performance were tested  
17 weekly using mixtures with known composition and isotope ratios. Precision of the system was  
18 determined by the reproducibility of peak areas from repeat measurements of the test mixtures.  
19 Accuracy of isotope ratio measurements was tested by comparing δ<sup>13</sup>C values determined  
20 online for the test mixtures and offline for the individual VOC.  
21

Comment [A1]: replace 'd' by 'δ'(=delta)

### 22 3. Results and discussion

23 The components of the sampling and analysis system were tested under different  
24 operational conditions. Here, the most important results of the optimization and validation tests  
25 are presented and discussed.  
26

#### 27 3.1 Reproducibility, accuracy and linearity of the TSPS-GC-IRMS system

28  
29 Performance assessment of the TSPS was based on injections of gaseous mixtures  
30 directly into the TSPS. On average, the relative standard deviation (RSD, %) of peak areas from  
31 repeat measurements (>10 repeats) was 5.3 %, with 4 % for alkanes and 6 % for aromatic  
32 compounds. For all targeted compounds, calibration curves were constructed with intercepts for  
33 regression set to zero, in most of the linear regression analysis the correlation coefficient (R<sup>2</sup>  
34 value) was >0.99. The intercept was set to zero since intercepts in linear regressions allowing  
35 non-zero intercepts did not significantly differ from zero. An example of a calibration curve is  
36 shown in Fig. 2. The sensitivity of the GC-IRMS measurements was in the range of (1.2 to 3.4)  
37 × 10<sup>10</sup> ions ng<sup>-1</sup> carbon.

1 Figure 3 shows the measured  $\delta^{13}\text{C}$  values as function of sample mass for aromatic VOC  
2 and n-alkanes, also shown are the  $\delta^{13}\text{C}$  values from the offline analysis of bulk VOC. The  
3 standard deviation for repeat measurements of online  $\delta^{13}\text{C}$  was mostly lower than 0.3 ‰, on  
4 average 0.2 ‰ for >10 ng of alkanes and 0.3 ‰ for >1 ng of aromatics. Generally the standard  
5 deviation was lower for higher masses. For masses exceeding 5 ng, the measured  $\delta^{13}\text{C}$  values  
6 were independent of sample mass, but for lower masses, the measured values systematically  
7 decreased with decreasing mass. The only exception was the combined measurement of p-xylene  
8 and m-xylene, where the decrease in measured  $\delta^{13}\text{C}$  is already significant at sample masses  
9 below 10 ng. This is most likely due to the necessity of a wider time window for integration of  
10 the two overlapping peaks and the resulting greater impact of baseline drift and baseline noise.

11 For aromatic VOC offline and online values usually agreed for masses of more than 3 ng  
12 to 5 ng per compound within the uncertainty of the measurements (Fig. 3), however there was a  
13 significant bias towards lower delta values for smaller masses. For alkanes a significant  
14 difference between offline and online data was observed, even though both offline and online  
15 tests displayed good reproducibility, with uncertainties well below the observed difference.

16 No certified standards or equivalent internationally accepted reference mixtures suitable  
17 for tests of the accuracy of isotope ratio measurement of atmospheric VOC are available and  
18 consequently the accuracy tests had to be conducted using mixtures prepared in-house. In the  
19 case of good agreement between GC-IRMS analysis and the isotope ratio of the test mixtures this  
20 gives confidence in the accuracy of the GC-IRMS measurements. However, in case of a  
21 systematic discrepancy between off-line and GC-IRMS results this makes it difficult to identify  
22 the source of bias. While isotope fractionation during sampling on the adsorbent cartridges can  
23 be ruled out since the isotope ratios obtained from direct analysis of the test mixtures agreed with  
24 isotope ratios determined from samples of the test mixtures collected on adsorbent cartridges,  
25 uncertainty in the carbon isotope ratio of the test mixture as well as biased GC-IRMS isotope  
26 ratio measurements are possible explanations.

27 Error in the  $\delta^{13}\text{C}$  in the test mixture can be due to loss or contamination during  
28 preparation of the mixture or error in the off-line  $\delta^{13}\text{C}$  determination for the bulk n-alkanes.  
29 Since the mixtures were not prepared for the purpose of calibrating concentration measurements  
30 it is not possible to exclude or verify the existence of losses or contaminations from comparison  
31 of peak areas or calibration curves. The uncertainty of conventional dual-inlet  $\delta^{13}\text{C}$   
32 measurements in carbon dioxide is much lower than the observed discrepancy in the n-alkane  
33 carbon isotope ratios. However, incomplete oxidation of n-alkanes during the off-line  
34 combustion of the bulk n-alkanes could result in significant carbon isotope fractionation. Since  
35 n-alkanes are less reactive than benzene and alkylbenzenes, this would be consistent with the  
36 absence of bias for aromatic VOC.

37 Bias in the GC-IRMS analysis can be caused by contamination or losses during  
38 adsorption or desorption of VOC in the TSPS, problems in the chromatographic separation such  
39 as peak tailing or baseline drift, and incomplete oxidation in the combustion interface. A

1 significant impact from contamination or artifacts can be ruled out since this would result in a  
2 decrease of bias for high sample masses, which is not found (Figure 3). Although in principle  
3 none of the other explanations can be ruled out, there is no direct evidence which would support  
4 them. The good reproducibility of the measured n-alkane carbon isotope ratios would require  
5 that any loss in the TSPS would be highly reproducible, which is unlikely and the tests of the  
6 TSPS showed no indication for significant loss of n-alkanes in the TSPS. Except for n-pentane  
7 there also is no indication for problems due to chromatographic separation, peak shape, or  
8 baseline drift for n-alkanes. Due to the higher chemical stability of n-alkanes incomplete  
9 oxidation in the combustion interface could result in isotope fractionation for n-alkanes without  
10 influencing the results for aromatic VOC, although this would require inverse isotope effects for  
11 the catalytic oxidation of n-alkanes in order to explain the bias towards higher carbon isotope  
12 ratios.

### 14 3.2 Cartridge tests

15 The VOC masses observed for analysis of blank cartridges using desorption temperatures  
16 in the range of 474 K to 573 K are listed in Table 1 together with their standard deviations and  
17 the resulting  $3\sigma$  detection limit. For comparison, the averages of masses of VOC collected in  
18 ambient samples are also given. Tests of individual components of the cartridges showed that the  
19 main source of the blanks was Carboxene 569. Stainless steel tubes did not produce any  
20 contamination; tests of cartridges containing only quartz wool at temperatures ranging from 295  
21 K to 590 K showed no or only marginal signals. As can be seen from Table 1, the blank values  
22 were small compared to typical VOC masses in ambient samples. For the typical sample volume  
23 of 30 L the detection limits correspond to mixing ratios in the range of 3 pptV to 15 pptV.

24 Optimum desorption conditions were determined by analysing cartridges loaded with  
25 approximately 30 ng to 60 ng of VOC using the gaseous test mixtures. Desorption times,  
26 temperatures and flow rates were varied (Table 2). A flow rate of 50 mL to 80 mL was used for  
27 most of the tests. Lower desorption temperatures and shorter desorption times (Table 2) resulted  
28 in poor recovery of compounds of low volatility. However, the option of using very high  
29 temperatures is limited since a prolonged exposure to high temperatures may result in a  
30 degradation of the cartridge materials. Since recovery was, within the uncertainty of the  
31 measurements, quantitative at 553 K and 40 min, these conditions were chosen as standard  
32 desorption conditions. Table 3 shows the average recovery and its reproducibility for the chosen  
33 standard conditions.

34 Sampling efficiency tests using two cartridges in series showed that for sample volumes  
35 of up to 80 L, the second cartridge only contained between 2 % and 5 % of the mass found on the  
36 first cartridge. These masses often were close to or below the detection limits and therefore only  
37 provide an upper limit for breakthrough during sampling.

1 Capped sample-containing cartridges could be stored both at room temperature for a short  
 2 period and in a freezer for longer time with no significant loss of VOC (Fig. 4). Generally the  
 3 change in mass during storage is less than 10 %. Similarly, the isotopic fractionation during  
 4 storage was less than 0.3 ‰ (Fig. 5). The masses of VOC loaded for these tests were in the range  
 5 of 25 ng to 75 ng (Table 3).

### 6 3.3 Ambient measurements

7  
 8  
 9 The newly developed method was applied in an ambient air quality field study in 2009 to  
 10 2010 (Kornilova et al., 2013) at Egbert, a semi-rural location in Southern Ontario. Details about  
 11 the sampling site can be found in Kornilova et al., 2013. This paper also contains a detailed  
 12 discussion of the use of the measured VOC carbon isotope ratios to quantitatively evaluate  
 13 photochemical VOC processing. Therefore, the discussion here will be limited to aspects related  
 14 to performance of the analytical method.

15 Example chromatograms for the test mixture and an ambient sample are provided in Fig.  
 16 6. For ambient samples, it was found that n-pentane was subject to substantial peak overlap,  
 17 which prevents reliable peak evaluation. Therefore, to minimize the risk of contamination of the  
 18 combustion interface and the IRMS, monitoring the column effluent by IRMS started with  
 19 hexane. Otherwise all target compounds were well separated with the exception of p-xylene and  
 20 m-xylene, which were integrated as one peak. It should be noted that the change in time window  
 21 of monitoring the column effluent changes the retention times, especially for peaks eluting early  
 22 in the time window which is monitored by the IRMS. This is the result of a small counter  
 23 pressure at the end of the separation column which results from the gas-flow that is used to  
 24 prevent the column effluent from entering the combustion interface and the resulting lower  
 25 pressure difference between the head and the end of the separation column.

26 For most of the samples the peaks for n-octane, n-nonane, and n-decane corresponded to  
 27 masses below the lower limit of the linear range and for these compounds the number of  
 28 available data points is too small to allow calculation of meaningful averages. Problems with  
 29 peak evaluation due to small peak areas were not limited to the C<sub>8</sub>-C<sub>10</sub> n-alkanes, but occurred  
 30 less frequently for aromatic VOC. The total number of samples that could be analysed for  $\delta^{13}\text{C}$   
 31 was 50 and the number of  $\delta^{13}\text{C}$  measurements available ranges from 46 for benzene to 26 for  
 32 hexane.

33 Table 4 gives an overview for concentrations and  $\delta^{13}\text{C}$ . A comparison between measured  
 34  $\delta^{13}\text{C}$  and  $\delta^{13}\text{C}$  of typical urban VOC emissions, which are also included in Table 7, shows that  
 35 the measured  $\delta^{13}\text{C}$  values generally are heavier than those of typical emissions. This is expected  
 36 since photochemical processing will typically result in enrichment of  $^{13}\text{C}$  (Rudolph et al., 2002;  
 37 Rudolph and Czuba, 2000). The only exception is hexane; here the lower end of the  $\delta^{13}\text{C}$  values  
 38 are about 1.4 ‰ lower than the  $\delta^{13}\text{C}$  values of urban emissions, which is outside of the  
 39 uncertainty range of known urban emissions and the measurement precision. The possible bias of

**Comment [A2]:** R2: Can you comment on the apparent positive artifacts for n-hexane and benzene recovery with storage?

**Comment [A3]:** Do we need to discuss this? These would be pure speculations, not sure if they are valid here.

1 n-hexane online measurements cannot explain this difference since a correction for a possible  
2 bias derived from the difference between online and offline results (Fig. 3) would result in an  
3 even larger difference, sometimes exceeding 3 ‰. This suggests the existence of a local or  
4 regional n-hexane source with a  $\delta^{13}\text{C}$  value lower than typical urban emissions. The existence of  
5 a substantial unidentified regional n-hexane source is consistent with n-hexane mixing ratios  
6 which are often high compared to that of other VOC. It should be noted that the results for n-  
7 hexane may be biased towards samples at the higher end of mixing ratios due to the significant  
8 number of samples for which reliable evaluation of  $\delta^{13}\text{C}$  was not possible.

9 Due to the relatively high sample mass needed for  $\delta^{13}\text{C}$  measurement by GC-IRMS bias  
10 due to the problem of reliable evaluation of  $\delta^{13}\text{C}$  in samples with low concentrations is a  
11 potentially significant problem. To evaluate this in more detail, we use results for aromatic VOC  
12 where between 80 % and 90 % of the samples could be analysed for  $\delta^{13}\text{C}$ . Figure 7 shows the  
13 dependence between measured  $\delta^{13}\text{C}$  and concentration for benzene and p,m-xylene. Compared to  
14 the  $\delta^{13}\text{C}$  values of emissions, the samples are enriched in  $\delta^{13}\text{C}$  to various degrees. The number of  
15 samples substantially enriched in  $\delta^{13}\text{C}$  is higher at low mixing ratios than at high mixing ratios.  
16 Although no statistically significant overall trend can be determined due to the substantial  
17 seemingly random variability of the data, the mean  $\delta^{13}\text{C}$  value for the lowest ten concentrations  
18 is  $-23.1 \text{ ‰} \pm 0.1 \text{ ‰}$  (error of mean) for benzene and  $-22.9 \text{ ‰} \pm 0.9 \text{ ‰}$  for p,m-xylene, which  
19 indicates a substantial difference to the average for all measurements, which is  $-25.2 \text{ ‰} \pm 0.1 \text{ ‰}$   
20 and  $-25.4 \text{ ‰} \pm 0.5 \text{ ‰}$  for benzene and toluene, respectively. This is a strong indication that  
21 limitations in making precise  $\delta^{13}\text{C}$  measurements at very low mixing ratios can bias the average  
22 measured  $\delta^{13}\text{C}$  towards lower values. The origin of such bias is easily understood, air with higher  
23 mixing ratios of VOC is more likely to have been subjected to significant recent emissions than  
24 clean air.

### 27 3.4 Comparison with existing methods

29 Due to the complexity of the existing methods used for sampling and measurements of  
30 stable carbon isotope compositions of ambient VOC, the number of publications on method  
31 development and ambient measurements and their interpretation is still quite limited. An  
32 overview of existing methods has been given in a recent paper by Gensch et al. (2014). All  
33 methods for ambient measurements and emission studies use GC-IRMS for analysis of sampled  
34 VOC, although often details of the chromatographic separation and combustion interface differ.  
35 The main difference between the methods is the sampling and VOC enrichment procedure. This  
36 is not surprising since the analysis of VOC carbon isotope ratios requires enrichment of VOC  
37 from large sample volumes, which is technically challenging and may require substantial logistic  
38 effort in field studies.

1 In principle, the lower limit of mixing ratios for which meaningful VOC carbon isotope  
2 ratios can be conducted depends on the sample volume that can be used. This volume typically  
3 ranges from several litres for whole air sampling to several hundred litres or more for cryogenic  
4 sampling (Gensch et al., 2014 and references therein). With sample volumes of up to 80 L, the  
5 method developed in this work allows collection of significantly larger samples than methods  
6 using whole air sampling, but not for the extremely large sample volumes that can be used by  
7 some cryogenic methods (Bahlmann et al., 2011; Bill et al., 2004; Iannone et al., 2007).  
8 However, the adsorptive sampling method described in this paper requires less expensive  
9 instrumentation than cryogenic sampling and can be conducted in the field with instrumentation  
10 that is similar in requirements (electrical power, weight, storage, shipment) to whole air sampling  
11 in pressurized containers.

12 While adsorptive sampling as described in this paper allows use of larger air sample  
13 volumes than whole air sampling, it has some other limitations. Each sample can only be  
14 analysed once, this can be a problem in cases where the mixing ratios in the sampled air is highly  
15 variable or the range of mixing ratios expected is unknown, such as in many emission studies.  
16 Due to the limited dynamic range of IRMS measurements, sampled VOC masses may be outside  
17 of the dynamic range and thus lost. Whole air samples, in principle, allow repeat runs of the  
18 same sample using different volumes or IRMS settings. This is very useful for samples with high  
19 VOC concentrations, but of limited value for very low concentrations since existing methods for  
20 whole air sampling do not provide sufficient sample volume for several measurements at sub-  
21 ppbV levels.

22 Collection of whole air samples typically requires less time than adsorptive or cryogenic  
23 sampling. This can be a significant advantage if sampling time is limited, for example in airplane  
24 based sampling. On the other hand, adsorptive sampling can easily provide samples that are  
25 integrated over long time periods. For example, our study at Egbert used 24 hour sampling  
26 periods, which allows determination of representative averages without the necessity to collect  
27 and analyse an extremely large number of samples. Finally, adsorptive sampling is not suitable  
28 for collection of VOC with very high volatility such as C<sub>2</sub> or C<sub>3</sub> hydrocarbons.

29 The reproducibility of isotope ratio measurements for our newly developed method is  
30 better than 0.5 ‰, similar to that of other state-of-the-art methods for measurement of  $\delta^{13}\text{C}$  in  
31 ambient VOC (Czapiewski et al., 2002; Eckstaedt et al., 2012; Kawashima and Murakami, 2014;  
32 Rudolph et al., 1997; Saito et al., 2009; Turner et al., 2006; Wintel et al., 2013), however the  
33 lower limit of concentrations for which this precision can be achieved is significantly lower than  
34 that achieved by these other methods.

35 Tests of accuracy for  $\delta^{13}\text{C}$  measurements of aromatic VOC demonstrate that there is no  
36 detectable bias, similar to the findings in several other studies (Czapiewski et al., 2002;  
37 Eckstaedt et al., 2012; Rudolph et al., 1997; Saito et al., 2009). However, for n-alkanes the GC-  
38 IRMS results differ from the offline values for bulk material used to prepare test mixtures by  
39 approximately 2 ‰. Unfortunately there are not many published tests of the accuracy of  $\delta^{13}\text{C}$

1 measurements for C<sub>6</sub> and heavier n-alkanes in air. Rudolph (2007) summarized results from  
2 similar sets of comparisons, which included a wide range of n-alkanes and concluded that in  
3 most cases offline and online  $\delta^{13}\text{C}$  values agree within the uncertainty of the GC-IRMS  
4 measurement. However, in a few cases differences in the range of 1 ‰ were observed, although  
5 these differences were statistically not significant.

6

#### 7 **4. Summary and conclusions**

8

9 The newly developed method for analysis of atmospheric VOC allows measurement of  
10  $\delta^{13}\text{C}$  of light aromatic hydrocarbons with an accuracy and precision of better than 0.5 ‰ for  
11 mixing ratios in the lowest pptV range. For n-alkanes a similar precision could be achieved, but a  
12 systematic bias of up to 2 ‰ is possible. The precision of concentration measurements is  
13 estimated to be better than 10 %. Sampling is done with inexpensive equipment suitable for field  
14 work and samples can be stored for more than one week under ambient conditions and at least  
15 six months in a freezer without detectable sample degradation. This allows transport and storage  
16 of samples with little logistic effort. Overall, these performance characteristics are a significant  
17 improvement over that of other currently available methods.

18 The methodology has been specifically developed and tested for the purpose of accurate  
19 measurement of light aromatic hydrocarbons, but this does not exclude its use for analysis of  
20 other VOC. Tests demonstrated that breakthrough of heavier than C<sub>5</sub> VOC is negligible for up to  
21 80 L of air, but since no tests have been conducted for larger volumes, collection of VOC from  
22 even larger volumes may be possible with no or little modification of the sampling cartridges.

23 Tests of the accuracy of carbon isotope ratio measurements of n-alkanes showed a bias of  
24 approximately 2 ‰ between the isotope ratio of test mixtures prepared in-house and GC-IRMS  
25 measurements. It was not possible to identify if this is due to uncertainty in the isotope ratio of  
26 the test mixture or bias of GC-IRMS measurements. Nevertheless, this demonstrates the  
27 importance of testing the accuracy of GC-IRMS measurements of atmospheric VOC.  
28 Unfortunately the number of published comparisons between artificial VOC mixtures with  
29 known isotope ratios and GC-IRMS measurements is limited, a consequence of the substantial  
30 resources and effort required to prepare test mixtures with known isotope ratios and the  
31 unavailability of certified or otherwise well characterised reference mixtures suitable for such  
32 tests. This supports the necessity for creating reference materials for isotope ratio analysis of  
33 organic compounds in the atmosphere as pointed out by Gensch et al. (2014).

34 The measurements conducted over a semi-rural area demonstrate the importance of  
35 accurate measurements at low mixing ratios to avoid bias and lack of representativeness of the  
36 measurements. Higher mixing ratios often are the result of substantial impact of nearby sources  
37 and therefore have carbon isotope ratios close to the isotopic composition of emissions. Inability  
38 to measure  $\delta^{13}\text{C}$  for VOC at low mixing ratios therefore will bias the  $\delta^{13}\text{C}$  observations towards

1 the source composition and therefore underestimate the role of chemical processing, which  
2 results in enrichment of  $^{13}\text{C}$ .

3 Overall, the simplicity and affordability of the developed sampling and sample  
4 processing system is a valuable step towards the possibility of a wider application of stable  
5 carbon isotope measurements in studies of ambient VOC. However, it should be noted that GC-  
6 IRMS instrumentation is still demanding and expensive. Nevertheless, due to the option to  
7 collect VOC from large volumes of air that results in larger sample masses, the need for highly  
8 sensitive GC-IRMS instrumentation might not be critical. This is an important step towards  
9 establishing VOC isotope ratio measurements as a standard technique in atmospheric chemistry.

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14

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- Anderson, R., Czuba, E., Ernst, D., Huang, L., Thompson, A., and Rudolph, J.: Method for measuring carbon kinetic isotope effects of gas-phase reactions of light hydrocarbons with the hydroxyl radical, *J. Phys. Chem. A*, 107, 6191-6199, 2003.
- Anderson, R., Huang, L., Iannone, R., Thompson, A., and Rudolph, J.: Carbon kinetic isotope effects in the gas phase reactions of light alkanes and ethene with the OH radical at 296 +/- 4 K, *J. Phys. Chem. A*, 11537-11544, 2004.
- Atkinson, R.: Atmospheric chemistry of VOCs and NO<sub>x</sub>: *Atmos. Environ.*, 2063-2101, 2000.
- Bahlmann, E., Weinberg, I., Seifert, R., Tubbesing, C., and Michaelis, W.: A high volume sampling system for isotope determination of volatile halocarbons and hydrocarbons, *Atmos. Meas. Tech.*, 4, 2073-2086, 2011.
- Bill, M., Conrad, M. E., and Goldstein, A. H.: Stable carbon isotope composition of atmospheric methyl bromide, *Geophys. Res. Lett.*, 31, L04109, 2004.
- Craig, H.: Isotopic standards for carbon and oxygen and correction factors for mass-spectrometric analysis of carbon dioxide, *Geochim. Cosmochim. Ac.*, 12, 133-149, 1957.
- Czapiewski, C. V., Czuba, E., Huang, L., Ernst, D., Norman, A. L., Koppmann, R., and Rudolph, J.: Isotopic composition of non-methane hydrocarbons in emissions from biomass burning, *J. Atmos. Chem.*, 43, 45-60, 2002.
- de Gouw, J. A., Middlebrook, A. M., Warneke, C., Goldan, P. D., and Kuster, W. C.: Budget of organic carbon in a polluted atmosphere, *J. Geophys. Res.*, 110, D16305, 2005.
- Eckstaedt, C., Grice, K., Ioppolo-Armanios, M., Chidlow, G., and Jones, M.:  $\delta D$  and  $\delta^{13}C$  analyses of atmospheric volatile organic compounds by thermal desorption gas chromatography isotope ratio mass spectrometry., *J. Chromatogr A*, 1218, 6511-6517, 2011.
- Eckstaedt, C. V., Grice, K., Ioppolo-Armanios, M., Kelly, D., and Gibberd, M.: Compound specific carbon and hydrogen stable isotope analyses of volatile organic compounds in various emissions of combustion processes, *Chemosphere*, 89, 1407-1413, 2012.
- Fisseha, R., Spahn, H., Wegener, R., Hohaus, T., Brasse, G., Wissel, H., Tillmann, R., Wahner, A., Koppmann, R., and Kiendler-Scharr, A.: Stable carbon isotope composition of secondary organic aerosol from  $\beta$ -pinene oxidation, *J. Geophys. Res.*, 114, D02304, 2009.
- Forstener, H. J., Flagan, R. C., and Seinfeld, J. H.: Secondary organic aerosol from the photooxidation of aromatic hydrocarbons: Molecular composition, *Environ. Sci. Technol.*, 31, 1345-1358, 1997.

- Gelencsér, A., Siszler, K., and Hlavay, J.: Toluene–benzene concentration ratio as a tool for characterizing the distance from vehicular emission sources, *Environ. Sci. Technol.*, 31, 2869-2872, 1997.
- Gensch, I., Kiendler-Scharr, A., and Rudolph, J.: Isotope ratio studies of atmospheric organic compounds: Principles, methods, applications and potential, *Int. J. Mass Spectrom.*, 365-366, 206-221, 2014.
- Giebel, B. M., Swart, P. K., and Riemer, D. D.:  $\delta^{13}\text{C}$  Stable isotope analysis of atmospheric oxygenated volatile organic compounds by gas chromatography-Isotope Ratio Mass Spectrometry, *Anal. Chem.*, 82, 6797-6806, 2010.
- Goldstein, A., and Shaw, S.: Isotopes of volatile organic compounds: An emerging approach for studying atmospheric budgets and chemistry., *Chem. Rev.* 103, 5025-5048, 2003.
- Guenther, A., Geron, C., Pierce, T., Lamb, B., Harley, P., and Fall, R.: Natural emissions of non-methane volatile organic compounds, carbon monoxide, and oxides of nitrogen from North America, *Atmos. Environ.*, 34, 2205-2230, 2000.
- Helmig, D., Tanner, D. M., Honrath, R. E., Owen, R. C., and Parrish, D. D.: Nonmethane hydrocarbons at Pico Mountain, Azores: 1. Oxidation chemistry in the North Atlantic region, *J. Geophys Res.*, 113, D20S91, 2008.
- Honrath, R. E., Helmig, D., Owen, R. C., Parrish, D. D., and Tanner, D. M.: Nonmethane hydrocarbons at Pico Mountain, Azores: 2. Event-specific analyses of the impacts of mixing and photochemistry on hydrocarbon ratios, *J. Geophys Res.*, 113, D20S92, 2008.
- Hurley, M. D., Sokolov, O., Wallington, T. J., Takekawa, H., Karasawa, Klotz, B., Barnes, I. A. N., and Becker, K. H.: Organic aerosol formation during the atmospheric degradation of toluene, *Environ. Sci. Technol.*, 35, 1358-1366, 2001.
- Iannone, R., Anderson, R. S., Rudolph, J., Huang, L., and Ernst, D.: The carbon kinetic isotope effects of ozone-alkene reactions in the gas-phase and the impact of ozone reactions on the stable carbon isotope ratios of alkenes in the atmosphere, *Geophys. Res. Lett.*, 30, 1684 10.1029/2003GL017221, 2003.
- Iannone, R., Anderson, R. S., Vogel, A., Eby, P. S., Whiticar, M. J., and Rudolph, J.: The hydrogen kinetic isotope effects of the reactions of n-alkanes with chlorine atoms in the gas phase, *J. Atmos. Chem.*, 50, 121-138, 2005.
- Iannone, R., Koppmann, R., and Rudolph, J.: Technique for atmospheric measurements of stable carbon isotope ratios of isoprene, methacrolein, and methyl vinyl ketone, *J. Atmos. Chem.*, 58, 181-202, 2007.

- Iannone, R., Koppmann, R., and Rudolph, J.:  $^{12}\text{C}/^{13}\text{C}$  kinetic isotope effects of the gas-phase reactions of isoprene, methacrolein, and methyl vinyl ketone with OH radicals, *Atmos. Environ.*, 43, 3103-3110, 2009.
- Iannone, R., Koppmann, R., and Rudolph, J.: Stable carbon kinetic isotope effects for the production of methacrolein and methyl vinyl ketone from the gas-phase reactions of isoprene with ozone and hydroxyl radicals, *Atmos. Environ.*, 44, 4135-4141, 2010.
- Irei, S., Huang, L., Collin, F., Zhang, W., Hastie, D., and Rudolph, J.: Flow reactor studies of the stable carbon isotope composition of secondary particulate organic matter generated by OH-radical-induced reactions of toluene, *Atmos. Environ.*, 40, 5858-5867, 2006.
- Jang, M., and Kamens, R. M.: Characterization of secondary aerosol from the photooxidation of toluene in the presence of  $\text{NO}_x$  and 1-propene, *Environ. Sci. Technol.*, 35, 3626-3639, 2001.
- Jenkin, M. E., and Clemitshaw, K. C.: Ozone and other secondary photochemical pollutants: chemical processes governing their formation in the planetary boundary layer, *Atmos. Environ.*, 34, 2499-2527, 2000.
- Jobson, B. T., Parrish, D. D., Goldan, P., Kuster, W., Fehsenfeld, F. C., Blake, D. R., Blake, N. J., and Niki, H.: Spatial and temporal variability of nonmethane hydrocarbon mixing ratios and their relation to photochemical lifetime, *J. Geophys. Res.*, 103, 13,557-13,567, 1998.
- Jordan, C., Fitz, E., Hagan, T., Sive, B., Frinak, E., Haase, K., Cottrell, L., Buckley, S., and Talbot, R.: Long-term study of VOCs measured with PTR-MS at a rural site in New Hampshire with urban influences, *Atmos. Chem. Phys.*, 9, 4677-4697, 2009.
- Kawashima, H., and Murakami, M.: Measurement of the stable carbon isotope ratio of atmospheric volatile organic compounds using chromatography, combustion, and isotope ratio mass spectrometry coupled with thermal desorption, *Atmos. Environ.*, 89, 140-147, 2014.
- Kleinman, L. I., Daum, P. H., Lee, L.-N., Nunnermacker, L. J., Springston, S. R., Weinstein-Lloyd, J., Hyde, P., Doskey, P., Rudolph, J., Fast, J., and Berkowitz, C.: Photochemical age determinations in the Phoenix metropolitan area, *J. Geophys. Res.*, 108, 4096, 2003.
- Kornilova, A., Saccon, M., O'Brien, J. M., Huang, L., and Rudolph, J.: Stable carbon isotope ratios and the photochemical age of atmospheric volatile organic compounds, *Atmosphere-Ocean*, 2013.
- Li, Q., Wang, W., Zhang, H.-W., Wang, Y.-J., Wang, B., Li, L., Li, H.-J., Wang, B.-J., Zhan, J., Wu, M., and Bi, X.-H.: Development of a compound-specific carbon isotope analysis

method for 2-methyltetrols, biomarkers for secondary organic aerosols from atmospheric isoprene, *Anal. Chem.*, 82, 6764-6769, 2010.

Lurmann, F. W., and Main, H. H.: Analysis of the ambient VOC data collected in the Southern California Air Quality Study, California Air Resources Board, 1992.

Matthews, D. E., and Hayes, J. M.: Isotope-ratio-monitoring gas chromatography-mass spectrometry, *Anal. Chem.*, 50, 1465-1473, 1978.

McKeen, S. A., and Liu, S. C.: Hydrocarbon ratios and photochemical history of air masses, *Geophys. Res. Lett.*, 20, 2363-2366, 1993.

McKeen, S. A., Liu, S. C., Hsie, E.-Y., Lin, X., Bradshaw, J. D., Smyth, S., Gregory, G. L., and Blake, D. R.: Hydrocarbon ratios during PEM-WEST A: A model perspective, *J. Geophys. Res.*, 101, 2087-2109, 1996.

Moukhtar, S., Saccon, M., Kornilova, A., Irei, S., Huang, L., and Rudolph, J.: Method for determination of stable carbon isotope ratio of methylnitrophenols in atmospheric particulate matter, *Atmos. Meas. Tech.*, 4, 2453-2464, 2011.

Niedojadlo, A., Kurtenbach, R., and Wiesen, P.: How reliable are emission inventories? Field observations versus emission predictions for NMVOCs, Springer, Netherlands, 2008.

Norman, A. L., Hopper, J. F., Blanchard, P., Ernst, D., Brice, K., Alexandrou, N., and Klouda, G.: The stable carbon isotope composition of atmospheric PAHs, *Atmos. Environ.*, 33, 2807-2814, 1999.

Parrish, D. D., Stohl, A., Forster, C., Atlas, E. L., Blake, D. R., Goldan, P. D., Kuster, W. C., and de Gouw, J. A.: Effects of mixing on evolution of hydrocarbon ratios in the troposphere, *J. Geophys. Res.*, 112, D10S34, 2007.

Piccot, S. D., Watson, J. J., and Jones, J. W.: A global inventory of volatile organic compound emissions from anthropogenic sources, *J. Geophys. Res.*, 97, 9897-9912, 1992.

Redeker, K. R., Davis, S., and Kalin, R. M.: Isotope values of atmospheric halocarbons and hydrocarbons from Irish urban, rural, and marine locations, *J. Geophys. Res.*, 112, D16307, 2007.

Reimann, S., and Lewis, A.: Anthropogenic VOCs, in: *Volatile Organic Compounds in the Atmosphere*, edited by: Koppmann, R., Blackwell Publishing Ltd, Oxford, 2007.

Roberts, J. M., Fehsenfeld, F. C., Liu, S. C., Bollinger, M. J., Hahn, C., Albritton, D. L., and Sievers, R. E.: Measurements of aromatic hydrocarbon ratios and NO<sub>x</sub> concentrations in the rural troposphere: Observation of air mass photochemical aging and NO<sub>x</sub> removal, *Atmos. Environ.* (1967), 18, 2421-2432, 1984.

- Rogers, K. M., and Savard, M. M.: Detection of petroleum contamination in river sediments from Quebec City region using GC-IRMS, *Org. Geochem.*, 30, 1559-1569, 1999.
- Rudolph, J., and Johnen, F. J.: Measurements of light atmospheric hydrocarbons over the Atlantic in regions of low biological activity, *J. Geophys. Res.*, 95, 20,583-20,591, 1990.
- Rudolph, J., Lowe, D. C., Martin, R. J., and Clarkson, T. S.: A novel method for the compound specific determination of  $\delta^{13}\text{C}$  in volatile organic compounds at ppt levels in ambient air, *Geophys. Res. Lett.*, 24, 659-662, 1997.
- Rudolph, J., and Czuba, E.: On the use of isotopic composition measurements of volatile organic compounds to determine the "photochemical age" of an air mass, *Geophys. Res. Lett.*, 27, 3865-3868, 2000.
- Rudolph, J.: Tropospheric Chemistry and Composition: Aliphatic hydrocarbons, in: *Encyclopedia Atmospheric Sciences*, edited by: Holton, J. R., Pyle, J., and Curry, J. A., Academic Press, London, UK, 2355-2364, 2002.
- Rudolph, J., Czuba, E., Norman, A. L., Huang, L., and Ernst, D.: Stable carbon isotope composition of nonmethane hydrocarbons in emissions from transportation related sources and atmospheric observations in an urban atmosphere, *Atmos. Environ.*, 36, 1173-1181, 2002.
- Rudolph, J., Anderson, R. S., Czapiewski, K. V., Czuba, E., Ernst, D., Gillespie, T., Huang, L., Rigby, C., and Thompson, A. E.: The stable carbon isotope ratio of biogenic emissions of isoprene and the potential use of stable isotope ratio measurements to study photochemical processing of isoprene in the atmosphere, *J. Atmos. Chem.*, 44, 39-55, 2003.
- Rudolph, J.: Gas Chromatography-Isotope Ratio Mass Spectrometry, in: *Volatile Organic Compounds in the Atmosphere*, edited by: Koppmann, R., Blackwell Publishing Ltd, Oxford, 2007.
- Saito, T., Tsunogai, U., Kawamura, K., Nakatsuka, T., and Yoshida, N.: Stable carbon isotopic compositions of light hydrocarbons over the western North Pacific and implication for their photochemical ages, *J. Geophys. Res.*, 107, ACH 2-1–ACH 2-9, 2002.
- Saito, T., Kawamura, K., Tsunogai, U., Chen, T.-Y., Matsueda, H., Nakatsuka, T., Gamo, T., Uematsu, M., and Huebert, B. J.: Photochemical histories of nonmethane hydrocarbons inferred from their stable carbon isotope ratio measurements over east Asia, *J. Geophys. Res.*, 114, D11303, 2009.
- Santrock, J., Studley, S. A., and Hayes, J. M.: Isotopic analyses based on the mass spectra of carbon dioxide, *Anal. Chem.*, 57, 1444-1448, 1985.

- Sawyer, R. F., Harley, R. A., Cadle, S. H., Norbeck, J. M., Slott, R., and Bravo, H. A.: Mobile sources critical review: 1998 NARSTO assessment, *Atmos. Environ.*, 34, 2161-2181, 2000.
- Smallwood, B. J., R. Philp, P., and Allen, J. D.: Stable carbon isotopic composition of gasolines determined by isotope ratio monitoring gas chromatography mass spectrometry, *Org. Geochem.*, 33, 149-159, 2002.
- Smith, D. F., McIver, C. D., and Kleindienst, T. E.: Primary product distribution from the reaction of hydroxyl radicals with toluene at ppb NO<sub>x</sub> mixing ratios (vol 30, pg 209, 1998), *J. Atmos. Chem.*, 31, 349-350, 1998.
- Stein, O., and Rudolph, J.: Modelling and interpretation of stable carbon isotope ratios of ethane in global chemical transport models, *J. Geophys. Res.*, 112, D14308, 2007.
- Thompson, A.: Stable carbon isotope ratios of nonmethane hydrocarbons and halocarbons in the atmosphere. Doctor of Philosophy Thesis, York University, 2003.
- Thompson, A., Rudolph, J., Rohrer, F., and Stein, O.: Concentration and stable carbon isotopic composition of ethane and benzene using a global three-dimensional isotope inclusive chemical tracer model, *J. Geophys. Res.-Atmos.*, 108, D13, 1029/2002JD002883, 2003.
- Turner, N., Jones, M., Grice, K., Dawson, D., Ioppolo-Armanios, M., and Fisher, S.: d<sup>13</sup>C of volatile organic compounds (VOCS) in airborne samples by thermal desorption-gas chromatography-isotope ratio-mass spectrometry (TD-GC-IR-MS), *Atmos. Environ.*, 40, 3381-3388, 2006.
- Wintel, J., Hösen, E., Koppmann, R., Krebsbach, M., Hofzumahaus, A., and Rohrer, F.: Stable carbon isotope ratios of toluene in the boundary layer and the lower free troposphere, *Atmos. Chem. Phys.*, 13, 11059-11071, 2013.
- Zuiderweg, A. T., Holzinger, R., and Roeckmann, T.: Analytical system for stable carbon isotope measurements of low molecular weight (C2-C6) hydrocarbons, *Atmos. Meas. Tech.*, 4, 1161-1175, 2011.

**Table 1.** Averages of background signals for cartridges containing  $(1 \pm 0.2)$  g of carboxene and 0.2 g to 0.5 g of quartz wool.

Compound	Blank $\pm$ SD <sup>a</sup> (ng)	DL <sup>b</sup> (ng)	Average mass per ambient cartridge (ng) <sup>c</sup>
n-Hexane	$0.2 \pm 0.3$	1.0	146
Benzene	$0.7 \pm 0.5$	1.5	16
n-Heptane	$1.1 \pm 0.1$	0.3	64
Toluene	$0.2 \pm 0.1$	0.4	61
Ethylbenzene	$0.2 \pm 0.2$	0.6	7
p,m-Xylene	$0.2 \pm 0.2$	0.5	17
o-Xylene	$0.2 \pm 0.1$	0.4	5

<sup>a</sup> SD: Standard deviation calculated from at least 4 repeat measurements.

<sup>b</sup>  $3\sigma$  detection limit (DL).

<sup>c</sup> average calculated from 50 ambient samples.

**Table 2.** Recoveries of VOC (%) under different desorption conditions (samples were analysed with GC-FID or GC-IRMS)<sup>a</sup>.

Test#	1	2	3	4	5	6	7	8
Duration (min)	15	15	15	20	25	30	30	40
Temperature (K)	523	553	573	673	573	573	583	553
Flow rate (mL/min)	55	53	54	78	58	58	53	60
n-Hexane	70	107	164	102	106	104	102	102
Benzene	73	105	148	104	106	105	108	113
n-Heptane	39	93	102	99	96	97	96	100
Toluene	43	96	110	103	100	100	102	103
Ethylbenzene	29	74	77	86	87	86	87	90
p,m-Xylene	24	65	67	82	90	87	97	95
o-Xylene	25	68	64	81	85	84	92	94

<sup>a</sup> Results are based on one measurement.

**Table 3.** Average recoveries<sup>a</sup> of VOC sampled on cartridges for desorption at 553 K for 40 min at (60 to 80) mL min<sup>-1</sup> carrier gas flow rate.

Compounds	Mass of loaded VOC (ng)	Average Recovery <sup>b</sup> %	SD <sup>c</sup> of delta values ‰
Benzene	25	112 ± 2 <sup>c</sup>	0.6
Heptane	50	93 ± 9	0.7
Toluene	60	102 ± 14	0.6
Ethylbenzene	50	84 ± 13	0.3
p,m-Xylene	75	99 ± 20	0.6
o-Xylene	35	94 ± 5	0.5

<sup>a</sup> Values were not blank corrected.

<sup>b</sup> Based on 3 to 4 measurements.

<sup>c</sup> SD: Standard deviation.

**Table 4.** Concentrations and delta values determined from ambient samples collected at Egbert.

Compound	Concentrations (ppbV) (ambient)			$\delta^{13}\text{C}_{\text{ambient}}$ (‰)			$\delta^{13}\text{C}_{\text{source}}$ (‰) <sup>a</sup>	
	Median	10 <sup>th</sup> percentile	90 <sup>th</sup> percentile	Median	10 <sup>th</sup> percentile	90 <sup>th</sup> percentile	Average	SD <sup>b</sup>
n-Hexane	1.53	0.61	5.18	-27.0	-28.1	-24.0	-26.7	0.4
Benzene	0.10	0.05	0.24	-25.6	-27.9	-22.2	-28.0	1.1
n-Heptane	0.40	0.14	1.80	-24.2	-26.1	-21.7	-26.4	1.1
Toluene	0.13	0.06	0.25	-24.8	-26.7	-22.5	-27.6	0.6
Ethylbenzene	0.02	0.01	0.06	-23.9	-27.5	-18.6	-27.7	0.4
p,m-Xylene	0.03	0.01	0.11	-23.8	-27.7	-19.8	-27.4	0.5
o-Xylene	0.01	0.004	0.03	-23.5	-26.2	-20.7	-27.2	0.1

<sup>a</sup> Stable carbon isotope composition of the sources based on Rudolph et al. (2002).

<sup>b</sup> SD: Standard deviation

## Figure Captions

**Fig. 1.** Schematic diagram of the Two Stage Preconcentration System (TSPS)

**Fig. 2.** Calibration curve for toluene (based on TSPS-GC-IRMS measurements, slope =  $1.46 \times 10^{10} \pm 0.03 \times 10^{10}$  ions  $\text{ng}^{-1}$ ,  $R^2 = 0.997$ )

**Fig.3.** Plot of  $\delta^{13}\text{C}$  values versus injected mass measured online using TSPS-GC-IRMS. Dashed lines are the reference (offline)  $\delta^{13}\text{C}$  values. For p,m-xylene the average of the p-xylene and m-xylene offline values is given since the mixtures contained equal masses of both xylenes.

**Fig. 4.** Impact of storage on sample recovery for storage of 1 to 2 weeks at ambient temperatures (diamonds) and 6 to 7 months in a freezer (circles). The recovered masses are given relative to samples analysed immediately after sampling. Both storage tests used the same reference values. Values used are the average of 3 to 4 measurements and the error bars show the variability of results for analysis after storage. Masses loaded range from 25 ng to 75 ng. Values were not corrected for blanks.

**Fig. 5.** Change in  $\delta^{13}\text{C}$  of VOC collected on cartridges during storage for 1 to 2 weeks at ambient temperatures (diamonds) and 6 to 7 months in a freezer (circles). The differences are given relative to the  $\delta^{13}\text{C}$  values of samples analysed without storage. Both storage tests used the same reference values. Values used are the average of 3 to 4 measurements and the error bars show the variability of results for analysis after storage. Masses loaded range from 25 ng to 75 ng. Values were not corrected for blanks.

**Fig.6.** Chromatogram obtained by GC-IRMS for a test mixture with 11 compounds (A) and an ambient sample (bottom B). Rectangular peaks are reference  $\text{CO}_2$  signals and peaks numbered from 1 to 11 correspond to n-pentane, n-hexane, benzene, n-heptane, toluene, n-octane, ethylbenzene, p,m-xylene, o-xylene, n-nonane and n-decane.

**Fig. 7.** Mixing ratio versus stable carbon isotope ratio plots for samples collected at Egbert. The range (average  $\pm \sigma$ ) of isotope ratios reported by Rudolph et al. (2002) for urban sources are shown as dashed lines.

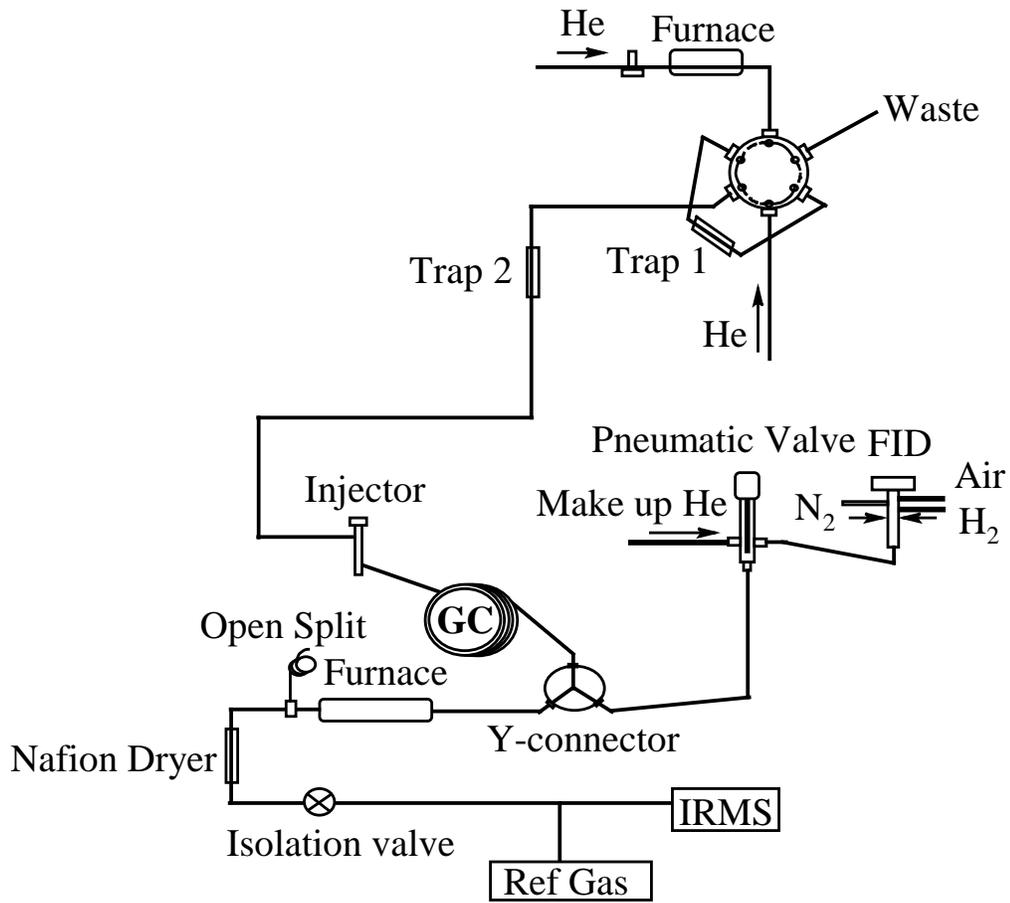


Fig. 1

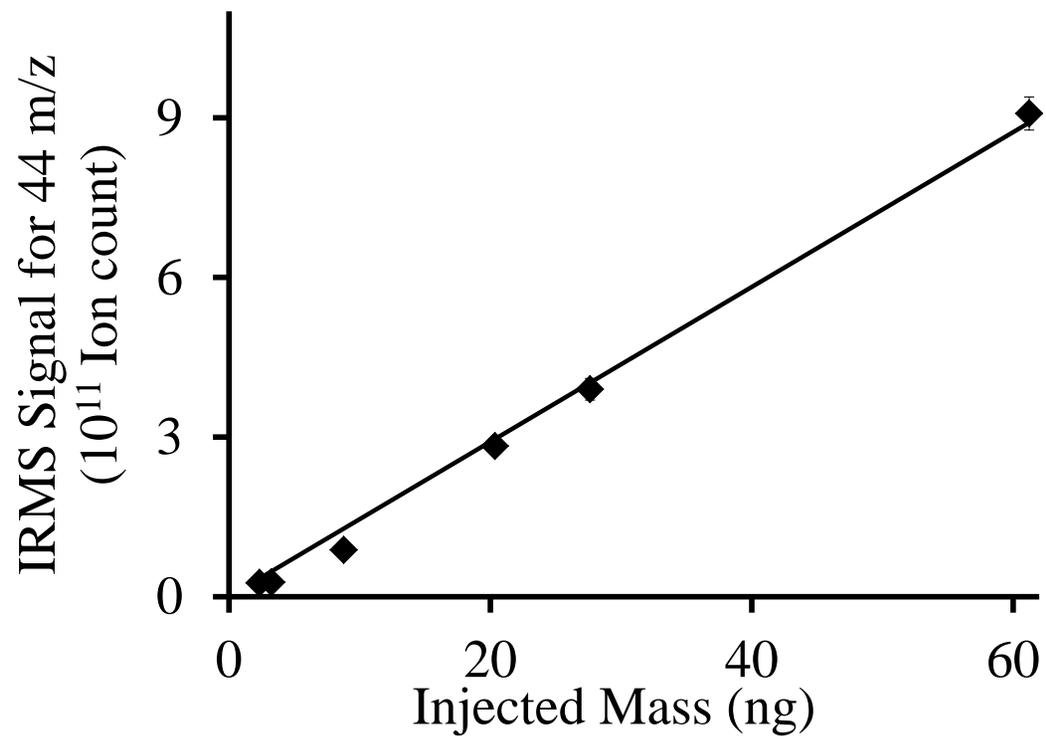
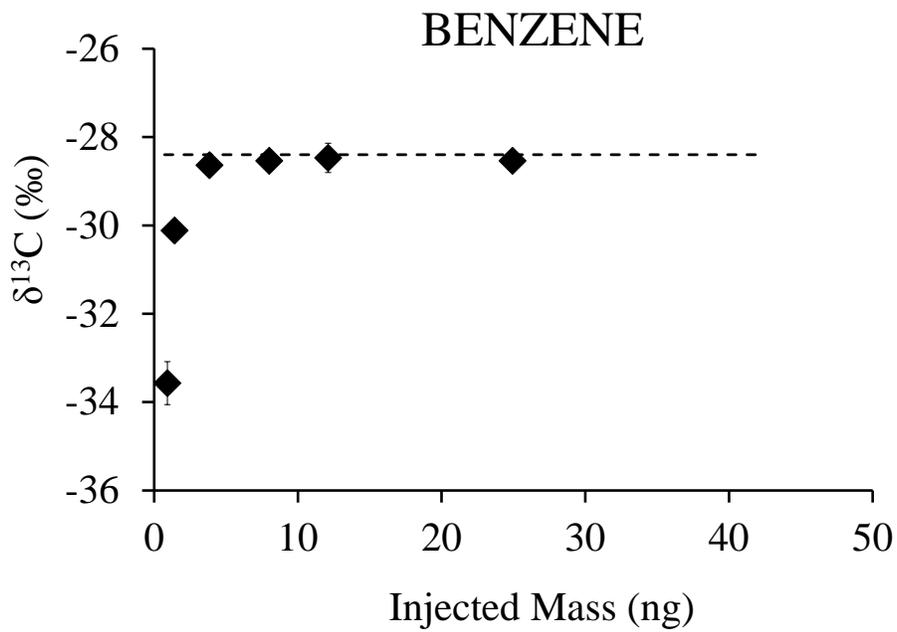
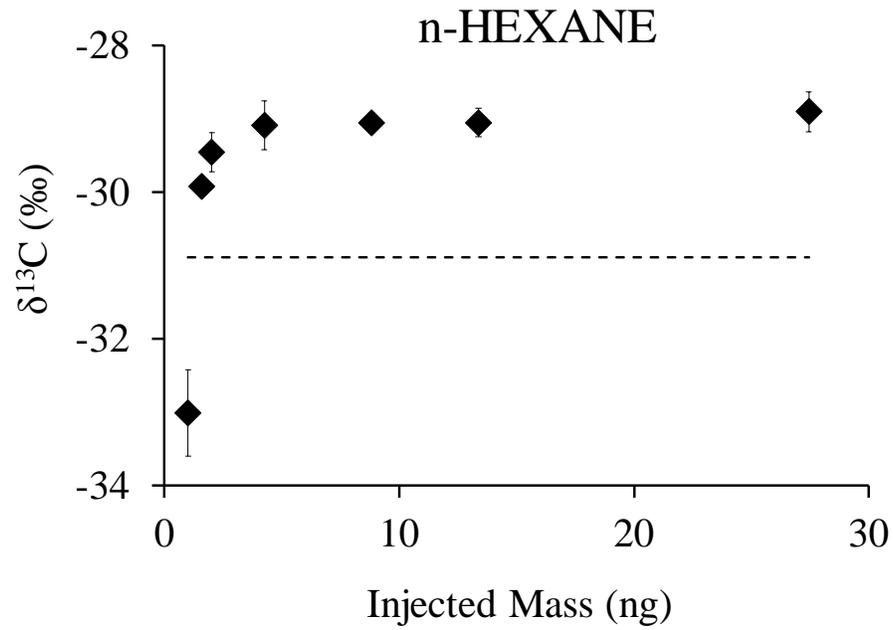
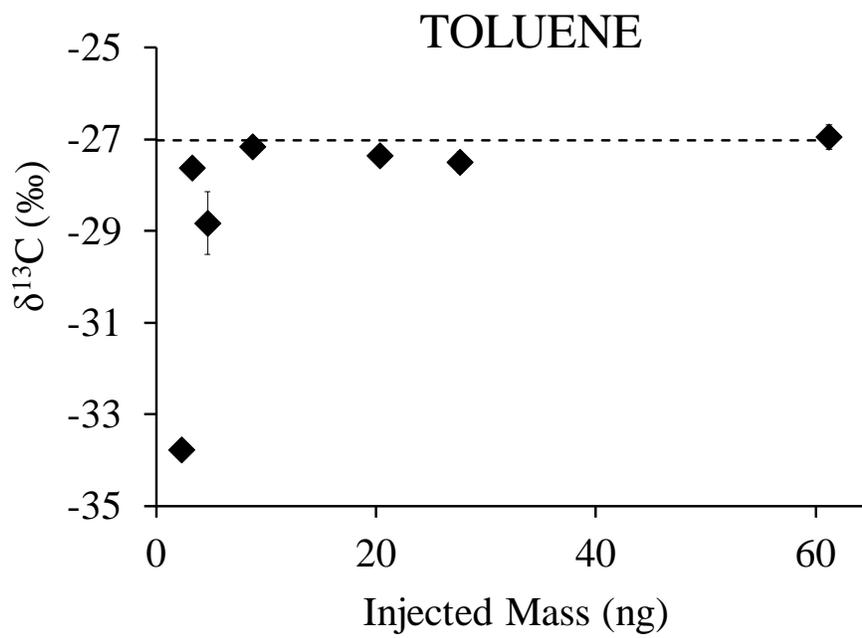
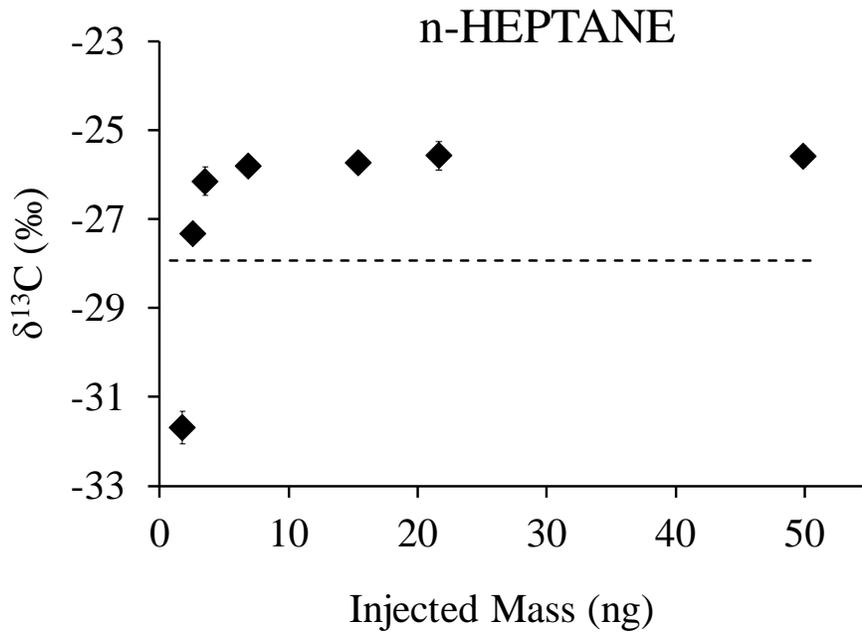
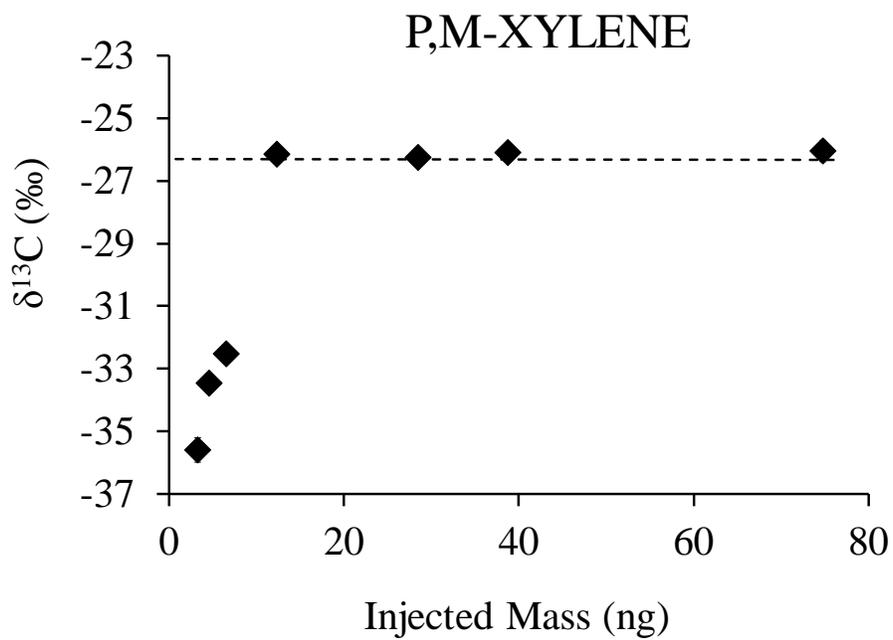
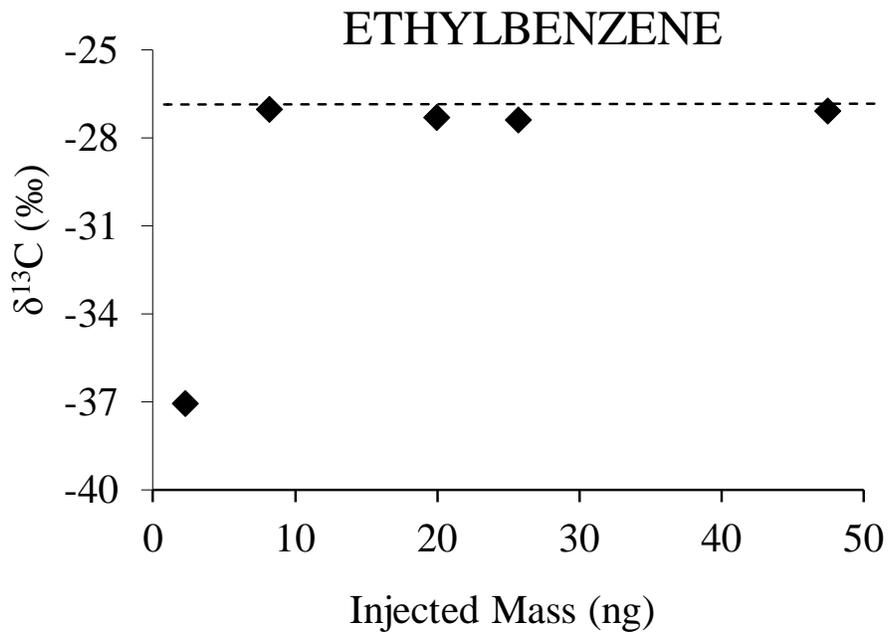


Fig. 2







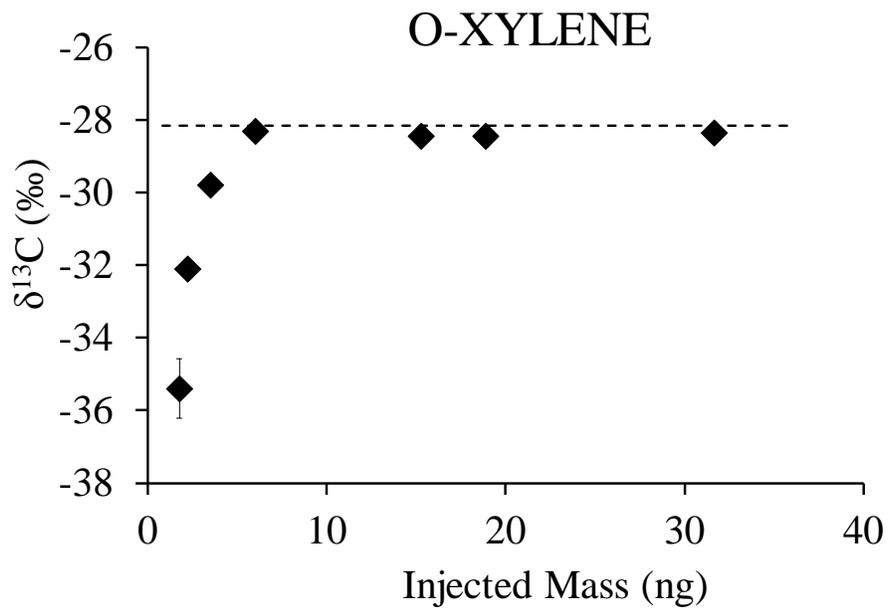


Fig. 3

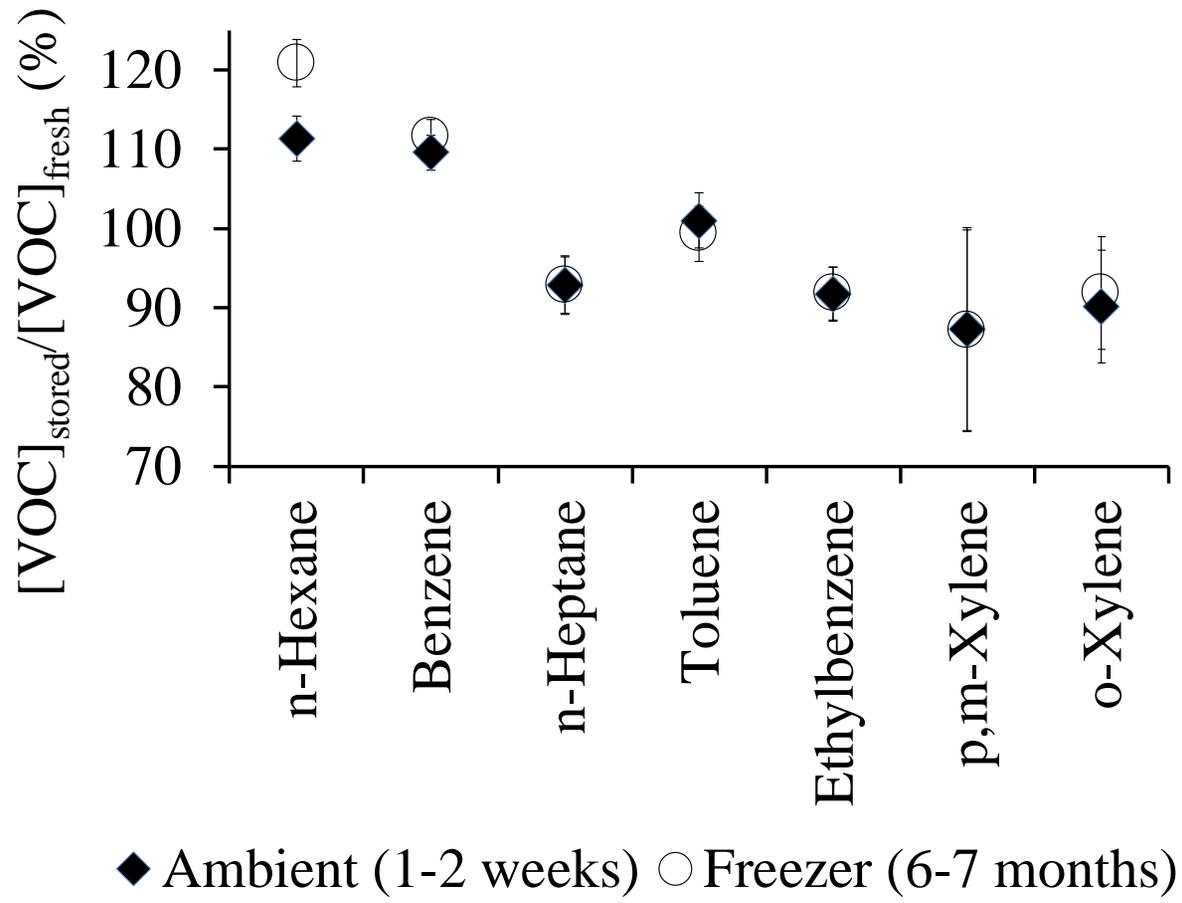


Fig. 4

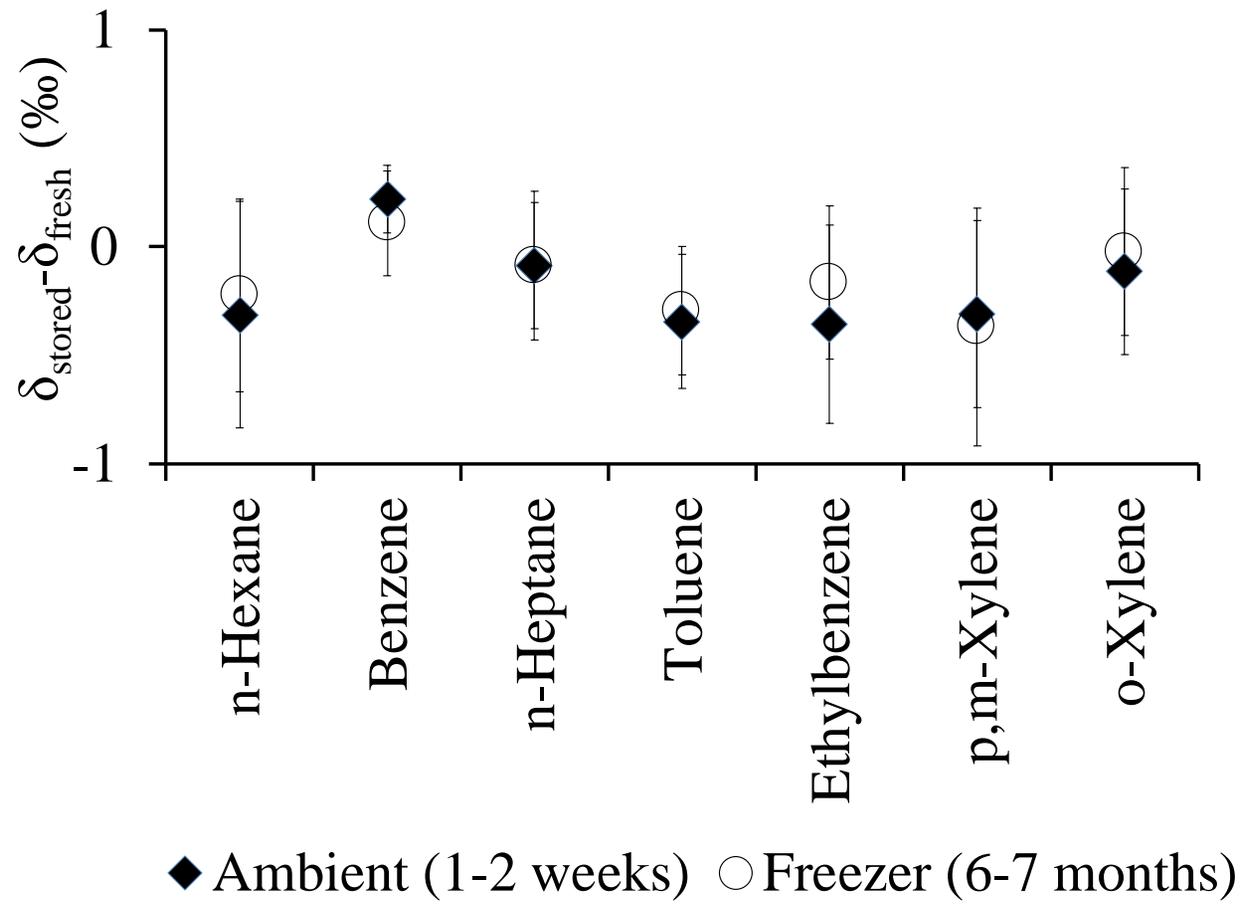
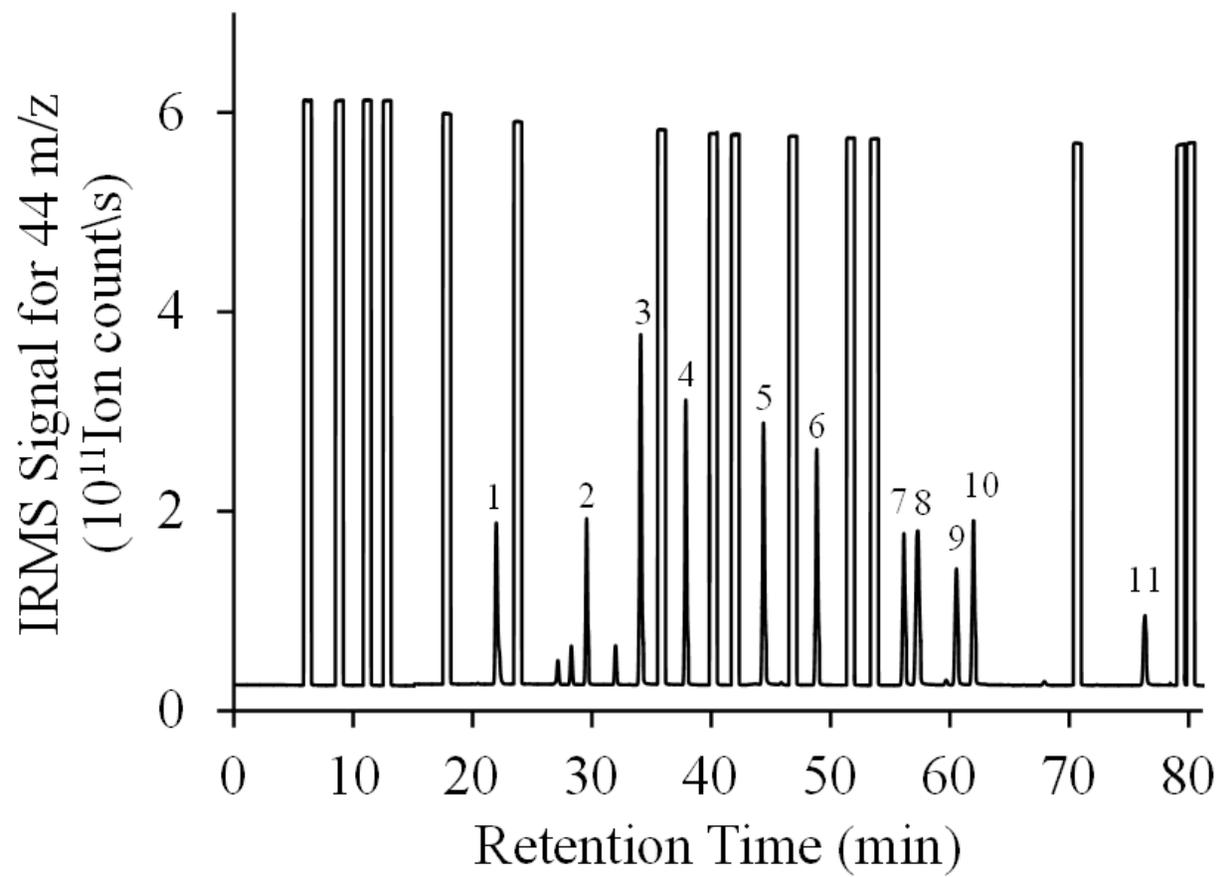


Fig. 5

A



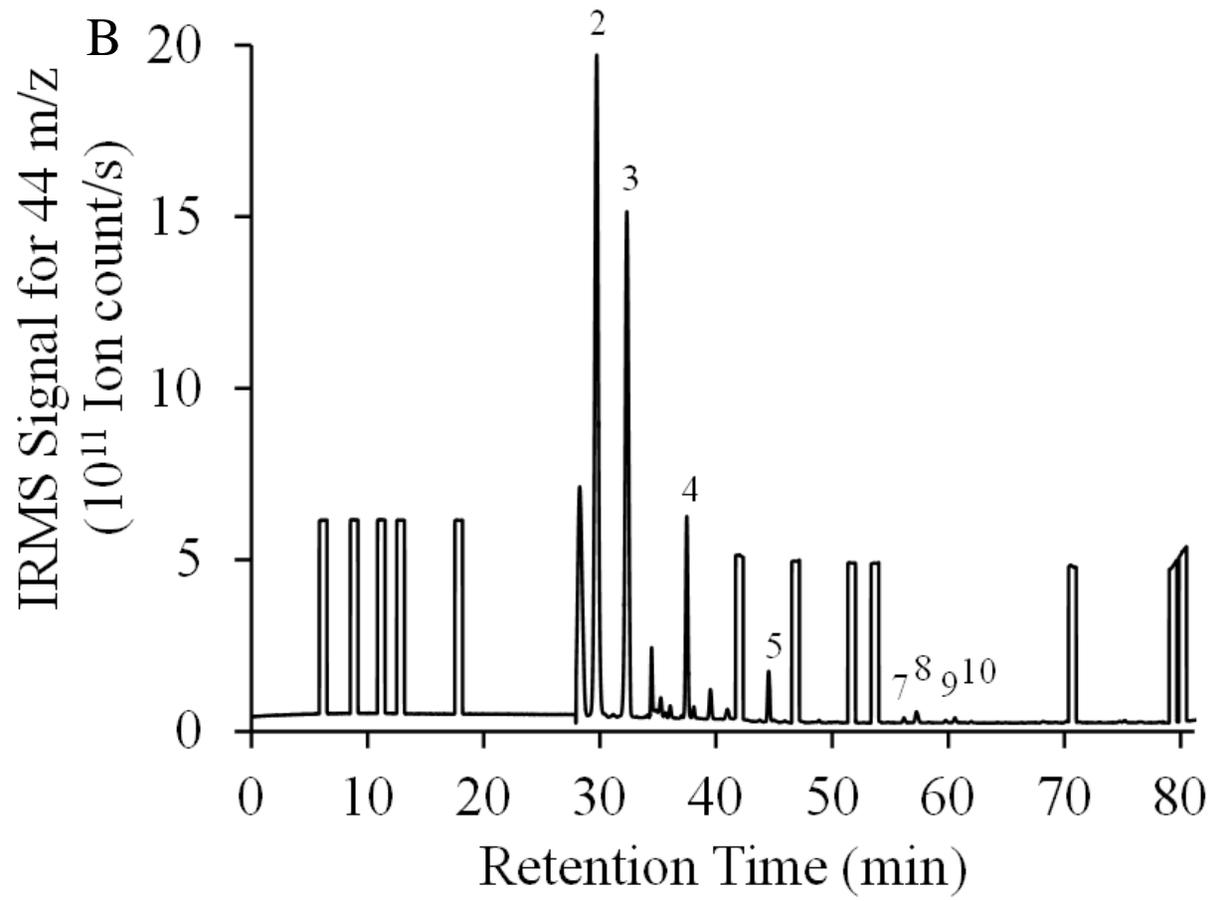
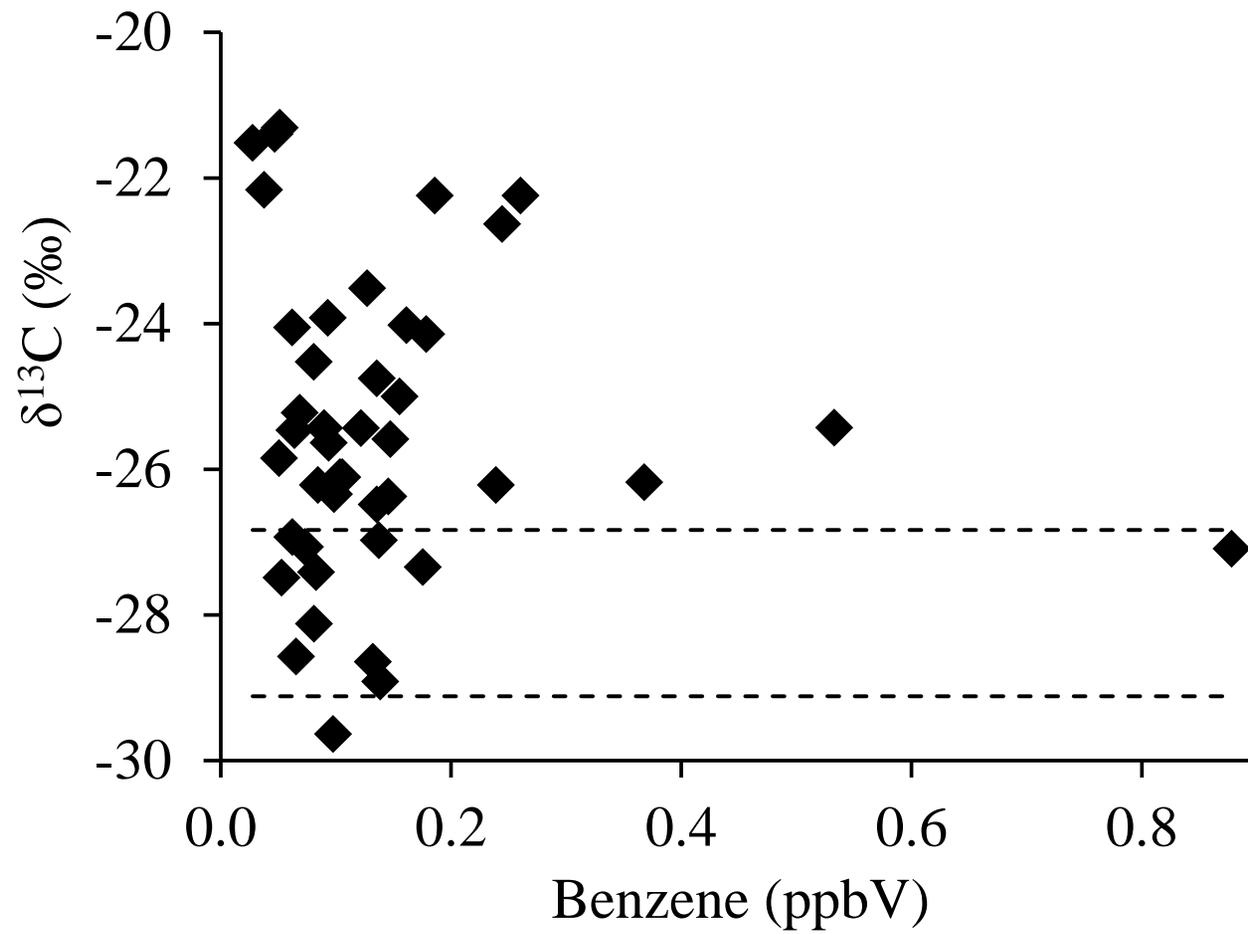
**B**

Fig. 6



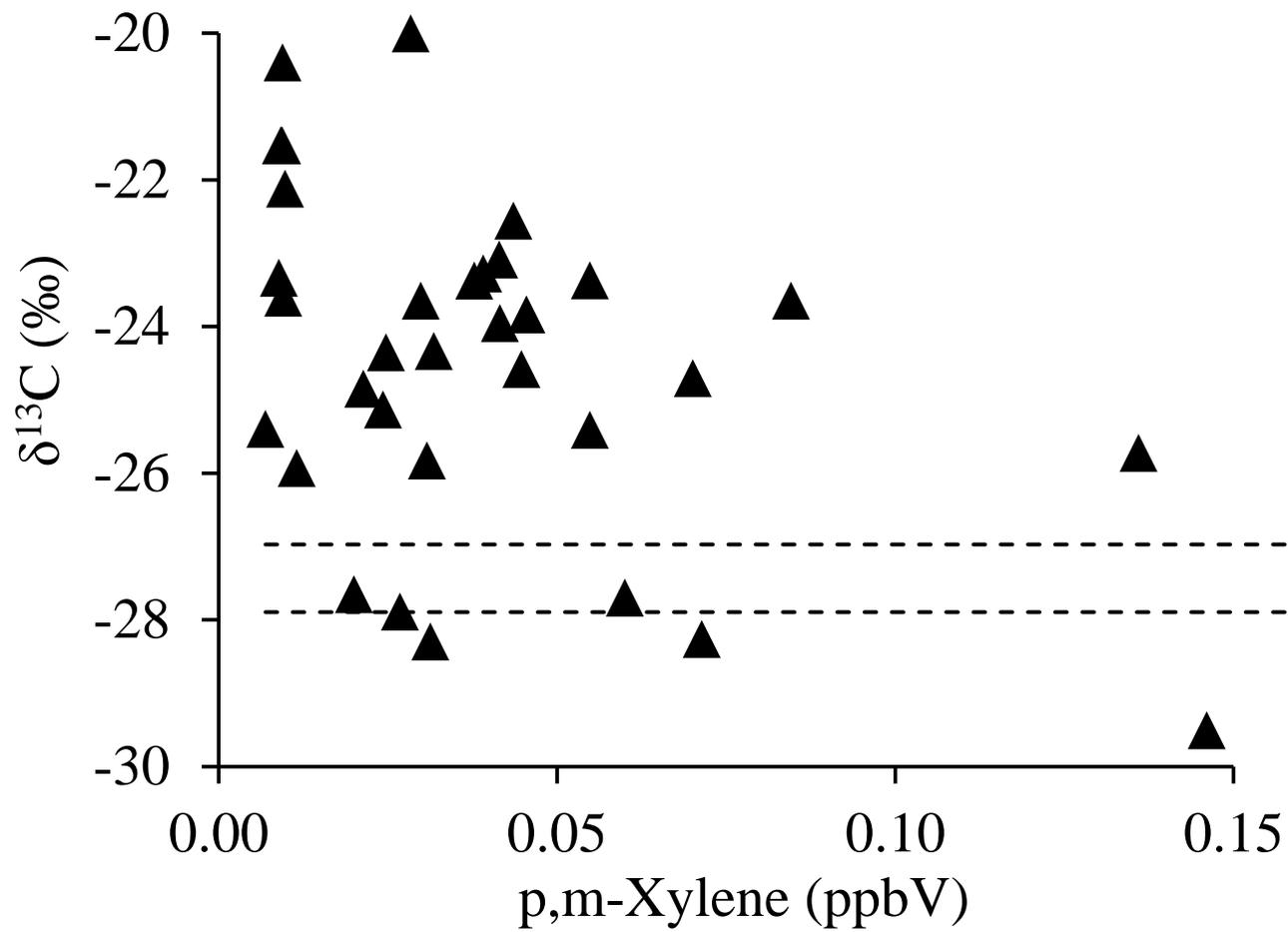


Fig. 7