Results from the first national UK inter-laboratory calibration for very short-lived halocarbons

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The first UK inter-laboratory calibration for VSLH

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Abstract

Very short-lived halocarbons (VSLH) such as CH$_3$I, CH$_2$Br$_2$ and CHBr$_3$ provide an important source of reactive halogens to the atmosphere, however high spatial and seasonal variability in their ambient mixing ratios and sea-air fluxes gives rise to considerable uncertainty in global scale emission estimates. One solution to improve global flux estimates is to combine the multitude of individually published datasets to produce a database of collated global halocarbon observations. Some progress towards this has already been achieved through the HalOcAt (Halocarbons in the Ocean and Atmosphere) database initiative, however the absence of a common calibration scale for very short-lived halocarbons makes it difficult to distinguish true environmental variations from artefacts arising from differences between calibration methodologies. As such, the lack of inter-calibrations for both air and seawater measurements of very short-lived halocarbons has been identified as a major limitation to current estimations of the global scale impact of these reactive trace gases. Here we present the key findings from the first national UK inter-laboratory comparison for calibrations of the halocarbons CH$_3$I, CH$_2$Br$_2$ and CHBr$_3$. The aim of this inter-calibration was to provide transparency between halocarbon calibrations from major UK research institutions, an important step towards enabling all measurements from these institutions to be treated as one coherent integrated dataset for global source term parameterisations.

1 Introduction

Very short-lived halocarbons or VSLH (defined as those with atmospheric lifetimes of less than six months, Law and Sturges, 2007) of predominantly biogenic marine origin are thought to supply a substantial fraction of reactive halogens to the marine boundary layer (MBL) (Carpenter et al., 2003; Jones et al., 2010), free troposphere (von Glasow et al., 2004; Yang et al., 2005) and lower stratosphere (Garcia and Solomon, 1994; Sturges et al., 2000; Liang et al., 2010 and references therein), potentially influencing both stratospheric and tropospheric photochemistry and climate.
Photo-destruction of volatile organic iodine- and bromine-containing trace gases to release I and Br atoms within the troposphere can initiate catalytic ozone depletion and impact upon NO/NO$_2$ and OH/HO$_2$ chemistry, which in turn affects the atmospheric lifetimes of other climatically important trace gases (Vogt et al., 1999; Bloss et al., 2005; Read et al., 2008; Whalley et al., 2010; Mahajan et al., 2010).

While the majority of organic iodine compounds (or iodocarbons) are broken down within the troposphere, the relatively longer lived bromocarbons (e.g. CH$_2$Br$_2$ and CHBr$_3$, with respective lifetimes of 3–4 months and ~3–5 weeks; Quack et al., 2007) are thought to contribute ~5–40% of stratospheric bromine (WMO report, 2010). Bromine-catalysed stratospheric ozone depletion is on average ~60 times more efficient than chlorine-initiated ozone destruction (Law and Sturges, 2007), yet to date sources of stratospheric bromine have been less well quantified.

The number of published VSLH datasets has increased substantially in recent years (including Quack and Wallace, 2003 and references therein; Chuck et al., 2005; Archer et al., 2007; Carpenter et al., 2007, 2009; Quack et al., 2007; Laube et al., 2008; Varner et al., 2008; O’Brien et al., 2009; Jones et al., 2010; Kurihara et al., 2010), and includes atmospheric and oceanic measurements, and sea-air flux estimates. However, there remains considerable uncertainty associated with global source term estimates for these gases (e.g. annual global CHBr$_3$ source estimated as ~10 Gmol Br yr$^{-1}$, but including the quantifiable uncertainties gives the range 3–22 Gmol Br yr$^{-1}$; Quack and Wallace, 2003). This high level of uncertainty is the result of several factors, including the spatial and seasonal variability in biological production and sea-air transport rates, their short tropospheric lifetimes, the sparse availability of data which may be considered representative of regional sources (i.e. made in the free troposphere or over the open ocean), and the lack of a common calibration scale for these gases. Improved quantification of the global emission budget for these short-lived halocarbons requires an integrated approach, combining as many datasets as possible, in order to maximise the temporal and spatial resolution.
The Halocarbons in the Ocean and Atmosphere (HalOcAt) database project is a current SOLAS/COST initiative hosted by IFM-GEOMAR, which aims to collate atmospheric and oceanic VSLH observations from across the international research community (https://halocat.ifm-geomar.de). However, the lack of formal comparisons between the various calibration techniques used by different research groups to quantify halocarbons has been identified as one of the major limitations to effectively utilising this type of integrated dataset to improve global emissions estimates. Indeed it is currently difficult to discern with any certainty whether observational differences are naturally occurring, or are simply artefacts of differences in measurement and calibration techniques. Given the potential benefits of compiling a coherent global database for measurements of very short-lived halocarbons within the atmosphere and ocean, there is a general consensus within the community for the need to link individual datasets to a common calibration scale (Butler et al., 2010). Such calibration scales are already in place for longer lived atmospheric trace gases, such as CH$_4$ and CO$_2$, however the comparative instability of shorter-lived gases in metal canisters for prolonged periods (Finlayson-Pitts and Pitts, 2000), coupled with the very limited number of stability studies, makes large scale intercalibrations for these gases less straightforward.

A meeting of international scientists from the halocarbon measurement community in London in February 2008 highlighted the need for more formal inter-laboratory calibrations for these short-lived gases (Butler et al., 2010). Whilst the ultimate goal would be to undertake an international inter-calibration, encompassing all research laboratories that routinely monitor very short-lived halocarbons, funding and logistical constraints mean that to date this has not been achieved. However, the UK halocarbon measurement community has taken an initial step towards this goal, and we report here the key findings from the first formal UK inter-laboratory comparison of calibrations for very short-lived halocarbons.
2 Methods

Research groups from the following institutions participated in the UK inter-laboratory comparison of calibrations for very short-lived halocarbons, as part of the UK National Environment Research Council (NERC) Quantifying and Understanding the Earth System (QUEST) initiative, between June and November, 2010:
- The University of Bristol (UoB, Mace Head Atmospheric Station, Ireland)
- The University of Cambridge (UoC)
- The University of East Anglia (UEA)
- Plymouth Marine Laboratory (PML)
- The University of York (UoY)

2.1 The inter-calibration gas standard

Each laboratory’s existing calibration technique was compared to the NOAA (National Oceanic and Atmospheric Administration, US) calibration scales, through the exchange of a compressed gas standard, SX-3570. The multi-component gas standard (in an electropolished stainless steel canister, Essex Cryogenics, St. Louis, MO, US) containing both VSLH and longer-lived halocarbons in modified continental background air (from Niwot Ridge, Colorado) was prepared and analysed by the NOAA Earth System Research Laboratory (ESRL) Global Monitoring Division in Boulder, Colorado. For those halocarbons with very low background air concentrations, the standard was spiked with additional volumes of high concentration halocarbons, in order to generate mixing ratios $\sim$2–5 ppt above ambient levels. Mixing ratios assigned by NOAA for the VSLH targeted for the comparison ($\text{CH}_3\text{I}$, $\text{CH}_2\text{Br}_2$ and $\text{CHBr}_3$) are given in Table 1. The SX-3570 gas standard also contains other short-lived polyhalogenated methanes, including $\text{CH}_2\text{I}_2$, $\text{CH}_2\text{ICl}$ and $\text{CHBrCl}_2$, however the stability of these gases in canisters has not been well characterised, and hence only provisional mixing ratios were assigned for these compounds.
2.2 Calibration and analysis methods

Details of the analytical systems and calibration techniques used by each group during the inter-comparison are outlined below.

2.2.1 University of Bristol (Mace Head)

All comparisons were carried out using the Medusa-MS (custom pre-concentration device coupled to an Agilent 6890 GC and Agilent 5973 MSD). 2l samples were pre-concentrated on a trap held at −150°C to −170°C and were subsequently analysed using a single main capillary chromatography column (CP-PoraBOND Q, 0.32 mm ID x 25 m, 5 µm, Varian Chrompack) with temperature and pressure ramping (Miller et al., 2008).

The MS was operated in EI-SIM (Electron Impact-Selected Ion Monitoring) mode using a single target ion and one or two qualifier ions as a check for the correct ion ratios. The starting column temperature was 40°C, this was ramped on injection of the sample to 200°C at 28°C min⁻¹. CH₃I was quantified using the CH₃I⁺ ion (m/z 142) with I⁺ (m/z 127) as a qualifier, whilst CH₂Br₂ was quantified via CH₂Br⁺² (m/z 174) with CH₂Br⁺ (m/z 93) and CH₂Br⁺ (m/z 95) as qualifier ions, and finally CHBr₃ on CH(²⁹Br)(³¹Br)⁺ (m/z 173) with CH(²⁹Br)(³¹Br)⁺ (m/z 171) as a qualifier ion.

The NOAA standard SX-3570 was compared against a tertiary standard (J-128), filled during relatively clean-air conditions at Trinidad Head using modified diving compressors (SA-6, RIX Industries), and measured against secondary standards at Scripps Institution of Oceanography (SIO) before and after usage at Mace Head. In the case of CH₂Br₂ and CHBr₃, preliminary calibrations values were assigned by comparison to a commercial gas standard prepared by Linde Gas Ltd. In the case of CH₃I, a preliminary calibration was assigned via comparison to a gas standard provided by Dr. Y. Yokouchi (National Institute for Environmental Studies, NIES, Japan). The precision of the SX-3570 to J-128 measurement was ≈2% for CH₃I, ≈0.7% for CH₂Br₂ and ≈1.2% for CHBr₃ (n = 10).
2.2.2 University of Cambridge

At Cambridge, the calibration comparison was carried out using a GC-ECD (gas chromatograph with electron capture detection) system similar to that described by Gostlow et al. (2010) but with improved sample pre-concentration (thermally regulated) and separation (longer column). The intercalibration standard SX-3570 was compared to the Cambridge working standard (NOAA cylinder no. SX3568) using a sequence of samples analysed over a 24 h period. The SX-3570 standard was treated as an unknown sample and the Cambridge working standard was used to achieve calibration. The sequence generated 8 sample chromatograms from analysis of SX-3570 and 9 calibration chromatograms of the Cambridge working standard (with sample volumes from 3–50 ml), to produce response curves which were used to calibrate the samples. Blank chromatograms were run following each sample or calibration chromatogram by passing helium carrier gas through the adsorbent bed (these served as a check on desorption efficiency and system impurities).

Each sample or calibration chromatogram was generated as follows: a flow (∼10 ml min⁻¹) of either sample or calibration air was passed through a dual bed adsorbent (1 mg each of Carboxen 1016/1001) held in an Ultimetal tube (1/32″ od×0.53 mm id, Varian) mounted across a 6 port, 2 position Valco valve set in the “load” position. The adsorbent tube was Peltier-cooled to 15°C, the target sample volume was 20 ml. Oxygen and residual moisture were purged from the adsorbent bed using dry helium before the Valco valve was switched to the “inject” position and the adsorbent tube was heated to 190°C for 15 s in the helium carrier flow. The column (Restek MXT 502.2, 20 m long, 0.18 m i.d., 1 µm film thickness) was held at 28°C for 6 min then heated to 132°C at 6°C per minute, the helium carrier pressure was increased to compensate for the increase in gas viscosity during the temperature program. The target peaks were directed into the ECD (Agilent micro volume model no. G2397-60510) running at 250°C. Target peaks in chromatograms from the intercalibration standard were calibrated using the response curve for each compound generated from the calibration
chromatograms using the Cambridge working standard. With the current chromatographic set-up, \( \text{CH}_2\text{Br}_2 \) and \( \text{CHBrCl}_2 \) co-elute (O’Brien et al., 2009). For this comparison, the SX-3570 \( \text{CHBrCl}_2 \) mixing ratio assigned by NOAA was subtracted from the combined value for the unresolved \( \text{CH}_2\text{Br}_2 \) and \( \text{CHBrCl}_2 \) peaks determined by Cambridge, in order to allow as close a comparison of the \( \text{CH}_2\text{Br}_2 \) mixing ratios as possible.

Measurement precision is calculated for each target compound as the standard deviation of the 8 replicates of the standard SX-3570. The uncertainty in the reported dry air mole fraction for each target compound is calculated directly from the estimated scale uncertainties from NOAA-ESRL: 10\% for \( \text{CHBr}_3 \) and 20\% for \( \text{CH}_3\text{I} \) and \( \text{CH}_2\text{Br}_2 \) (two standard deviations).

### 2.2.3 University of East Anglia

All samples were dried on-line prior to analysis using a glass tube filled with \( \text{Mg(ClO}_4\text{)}_2 \). Subsequently the condensable trace gases were pre-concentrated from ~250 ml of air at ~78°C in a 1/16” sample loop filled with an adsorbent (Hayesep D, 80/100 mesh) which was heated to 100°C immediately after injection. Separation was carried out using an Agilent 6890 gas chromatograph with an Agilent GS-GasPro column (length 30 m, ID 0.32 mm) coupled to a high sensitivity tri-sector (EBE) mass spectrometer from Micromass/Waters Corporation (AutoSpec Premier). The MS was operated in EI-SIR (Electron Impact-Selected Ion Recording) mode, at a mass resolution of 1000. The column temperature was held at −10°C for two minutes and then ramped from −10°C to 200°C at 10°C min\(^{-1}\) (Laube et al., 2010). \( \text{CH}_3\text{I} \) was quantified using the \( \text{I}^+ \) ion \((m/z \ 126.91)\) whilst \( \text{CH}_2\text{Br}_2 \) was quantified via \( \text{CH}_2(79\text{Br})^+ \) \((m/z \ 92.93)\) and \( \text{CHBr}_3 \) on \( \text{CH}(79\text{Br})(81\text{Br})^+ \) \((m/z \ 172.84)\).

The intercalibration standard SX-3570 was measured four times against an internal standard (Aculife-treated aluminium cylinder) filled with remote tropospheric air from Niwot Ridge near Boulder, USA. The latter was brought inline with NOAA calibration scales by measuring it against two tertiary standards (35 l electropolished stainless
VSLH in seawater were analysed based upon the methods outlined by Hughes et al. (2008). Calibration and quantification of VSLH were performed using laboratory-prepared liquid standards, by dilution of the pure compounds into HPLC-grade methanol. The primary standards were prepared gravimetrically, the secondary and working standard by serial dilution. Microlitre volumes of the working standards were injected into seawater, and these standards were used to achieve multi-point calibrations, taking into account individual purge efficiencies of the VSLH.

Halocarbons were extracted from the water phase by purging with ultra high purity (BIP) nitrogen at 90 ml min\(^{-1}\) for 10 min. Aerosols and moisture were removed from the purge gas stream using glass wool contained within a section of glass tubing, and a counterflow Nafion drier using oxygen-free nitrogen at 180 ml min\(^{-1}\). Halocarbons were trapped on triple-bed stainless steel solid sorbent tubes (Markes International Ltd.) containing Tenax, Carbograph and Carboxen. During sample collection, the tubes were held in a Peltier temperature-controlled block held at 1–2 °C. For the purposes of the intercalibration, sample volumes of 0.5, 1.0, 1.5 and 2.0 l of the NOAA SX-3570 standard were passed through a Nafion drier (Permapure™) before being trapped on Markes sorbent tubes as described above. Molar concentrations were derived for components of SX-3570 based upon aqueous phase calibrations, and converted to ppt in order to allow comparison with the mixing ratios reported by other institutions.

Seawater calibration standards and SX-3570 samples were analysed using an Agilent GC-MS (5973N), coupled to a Markes Unity thermal desorption (TD) platform. The GC is fitted with a 60 m DB-VRX capillary column (0.32 µm film thickness, J & W Ltd.), and the MS was operated in electron ionization (EI)/single ion mode (SIM) throughout the analyses. CH\(_3\)I was quantified using the CH\(_3\)I\(^+\) ion (m/z 142), CH\(_2\)Br\(_2\)
via $\text{CH}_2(^{79}\text{Br})(^{81}\text{Br})^+ (m/z \ 174)$ and $\text{CHBr}_3$ on $\text{CH}(^{79}\text{Br})(^{81}\text{Br})^+ (m/z \ 173)$. Within unity, the sample tube was heated to 200°C, and the sample refocused onto a cold trap held at −10°C. Following this, the cold trap was rapidly heated at 100°C s$^{-1}$ to 290°C, introducing the sample to the GC column with a helium carrier gas flow rate of 2 ml min$^{-1}$. The GC oven temperature was held at 40°C for 5 min, increased at 20°C min$^{-1}$ to 200°C and held for 2 min, before increasing at 20°C min$^{-1}$ to 240°C, and held for 4 min. GC-MS sensitivity drift was monitored and corrected for by injecting constant volumes of deuterated halocarbons (CD$_3$I, (CD$_3)_2$CDI) into each sample (prepared as for calibration standards from pure compounds) (Hughes et al., 2006, 2008). Total measurement uncertainties for each compound were estimated from the combined uncertainties (root sum of squares) of the following; volumes of liquid halocarbon standard additions at each stage during serial dilution, volumes of liquid standard additions to seawater samples, purge flow rate, and the uncertainty in the calibration linear regression. Precision was estimated as the standard deviation of repeat measurements performed during the calibration.

### 2.2.5 University of York

Analyses were carried out using an Agilent Technologies 6850 gas chromatograph coupled to a 5975C mass selective detector with an electron ionization source and operating in single ion mode (SIM). All samples were passed through a Nafion drier (Permapure™) prior to pre-concentration of volatile components onto a Peltier cooled (−20°C) adsorbent trap (UNITY2™ & CIA8, Markes International Ltd.). Analytes were thermally desorbed by heating the trap to 300°C, and separated using an HP5-MS column (25 m, 0.25 mm ID, 0.25 micron film, Agilent J&W). CH$_3$I was quantified using the CH$_3$I$^+$ ion ($m/z \ 142$) with I$^+$ ($m/z \ 127$) as a qualifier, whilst CH$_2$Br$_2$ was quantified via CH$_2(^{79}\text{Br})(^{81}\text{Br})^+ (m/z \ 174)$ with CH$_2(^{79}\text{Br})(^{79}\text{Br})^+ (m/z \ 172)$ as a qualifier ion, and finally CHBr$_3$ on CH($^{79}\text{Br})(^{81}\text{Br})^+ (m/z \ 173)$ with CH($^{79}\text{Br})(^{79}\text{Br})^+ (m/z \ 171)$ as a qualifier ion.
Sample volumes of 1.0, 1.5 and 2.0 l of the NOAA intercalibration standard were analysed to produce a response curve. Calibration of VSLH was achieved using a permeation oven based dynamic dilution technique based on Wevill and Carpenter (2004). Individual permeation tubes (Eco Scientific) containing pure VSLH liquids are contained in one of two temperature controlled ovens (thermostatic at 40 and 70 °C) and each tube permeates a single gaseous VSLH at a steady rate. Zero grade nitrogen flows through the system at 100 ml min \(^{-1}\) and dilutes the VSLH permeation gas to ∼ppm (parts per million) levels. Loop injections of 25 µl volumes of this gas are diverted to the instrument to achieve a multi-point calibration at ppt levels, and these calibrations were used to determine mixing ratios of VSLH in the intercalibration standard SX-3570. The permeation tubes are weighed periodically (using a KERN 770 mass balance, accurate to 0.01 mg) typically every 6 weeks over a period of ∼1 yr, to calculate the permeation rates. The estimated uncertainty of this calibration technique is ±12%, calculated from the combined uncertainties due to the variance of the permeation tube weighings, the calibration linear regression, and uncertainties in loop volumes and flow rates.

3 Results

Butler et al. (2010) recommend that in the first instance, inter-laboratory comparisons should focus upon the species CH\(_3\)I, CH\(_2\)Br\(_2\) and CHBr\(_3\), since these are the most widely reported of the short-lived halocarbons in both the atmosphere and ocean (e.g., Quack and Wallace, 2003; Chuck et al., 2005; Yokouchi et al., 2005; Butler et al., 2007; Quack et al., 2007; Carpenter et al., 2009). In addition, long term studies have characterised the stability of these species in compressed gas cylinders over prolonged periods (Butler et al., 2010), whilst the stability of other VSLH such as CH\(_2\)I\(_2\) and CH\(_2\)ICl has not been well studied. As such, although the iodine containing dihalomethanes CH\(_2\)I\(_2\) and CH\(_2\)ICl arguably play an equally important role in tropospheric halogen photochemistry, the lack of information regarding their long-term stability in pressurised gas
canisters means that at this stage, analyses of these gases have been excluded from the following discussion.

Each laboratory that participated in the inter-comparison used their existing measurement and calibration technique (detailed above) to determine the mixing ratios of CH$_3$I, CH$_2$Br$_2$ and CHBr$_3$ within the NOAA standard SX-3570. The results of these analyses are summarised in Table 1 and Fig. 1. The comparisons were carried out blind (i.e. the halocarbon mixing ratios in the NOAA SX-3570 gas standard were unknown), with the exception of the University of York calibration, where, as coordinators of the comparison, the scientists carrying out the analyses had prior knowledge of the certified mixing ratios.

It is important to stress that several different calibration scales exist for these gases, and there is no formal consensus as to which is the most accurate. As such, the data points with greater deviation from the NOAA-assigned values should not necessarily be considered “incorrect”. The purpose of this study is simply to identify systematic offsets between observations from the different institutions, in order to improve comparability between datasets.

The two institutions which already use the NOAA scale for VSLH calibration (University of Cambridge and University of East Anglia) reported mixing ratios which were most consistently in good agreement with the NOAA-assigned values across all compounds (ranging from $\sim$0.5–25% depending upon the species). Although both institutions report CHBr$_3$ mixing ratios $\sim$0.8 ppt larger than the NOAA-assigned value, and in excellent agreement with each other (which could indicate some degradation of CHBr$_3$ in the UoC and UEA standards, or an upward drift of CHBr$_3$ in SX-3570), all three values are consistent within the uncertainties of the measurements. The UoC, UEA and SX-3570 VSLH gas standards were all prepared and analysed by the NOAA laboratory at around the same time (August–September 2009).

Despite not currently using the NOAA scale for VSLH calibrations, both the Universities of York and Bristol report CHBr$_3$ mixing ratios in excellent agreement with the NOAA-assigned value, and CH$_3$I mixing ratios equivalent to the NOAA value, within
the measurement uncertainties. UoY calibrated \( \text{CH}_2\text{Br}_2 \) in SX-3570 slightly higher than NOAA (although in reasonable agreement taking into account the measurement uncertainties), whilst UoB calibrated \( \text{CH}_2\text{Br}_2 \) some \( \sim2-3 \) times lower than NOAA.

The group based at Plymouth Marine Laboratory primarily analyse VSLH concentrations in seawater, and do not routinely make atmospheric measurements of these gases. As such, their normal sampling methods had to be modified in order to carry out the gas-phase calibrations, and this should be taken into consideration when interpreting their reported values. The PML-assigned \( \text{CH}_2\text{Br}_2 \) mixing ratio was nevertheless in excellent agreement with the NOAA analyses, however, the PML values for \( \text{CH}_3\text{I} \) and \( \text{CHBr}_3 \) were in less good agreement.

The mean and median of all the independently derived \( \text{CH}_3\text{I}, \text{CH}_2\text{Br}_2 \) and \( \text{CHBr}_3 \) mixing ratios are in reasonably good agreement (within 15% for median) with the NOAA-assigned values. The average \( \text{CHBr}_3 \) mixing ratio is higher than the figure reported by NOAA, however this is skewed by one high value, and removing the PML mixing ratio gives a mean (and median) of \( \sim5.8 \) ppt \( \text{CHBr}_3 \), in very close agreement with the NOAA mixing ratio.

The largest deviation between the reported mixing ratios was observed in the analysis of \( \text{CH}_2\text{Br}_2 \) (41%), whilst the closest agreement was between \( \text{CHBr}_3 \) measurements (25% deviation). Excluding the PML values, the deviation between the \( \text{CHBr}_3 \) mixing ratios was reduced to \( \sim8\)%, and the spread in reported \( \text{CH}_3\text{I} \) values was also reduced, from 33% to \( \sim20\)%

### 4 Conclusions and future work

The results from the first UK national comparison for short-lived halocarbon calibrations provide a significant step forward to achieving the long-term goal of an internationally recognised calibration scale. This in turn will facilitate use of a coherent central database, containing a vast quantity of globally distributed VSLH measurements, and hence potentially lead to improved global source term parameterisations for these gases.
Since the ultimate long-term aim of these comparisons is to reduce the uncertainty in global emission estimates, we consider the spread in calibration scales for CHBr$_3$, CH$_2$Br$_2$ and CH$_3$I found in this study compared with the current uncertainty in their global budgets. Current global oceanic CHBr$_3$ flux estimates range from 240–1760 Gg Br yr$^{-1}$ (Quack and Wallace, 2003; Warwick et al., 2006; Butler et al., 2007), equivalent to $\sim$1000 Gg Br yr$^{-1}$±75%, an uncertainty considerably larger than the variation in the CHBr$_3$ calibration scales identified in this study. Likewise, the CH$_3$I global budget is estimated as $\sim$330 Gg I yr$^{-1}$±65% (ranging from 114–546 Gg I yr$^{-1}$, Moore and Grozsko, 1999; Bell et al., 2002; Butler et al., 2007; Jones et al., 2010), which equates to an uncertainty substantially greater than the spread in CH$_3$I calibrations. Despite exhibiting the most variability between calibration scales, the $\sim$66% uncertainty in the estimated global CH$_2$Br$_2$ source term ($\sim$169 Gg Br yr$^{-1}$±66%; Butler et al., 2007; Liang et al., 2010) still outweighs the discrepancies which might result from use of the different calibration scales reported here.

The results from this study suggest that gas phase CHBr$_3$ calibrations made by UK atmospheric research institutes are in good agreement, however this comparison also indicates that cross-calibrating between gaseous and dissolved VSLH is not straightforward. Compared to CHBr$_3$, there are larger discrepancies in CH$_3$I calibration scales, while CH$_2$Br$_2$ calibrations proved the least consistent of the VSLH analysed during this comparison. Differences in the analytical techniques employed by the institutions involved in this study (see Sect. 2.2) may also have contributed to differences in their reported mixing ratios. For example, different detection techniques (ECD vs. MS), separation methods (type of GC column and conditions) and sample collection and preparation methods are all factors which could potentially influence the calibrations to some extent.

On a national scale, the next desirable stage in this process would be for each institution to submit their CH$_3$I, CH$_2$Br$_2$ and CHBr$_3$ observations to the HalOcAt database, calibrated both according to their existing calibration scale, and adjusted for the common, NOAA, scale. As well as achieving comparability on a national level, this will
also enable measurements from the UK research groups that took part in this inter-
calibration to be interpreted alongside the large database of CH$_3$I, CH$_2$Br$_2$ and CHBr$_3$
observations already published using the NOAA scale (Butler et al., 2007). A summary
of relevant VSLH publications from the institutions that took part in this inter-calibration
is provided as an Appendix.

In order to achieve a reliable inter-comparison of those VSLH that are potentially less
stable in canisters and have even shorter lifetimes than CH$_3$I, CH$_2$Br$_2$ and CHBr$_3$ (such
as CH$_2$I$_2$ and CH$_2$ICl), the general consensus is that an in-situ field-based comparison,
with groups simultaneously measuring VSLH in ambient air, would be the best ap-
proach. As some of the poly-halogenated VSLH are rapidly photolysed by sunlight and
can also be less volatile and susceptible to wall losses, the sampling technique used
when quantifying these species is potentially as important as the calibration method
itself. Furthermore, this approach would also provide a mechanism through which
inter-laboratory comparisons could be extended to include measurements of VSLH
in seawater. Comparisons of dissolved VSLH may prove more challenging than gas
phase intercalibrations (since sparging efficiencies add an additional level of complex-
ity, and degradation within the water may make sequential sampling of one standard
over a period of a few months impractical), however these studies are necessary, since
seawater measurements form a critical component of the global VSLH budget.

In the long-term, this national inter-laboratory comparison of VSLH calibrations
should be repeated at frequent intervals (ideally on an annual basis), in order to main-
tain confidence in the level of comparability, and should ideally also be extended to
include other international institutions. In addition, explanations for the differences in
the calibration scales highlighted here should be explored, and the VSLH standard SX-
3570 should be returned to the NOAA-ESRL for reanalysis, in order to assess any drift
in mixing ratios which may have occurred since the initial analyses. Funding for this
type of activity is generally not readily available, yet establishing a robust, long-term
inter-laboratory comparison protocol for VSLH calibration is crucial to creating a coher-
ent, integrated global database for these gases, and will underpin any future studies
aiming to use such a database to achieve improved estimates for global VSLH source terms.

Supplementary material related to this article is available online at: http://www.atmos-meas-tech-discuss.net/4/765/2011/amtd-4-765-2011-supplement.pdf.

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**References**


Miller, B. R., Weiss, R. F., Salameh, P. K., Tanhua, T., Greally, B. R., Mühle, J., and Sim-


Table 1. Summary of CH\(_3\)I, CH\(_2\)Br\(_2\) and CHBr\(_3\) mixing ratios as certified by NOAA and determined by the individual UK research laboratories, together with the 1\(\sigma\) precision in the intercalibration analyses and the estimated 2\(\sigma\) uncertainty in the reported mixing ratio (including systematic uncertainties in the measurement, but excluding precision), \(U\). The number of data points used to derive the reported mixing ratio are given in brackets.

<table>
<thead>
<tr>
<th>Institution</th>
<th>Date of analysis</th>
<th>Calibration scale</th>
<th>CH(_3)I ((\pm 1\sigma)) (ppt)</th>
<th>(U) (ppt)</th>
<th>CH(_2)Br(_2) ((\pm 1\sigma)) (ppt)</th>
<th>(U) (ppt)</th>
<th>CHBr(_3) ((\pm 1\sigma)) (ppt)</th>
<th>(U) (ppt)</th>
</tr>
</thead>
<tbody>
<tr>
<td>NOAA</td>
<td>9 Sep 2009</td>
<td>NOAA</td>
<td>3.8 (\pm 0.10) (13)</td>
<td>0.76</td>
<td>2.6 (\pm 0.10) (10)</td>
<td>0.52</td>
<td>5.4 (\pm 0.10) (10)</td>
<td>0.54</td>
</tr>
<tr>
<td>University of Bristol (Mace Head)</td>
<td>23 Jun 2010</td>
<td>Linde (CH(_2)Br(_2), CHBr(_3)); NIES (CH(_3)I)</td>
<td>3.33 (\pm 0.07) (13)</td>
<td>0.52</td>
<td>0.70 (\pm 0.01) (10)</td>
<td>0.56</td>
<td>5.51 (\pm 0.07) (10)</td>
<td>0.62</td>
</tr>
<tr>
<td>University of Cambridge</td>
<td>12 Oct 2010</td>
<td>NOAA</td>
<td>4.76 (\pm 0.91) (8)</td>
<td>0.95</td>
<td>(\ast) 2.66 (\pm 0.21) (8)</td>
<td>0.53</td>
<td>6.26 (\pm 0.32) (8)</td>
<td>0.63</td>
</tr>
<tr>
<td>University of East Anglia</td>
<td>9 Jun 2010</td>
<td>NOAA</td>
<td>3.91 (\pm 0.02) (4)</td>
<td>0.94</td>
<td>2.62 (\pm 0.03) (4)</td>
<td>0.66</td>
<td>6.20 (\pm 0.44) (4)</td>
<td>1.27</td>
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<tr>
<td>Plymouth Marine Laboratory</td>
<td>9 Nov 2010</td>
<td>PML (liquid standards)</td>
<td>1.78 (\pm 0.10) (8)</td>
<td>0.21</td>
<td>2.93 (\pm 0.07) (8)</td>
<td>0.12</td>
<td>9.28 (\pm 0.57) (8)</td>
<td>0.93</td>
</tr>
<tr>
<td>University of York</td>
<td>25 Oct 2010</td>
<td>York (permeation oven)</td>
<td>2.97 (\pm 0.04) (3)</td>
<td>0.33</td>
<td>3.18 (\pm 0.05) (3)</td>
<td>0.41</td>
<td>5.30 (\pm 0.10) (3)</td>
<td>0.60</td>
</tr>
<tr>
<td>Mean</td>
<td></td>
<td></td>
<td>3.35</td>
<td>2.42</td>
<td>3.18</td>
<td>5.30</td>
<td></td>
<td></td>
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<tr>
<td>Median</td>
<td></td>
<td></td>
<td>3.33</td>
<td>2.66</td>
<td>4.14</td>
<td>6.20</td>
<td></td>
<td></td>
</tr>
<tr>
<td>% standard deviation</td>
<td></td>
<td></td>
<td>33%</td>
<td>41%</td>
<td>25%</td>
<td></td>
<td></td>
<td></td>
</tr>
</tbody>
</table>

\(\ast\) Note that the University of Cambridge GC-ECD was not able to distinguish between CH\(_2\)Br\(_2\) and CHBrCl\(_2\), which co-elute on their system. However, subtraction of the NOAA-assigned CHBrCl\(_2\) mixing ratio from the combined CH\(_2\)Br\(_2\) and CHBrCl\(_2\) signal (and assuming the same instrument response for CHBrCl\(_2\) and CH\(_2\)Br\(_2\)) gives the value indicated.
Fig. 1. CH$_3$I, CH$_2$Br$_2$ and CHBr$_3$ mixing ratios determined by each research group, with estimated measurement uncertainties. The solid line in each plot corresponds to the NOAA-assigned mixing ratio, and the grey dotted lines indicate the estimated uncertainty in the NOAA values.