

# Determining stages of cirrus ~~life-cycle~~ evolution: A cloud classification scheme

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**Abstract.** Cirrus clouds impose high uncertainties on climate prediction, as knowledge on important processes is still incomplete. For instance it remains unclear how cloud microphysical and radiative properties change as the cirrus evolves. Recent studies classify cirrus clouds into categories including “in situ”, “orographic”, “convective” and “liquid origin” clouds and investigate their specific impact. Following this line, we present a novel scheme for the classification of cirrus clouds that addresses the need to determine specific stages of cirrus ~~life-cycle~~ evolution. Our classification scheme is based on airborne Differential Absorption and High Spectral Resolution Lidar measurements of atmospheric water vapor, aerosol depolarization, and backscatter, together with model temperature fields and simplified parameterizations of freezing onset conditions. It identifies regions of supersaturation with respect to ice (ISSR), heterogeneous and homogeneous nucleation, depositional growth, and ice sublimation and sedimentation with high spatial resolution. Thus ~~the whole cirrus life-cycle can be traced~~all relevant stages of cirrus evolution can be classified and characterized. In a case study of a gravity lee wave influenced cirrus cloud, encountered during the ML-CIRRUS flight campaign, the applicability of our classification is demonstrated. Revealing the structure of cirrus clouds, this valuable tool might help to examine the influence of ~~life-cycle~~evolution stages on the cloud’s net radiative effect and to investigate the specific variability of optical and microphysical cloud properties in upcoming research.

## 1 Introduction

Cirrus play an important role for weather and climate: besides their influence on the water vapor budget in the upper troposphere through condensation and evaporation (Dinh et al., 2014) and dynamics due to latent heat (Spichtinger, 2014), they modify the radiation balance of the Earth and atmosphere. Thin, opaque cirrus clouds transmit most of the incident solar radiation and absorb long-wave radiation from the Earth’s surface. As they are typically high and cold, they only emit little long-wave radiation into space, and thus cause a trapping of radiative energy in the Earth-atmosphere system, which eventually contribute to a rising surface temperature. If the cloud is thick, reflection of solar radiation back to space can get greater than the long-wave absorption, and consequently can cause the surface of the Earth to cool (Baran, 2009). This net radiative effect depends on macroscopic cloud properties like optical thickness, ice water content, and geometric extent as well as on its microphysical parameters such as ice crystal number, size, and shape (Schnaiter et al., 2016; Gallagher et al., 2012; Zhang et al., 1999).

Today many factors are known that determine Recent studies have investigated factors that affect these properties: the  
The amount and composition of natural and anthropogenic aerosol particles in the troposphere and their ability to nucleate  
ice crystals (DeMott et al., 2010), the exact freezing condition and mechanism (Cziczo et al., 2013), are determining details  
of heterogeneous freezing (DeMott et al., 2010), but aerosol distribution and properties are not well known until today, in  
5 particular in the temperature range of cirrus clouds. Exact freezing conditions and mechanisms were studied by Cziczo et al. (2013) from  
in-situ measurements taken during flight campaigns conducted in Middle and North America. They found a dominance of  
heterogeneous freezing with mineral dust and metallic particles as the main source of residual particles. Although they managed  
to span a range of geographic regions and seasons, it is not clear, if those results are globally valid.

High water vapor supersaturations inside and outside of cirrus clouds point to several microphysical processes and often  
10 occur together with low ice crystal numbers (Krämer et al., 2009). Sublimation of ice crystals was reported to result in the  
disappearance of facets and corners, changing the crystal symmetry (Korolev and Isaac, 2006). The dependence of freezing on  
the updraft velocity during cloud formation (Kärcher and Lohmann, 2002), water vapor supersaturation (Krämer et al., 2009),  
cloud age (Korolev and Isaac, 2006), etc is theoretically known (Kärcher and Lohmann, 2002), but measurements are difficult  
and rare. In idealized settings, there is also a good understanding of the relevant processes in cloud formation and break up.  
15 But in nature no two clouds are alike and there exists a confusing variability of conditions under which they occur. This makes  
it difficult to represent cirrus clouds adequately in global circulation models for weather and climate prediction (Baran, 2012;  
Zhang et al., 2005).

In order to gain more insight into the particular role of different cirrus clouds, great efforts were made to classify cirrus by  
the meteorological contexts in which they occur (Jackson et al., 2015; Muhlbauer et al., 2014). Categories include “synoptic”,  
20 “orographic”, “lee wave” and “anvil” cirrus. Recently a more general classification was introduced that distinguishes the  
groups of “liquid origin” and “in situ” clouds that describe whether the cirrus formed from a pre-existing liquid cloud or from  
cloud-free air by freezing of existing water droplets at water saturation, or by freezing of solution droplets or heterogeneous  
nucleation at ice supersaturation but below water saturation (Krämer et al., 2016; Wernli et al., 2016). Such a classification  
classifications of recorded data is-are a prerequisite for statistically investigating the specific properties and influences of  
25 different clouds, and to extract the governing mechanisms and parameters from remote sensing and in situ measurements.

Likewise, detailed knowledge of cloud properties at different stages of evolution is yet to be gained, as a cloud is expected to  
show different properties at the time of formation and break-up/dissipation. During the evolution of a cloud ice particle number,  
effective radius and particle shape evolve as well. For instance, Korolev and Isaac (2006) showed that during dissipation, ice  
crystals loose-lose single facets and corners, thus changing their geometry significantly. Besides such changes of microphysical  
30 properties, also macrophysical qualities like the exact location, altitude and extent of cloud parts in a specific evolution state  
may result in a different net radiative effect.

A statistical study (Comstock et al., 2004), evaluating ground-based Lidar data taken in France over a time period of one  
year, and in-situ measurements (Korolev and Isaac, 2006; Heymsfield and Miloshevich, 1995; Heymsfield, 1975) indicated  
A classification scheme that reveals the spatial distribution of evolution stages would facilitate the investigation of possible  
35 dependencies on cirrus evolution.

Heymsfield (1975) first illustrated and documented the vertical and dynamical structure of ice generating cirrus uncinus clouds. This early work and following in-situ measurements (Heymsfield and Miloshevich, 1995) indicate that there is a vertical order of cirrus ~~life-cycle evolution~~ stages with ice nucleation near cloud top level, deposition of water vapor onto ice crystals and thus particle growth in the middle, and sublimation and sedimentation at cloud base level. ~~Due to its statistical origin, this proposed structure must be seen as an averaged reference order, where individual clouds~~ A more recent, statistical study by Comstock et al. (2004) evaluated an extensive data set of ground-based Lidar measurements taken at the ARM Southern Great Plains site (Oklahoma, USA) over a time period of one year. Vertical profiles of determined relative humidity with respect to ice ( $RH_i$ ) inside of cirrus clouds were divided into the upper most 25 %, the middle 50 % and the lower 25 % of total cloud depth. The frequency distribution of  $RH_i$  of the upper 25 % show a considerable amount of supersaturated regions with high  $RH_i$  values up to 160 %, associated with ice nucleation. The distribution of the lower 25 % is shifted towards subsaturation with a maximum between 70 % and 80 % and values down to 10 %, clearly dominated by crystal sedimentation and sublimation. Therefore they showed that the generally accepted vertical order of evolution stages dominated the majority of measured clouds while individual clouds, depending on cloud type, generation mechanism, cloud age, and atmospheric dynamics, may show strongly differing distributions (Groß et al., 2014). ~~An evolution classification of individual cloud parts~~ (Comstock et al., 2004; Groß et al., 2014). The classification scheme that we present is based on atmospheric Lidar cross-sections and therefore facilitates the detailed investigation of ~~all the above mentioned influences~~ evolution stages, their vertical and horizontal order, the impact of atmospheric dynamics, and their specific optical properties.

Such a classification needs information on relative humidity with respect to ice ( $RH_i$ ) and temperature, as they are two governing variables in ice particle formation, growth, and disappearance (Pruppacher and Klett, 2010). In order that ice can form in the atmosphere,  $RH_i$  must reach or surpass 100 %. It is well known that tropospheric air masses often show substantial supersaturations with respect to ice (Spichtinger et al., 2003). These so called ice-supersaturated regions (ISSR) result mainly from upward motion of air masses originating in diverse atmospheric dynamics like large scale synoptic ~~ascends~~ ascent (e.g. warm conveyor belt), convective systems or meso-scale gravity waves (Spichtinger et al., 2005a, b; Field et al., 2001).

The existence of an ISSR does not automatically imply the existence or formation of a cirrus cloud. For the homogeneous freezing (HOM) of solution droplets at cirrus temperatures (typically below 235 K), high supersaturations in the order of 140 % are necessary (Koop et al., 2000). Additionally, solid aerosol particles can act as ice nuclei and thus lead to freezing under a much broader range of conditions. Heterogeneous freezing onset temperatures and saturation ratios depend strongly on aerosol type, coating of the particles, and their size and are still subject to current research (Hoose and Möhler, 2012).

Once ice particles are present ~~, and no cooling, e.g. from updraft motion takes place,~~ remaining supersaturation is quickly depleted by deposition of water vapor onto existing crystals. ~~Depending on the particle number and average radius~~ In updraft regions of ice clouds however, it may take a few minutes to a few hours ~~for the equilibrium of 100 % to be reached until the air reaches the quasi-steady supersaturation which is a function of vertical velocity and ice particle size distribution and higher than 100 % for positive vertical velocities~~ (Korolev and Mazin, 2003). Furthermore, effects that hinder phase relaxation are discussed, stemming from e.g. liquid coating around ice particles (Bogdan and Molina, 2009), or dynamic processes (Spichtinger and Krämer, 2013). Thus, supersaturation inside of ice clouds can persist for a substantial amount of time. Likewise, regions

of subsaturation with respect to ice can emerge when heavy ice particles sediment out of the ISSR, or  $RH_i$  is reduced by warming. These regions are dominated by sublimation of ice crystals (Korolev and Isaac, 2006; Kübbeler et al., 2011).

It can be seen that detailed knowledge of humidity and temperature in and around the cloud, as well as knowledge about freezing onset conditions is necessary in order to identify different stages of cloud evolution. The airborne Differential Absorption Lidar WALES ("Water vapour Lidar Experiment in Space"), flown aboard the German research aircraft HALO ("High Altitude and Long range", model: Gulfstream G550), makes major parts of the needed data available. It provides an unique dataset of collocated, high spatial resolution measurements of atmospheric backscatter, depolarization and water vapor, enabling us to distinguish in-cloud and cloud-free regions, to identify the relevant aerosol type in the vicinity of the cloud and to calculate relative humidity.

In this paper, we present a detailed classification scheme for the evolution of cirrus clouds. Besides WALES measurements we use complementary model temperature fields from the European Centre for Medium-range Weather Forecasts (ECMWF). Provided with high resolution, two-dimensional Lidar cross-sections of the atmosphere, we are able to study the structure of clouds and the spatial distribution of classified evolution stages. By setting in-situ and remote sensing data in perspective to cirrus evolution, it facilitates the study of the specific optical, microphysical and radiative properties of evolution stages. We apply our scheme in a first case study of a lee wave influenced cirrus cloud over France, encountered during the ML-CIRRUS 2014 campaign (Voigt et al., 2016), demonstrate its applicability, and investigate the impact of both ~~small-scale~~ mesoscale and large-scale dynamics on the cloud structure. We end with a summary of the classification scheme and a brief outline of its potential in cirrus cloud research.

## 2 Water vapor remote sensing during ML-CIRRUS

In spring 2014, the Mid-Latitude Cirrus experiment ML-CIRRUS was conducted. It was designed to investigate natural cirrus and anthropogenic contrail cirrus with regard to their nucleation, life-cycle, and climate impact. In this campaign, the German research aircraft HALO, equipped with a combined in situ and remote sensing payload, performed 16 measurement flights above Europe. The on-board cloud probes, WALES Lidar and novel ice residual, aerosol, trace gas and radiation instruments probed mid-latitude cirrus clouds originating from e.g. air traffic, warm conveyor belts, jet streams, or mountain waves (Voigt et al., 2016).

WALES is an airborne Differential Absorption Lidar that measures the tropospheric water vapor concentration below the research aircraft by simultaneously emitting laser pulses at three online and one offline wavelength in the water vapor absorption band around 935 nm (Wirth et al., 2009). The averaged pulse energy is 35 mJ with a repetition rate of 200 Hz. The three online wavelengths provide the needed sensitivity to compose a complete water vapor profile from the partly overlapping line contributions that ranges from just below the aircraft down to ground level. Additionally, WALES is equipped with one channel at 1064 nm and one High Spectral Resolution channel at 532 nm using an iodine filter. Both receiver channels are designed to detect the depolarization of the backscattered light (Esselborn et al., 2008).

WALEs is capable to provide collocated measurements of humidity in the form of water vapor volume mixing ratio  $r_w$ , backscatter ratio (BSR), and aerosol depolarization ratio (ADEP). Those measurements form a two dimensional curtain along the flight track of the research aircraft intersecting the atmosphere below. The Lidar data we use in this paper has a vertical resolution of 15 m. Raw data is sampled at a rate of 5 Hz. At HALO's typical ground speed of  $210 \text{ ms}^{-1}$  and after averaging  
5 for a better signal-to-noise ratio, horizontal resolution is 2.5 km for humidity and 210 m for BSR and ADEP.

We use ECMWF analysis ~~and forecast temperature data~~ temperature data (available every 6 hours), with a horizontal resolution of  $0.25^\circ$  and 91 vertical levels that we interpolate ~~both temporally and spatially~~ linearly in time and bi-linearly in space onto the Lidar measurement cross-section. Then we calculate relative humidity with respect to ice from this temperature information and the measured absolute humidity:

$$10 \quad RH_i = \frac{r_w \cdot n_{air} \cdot T \cdot k_B}{e_{sat,i}(T)}, \quad (1)$$

with temperature  $T$ , volume number density of air  $n_{air}$ , and Boltzmann constant  $k_B$ . We use the parameterization for water vapor saturation pressure over ice  $e_{sat,i}$  by Murphy and Koop (2005).

The accuracy of calculated relative humidity relies strongly on the quality of absolute humidity and temperature data. WALEs humidity measurements exhibit a mean statistical uncertainty of 5 %. The applicability of ECMWF temperature in  
15 this calculation was investigated by Groß et al. (2014). They showed that during ascent and descent of a similar research flight in 2010 the mean temperature difference between ECMWF and on-board temperature sensors was ~~always less than  $10.8$  K~~ and estimated a resulting maximum relative uncertainty of  ~~$10$ – $15$  % in the calculated  $RH_i$~~   $10$  to  $15$  % as an upper boundary for the calculated  $RH_i$  at typical cirrus temperatures. In cases where collocated radiosonde or dropsonde measurements are available, their temperature profiles can be used to calibrate ECMWF temperature fields, eliminating possible offsets. As modern sondes  
20 feature measurement uncertainties of down to  $0.2$  K, the total relative uncertainty of  $RH_i$  therefore potentially can be reduced to values as low as  $6$  %.

### 3 Cirrus evolution classification scheme

With atmospheric ~~lidar~~ Lidar cross-sections at hand, we are able to identify in-cloud and cloud-free regions by applying a threshold for the backscatter ratio (see Fig. 1). As there is no sharp boundary between a cloud and its surrounding, this  
25 threshold value holds a certain arbitrariness. In the case study, we use a value of 2, but in cases where e. g. thick aerosol layers are present this threshold might need to be increased to avoid classifying parts of the aerosol layer as in-cloud regions.

Looking at cloud-free parts of the cross-section, regions that might possibly lead to cirrus cloud formation can be identified by searching for data points exhibiting ice supersaturation ( $RH_i > 100$  %). Moderately supersaturated cloud-free parts are classified as ISSR<sub>out</sub>. With higher supersaturations, the chances for the imminent nucleation of ice particles get increasingly  
30 higher. Therefore we introduce the classes HET<sub>out</sub> and HOM<sub>out</sub> in our classification. They represent regions outside of the cloud where onset condition for heterogeneous and homogeneous freezing get surpassed, respectively. Their classification is implemented via temperature dependent humidity thresholds.

It should be noted that ice is forming latest, as soon as conditions for homogeneous freezing get reached, as there is always a sufficient amount of solution droplets in the atmosphere (Minikin et al., 2003). Therefore, a cloud classification should not feature considerable regions of  $\text{HOM}_{\text{out}}$ . This fact should be kept in mind when choosing a BSR threshold value for the cloud border detection, making sure that HOM regions lie inside the cloud. ~~HET<sub>out</sub> regions, however, may exist in cases with no~~  
 5 ~~sufficient amount of aerosol ice nuclei. For~~ However, this might not always be completely achievable without misclassifying aerosol layers (see above).

Koop et al. (2000) found that the homogeneous freezing temperatures of numerous aqueous solution droplets (1-10  $\mu\text{m}$ ) would fall on one single solute-independent curve, when plotted in terms of water activity  $a_w$ . They suggested that this curve could be constructed by shifting the melting-point curve by  $\Delta a_w$ . From experiments on homogeneous freezing, ~~we they~~  
 10 ~~determined a shift by  $\Delta a_w = 0.305$ . For atmospheric applications, water activity is equal to relative humidity, when the droplet is in equilibrium with the water vapor pressure of the surrounding air (Kärcher and Lohmann, 2002). We use these findings to~~ extract a parameterization of the temperature dependent onset humidity for HOM ~~from Koop et al. (2000, their Fig. 3) for a~~ nucleation rate  $\omega = 0.0167 \text{ s}^{-1}$  and a droplet size of 0.5  $\mu\text{m}$  (see Table 1,  $RH_{i,\text{HOM}}(T)$ ).

To determine a humidity threshold for HET, detailed information of the involved aerosol type, its coating, and size distribution would be required. Then results from laboratory experiments on onset freezing temperatures and saturations for this kind  
 15 of aerosol could be used. As heterogeneous freezing conditions are still subject to current research (Hoose and Möhler, 2012) and as comprehensive aerosol information is difficult to acquire solely from remote sensing, we make only a coarse distinction between two important aerosol types: mineral dust (MD) and coated soot (CS).

Together with a synergistic analysis, WALEs ~~measurements of aerosol linear depolarization ratio (ADEP) and Lidar ratio~~  
 20 ~~Lidar data~~ can be used to identify the relevant aerosol type in the measurement area. To this end, we ~~use apply~~ an aerosol classification suggested by Groß et al. (2013). ~~It uses the fact that aerosol types can be distinguished by three intensive optical properties, aerosol Lidar ratio, aerosol linear depolarization ratio, and color ratio. Mineral dust shows linear depolarization values of more than 20 %, coated soot less than 20 %.~~ Then we employ simplified onset parameterizations  $RH_{i,\text{HET}}^{\text{MD}}(T)$  and  $RH_{i,\text{HET}}^{\text{CS}}(T)$  (see Table 1 and Krämer et al. (2016, their Fig. 4)). Until more detailed parameterizations are available, this  
 25 imposes ~~a an~~ uncertainty for the determination of the exact border of heterogeneous freezing regions. ~~In contrast to HOM<sub>out</sub>, HET<sub>out</sub> regions may exist in cases with no sufficient amount of aerosol ice nuclei, or due to simplifications in the utilized  $RH_{i,\text{HET}}(T)$  threshold.~~ The classes  $\text{ISSR}_{\text{out}}$  and  $\text{HET}_{\text{out}}$  represent pre-stages of ~~Cirrus cirrus~~ formation and indicate regions where a cirrus cloud is likely to develop.

Inside of a cloud ( $\text{BSR} > 2$ ), we proceed in the same manner. When the  $RH_{i,\text{HOM}}(T)$  threshold is surpassed we classify  
 30 as  $\text{HOM}_{\text{in}}$ . This region shows active ice nucleation. Together with  $\text{HET}_{\text{in}}$ , that we classify analogously, it ~~represent represents~~ the youngest evolution stage of a cirrus cloud.  $\text{HET}_{\text{in}}$  ~~will also is also expected to~~ show active nucleation as long as ice nuclei are present in the freezing region. ~~However, due to the limitations mentioned above, the border of HET towards lesser supersaturated areas must be interpreted with caution. Also ice crystals found in HET must not necessarily be formed by heterogeneous freezing, as sedimentation from higher levels featuring different nucleation conditions may take place. Still,~~

heterogeneous freezing is an important freezing mechanism in mid-latitudes (Cziczo et al., 2013) and the class HET<sub>in</sub> adds more information to the classification leading to a more complete characterization of cirrus clouds.

When relative humidity inside the cloud is lower than the freezing thresholds, we classify as DEP, as the remaining supersaturation is depleted by deposition of water vapor onto the existing ice particles. This intermediate evolution stage is dominated by depositional growth of ice crystals. The final evolution stage of a cloud sets in, when relative humidity falls below 100 %. In such an environment ice inevitably must sublime. We classify this region as SUB.

~~This classification scheme is applied independently to every recorded data point.~~ During the life-cycle of a cloud, nucleation, growth and sublimation events may occur more than once, e.g. when atmospheric dynamics cause renewed updrafts and a second freezing event on top of pre-existing ice takes place. As described, our method is able to identify nucleation, growth, sublimation regions and pre-stages of cloud formation. However on its own, it does not yield any information about earlier developments of those regions. Its very strength is to reveal the actual atmospheric state with regards to cirrus evolution at the time of measurement. This is done on a high spatial resolution that exceeds typical resolutions of GCMs, enabling the detailed study of individual cloud parts.

#### 4 Case study ML-CIRRUS 2014-03-29

We demonstrate the applicability of our classification scheme in a cirrus case that was obtained during the ML-CIRRUS field campaign on 29 March 2014. The meteorological situation over Western Europe and the Iberian Peninsula on the flight day is dominated by a trough extending from west of Ireland to the Iberian Peninsula and further to the western part of North Africa (Fig. 2 a). At 300 hPa, high southerly winds with wind speeds up to  $35 \text{ ms}^{-1}$  are observed on the leading edge over Southern France and Spain. ~~Model~~ Two days before the research flight, model forecasts indicated the existence of cirrus forming from high updrafts over the Pyrenees, as well as cirrus influenced by lee waves north of the mountain ridge. Additionally ~~high dust concentrations of Saharan mineral dust~~ highly dust loaded air masses over Southern France, originating from Saharan dust events over Algeria were expected.

In this meteorological setting the research flight was performed with the aim to sample all stages of cirrus evolution that resulted from an overflow of the Pyrenees with high wind speeds and consequent gravity wave excitations in the lee of the mountain ridge. Therefore the flight path in the relevant measurement region was chosen to run along the main wind direction, sampling the clouds along their path of advection.

The flight (Fig. 2, red flight path) started in Oberpfaffenhofen, Germany at 12:37 UTC and first went westward towards Paris, followed by a southward flight leg towards Spain at an altitude of 11200 m. The investigated cirrus cloud was encountered over Southern France during this leg that is running with a bearing of  $190^\circ$  (white flight leg). Inside cirrus clouds, over the Pyrenees mountains, three legs at different lower altitudes followed. From the Mediterranean coast the aircraft turned eastward and probed cirrus at several altitudes near the Balearic islands before it went northward towards Oberpfaffenhofen (landing at 19:50 UTC).



## 4.1 Cirrus leg overview

The following discussion of the classification scheme focuses on the southward flight leg stretching about 400 km to the north and 200 km to the south of the Pyrenees (Fig. 2, white flight path). Fig. 2 b shows a false color image of the Pyrenees area derived from SEVIRI ("Spinning Enhanced Visible and InfraRed Imager") data at 14:30 UTC. ~~Higher, cooler clouds have~~ The red, green and blue color channels of the image take SEVIRI's 635 nm, 850 nm and inverted 10.8  $\mu$ m channel data. This way, the high and therefore cool cirrus clouds stand out by a bluish color; low clouds are. Low clouds get depicted in yellow and the surface of the Earth has a green tone. Coming from the North, the flight path intersects an ice cloud that is part of a larger cloud regime expanding from Southern France towards the Algerian coast. This cloud is followed by a completely cloud-free area north of the Pyrenees. Over the mountain ridge some localized high clouds are crossed.

Earlier satellite pictures and ECMWF data (not shown) indicate that the cirrus cloud regime formed under the influence of a highly humid air mass stemming from Southern Spain that also shows the tendency to produce lower liquid clouds. As the air mass is transported to the North, it is impacted by gravity waves generated when crossing the Pyrenees. The analysis of preliminary temperature data (not shown) provided by the on-board, passive "Microwave Temperature Profiler" (Denning et al., 1989) shows that temperature oscillations, caused by gravity waves, continue to exist inside the cirrus cloud intersected by the flight path.

In Fig. 3 we plot a cross-section showing backscatter ratio at a wavelength of 532 nm along the chosen part of the flight path. Here atmospheric features apparent in Fig. 2 can be studied in greater detail. On the lee side, north of the Pyrenees (14:19-14:34 UTC), a high cirrus cloud is observed that extends from a height of 7 km to ~~11~~ 10.4 km. The southern and middle parts are dominated by high BSR values from 50 up to 200, whereas the northern section exhibits lower values. Aerosol linear depolarization ratios of more than 30 % inside the cloud (not shown) and temperatures below 240 K clearly indicate a pure ice cloud. Over the Pyrenees (14:42-14:53 UTC) a lower cirrus cloud is located at an altitude of about 6 km. Its spatially restricted occurrence over the mountain ridge indicates a formation due to forced updrafts stemming from the southerly cross-mountain flow. Even lower, at a height of 4 km a thick aerosol layer is discernible. An analysis of ADEP ~~indicates that this layer contains~~ shows values between 20 % and 30 %, typical for Saharan mineral dust (Groß et al., 2013) which is consistent with the origin of the air masses in North Africa, as forecasted by dust models.

Furthermore, in the region between the two clouds (14:34-14:43 UTC) gravity lee waves are discernible at an altitude of about 9500 m and also in the lower aerosol layer. These waves are expected to influence at least parts of the northern cirrus cloud. We will investigate them in more detail in Sect. 4.3.

## 4.2 Classifying evolution stages

In the following we will apply our classification scheme to the high cirrus cloud north of the Pyrenees. ~~Fig. Figures~~ Fig. 4 and Fig. 5 give a close-up view of the selected data marked with a black rectangle in Fig. 3. Water vapor volume mixing ratio  $r_w$  measured by WALES is plotted in Fig. 4 a, together with red isolines of ECMWF temperature. A black contour line (BSR = 2) marks the cloud border. Being an absolute humidity measure,  $r_w$  generally decreases with increasing altitude, as temperature is



decreasing and almost all sources of water vapor are located at the Earth's surface. Contrastingly, a humid layer, surrounded by dryer air at a height of approximately 9000 m, can be found upstream of the cirrus cloud (14:34 – 14:40 UTC). In this region, the water vapor data exhibits the same oscillations as previously seen in the BSR data.

Relative humidity with respect to ice (Fig. 4 b) is calculated from this data using the ECMWF model temperature field.

5 As expected, supersaturated regions (blue) are found mostly inside of the cirrus. There are also major subsaturated regions (red) in the northern part of the cloud. South of the cirrus high supersaturations exist in cloud-free air, mostly in the crests of the gravity waves (14:34-14:36 UTC). The highest supersaturations are measured in the most southern part of the cloud (14:33-14:34 UTC). They indicate a nucleation region.

To investigate individual parts of the cloud in more detail, we apply our classification and visualize the result in Fig. 5. Data  
10 pixels are classified (Sect. 3.2) and marked in color accordingly, and in-cloud and cloud-free regions can be distinguished by the black contour line for a BSR value of 2. Heterogeneous freezing is identified using the  $RH_{i,HET}^{MD}(T)$  threshold. HET threshold for coated soot ( $RH_{i,HET}^{CS}(T)$ ). We base this decision on 48-hour backward trajectory calculations of air masses at 8.5 km to 9.5 km (not shown) that indicate no influence from lower Saharan dust loaded levels. Subsaturated regions outside of the cloud are left blank and areas where no valid data is available are indicated by black hatching.

15 The above mentioned humid layer, discernible in Fig. 4, reaches ice supersaturation (ISSR) in the two crests of the gravity lee wave to the south of the cloud (14:34-14:36 UTC). ~~Here, values of  $RH_i$  are even higher than the threshold for HET freezing.~~ At the cloud edge, also the ~~HOM freezing threshold is HET~~ and HOM freezing thresholds get surpassed (14:33-14:34 UTC). ~~The~~ As no relevant isolated HET region exists and HOM freezing also sets in at the cloud edge, we assume that homogeneous freezing might be dominant freezing mechanism here. This upper, southern section of the cirrus ~~(14:32-14:34 UTC)~~ is dominated by ice nucleation and represents the youngest part of the cloud.

~~In the~~ The top level of the cloud and the ISSR south of it climbs from about 9,3 km to 10,4 km at 14:31:50 UTC. Considering the speed of the aircraft ( $200 \text{ ms}^{-1}$ ) and the wind speed (about  $30 \text{ ms}^{-1}$  in opposite direction), and assuming that the found pattern is stationary (more in Sect. 4.3), this corresponds to a vertical velocity of about  $50 \text{ cms}^{-1}$ . In the temperature range between 200 K and 220 K, such high vertical velocities typically lead to high ice crystal number densities with small ice  
25 particles ( $r_0 < 10 \text{ }\mu\text{m}$ ; Kärcher and Lohmann (2002)). As the updraft region is observed between 14:32 UTC and 14:36 UTC, we assume that freezing takes place at least between the cloud edge and the peak of the cloud top at 14:31:50 UTC, although no valid data is available for parts of the upper cloud, due to detector saturation.

In the middle (14:26-14:32 UTC), a section of moderate supersaturation (DEP) is located. This is an already well developed part of the cirrus that is dominated by depositional growth of ice crystals. After ~~an the~~ initial ascent (14:32-14:34 UTC), the  
30 cloud top ~~level~~ slopes from over 10 km down to under 9 km at the northern edge. ~~This indicates a corresponding to a vertical velocity of about  $-30 \text{ cms}^{-1}$ . Sedimentation may contribute to this, but sub  $10 \text{ }\mu\text{m}$  particles are rather associated with fall speeds of under  $2 \text{ cms}^{-1}$  (Gasparini et al., 2016, their Fig. 1). This~~ large-scale descent ~~reducing supersaturation and evoking,~~ also apparent in ECMWF model data (not shown), reduces supersaturation and evokes the intermediate DEP region as well as large connected regions of subsaturation (SUB) in the northern part of the cloud (14:19-14:26 UTC). Here the cloud is starting  
35 to break up, as ice particles are sublimating.

From these results all cirrus ~~life-cycle evolution~~ stages can be identified: from ~~cloud-free ISSR and~~ ice nucleation (HET, HOM) aided by vertical displacements in a gravity lee wave, to crystal growth by deposition of water vapor in a moderately supersaturated region DEP, to the dissolving of the cloud in a subsaturated region (SUB), ~~probably~~ caused by a large-scale descent.

- 5 The ~~detailed spatial~~ distribution of these major stages of cirrus evolution ~~features a horizontal order instead of a general vertical structure found in cirrus clouds over France (Comstock et al., 2004). Surprisingly we even feature a horizontal component with ice nucleation in the south and dissipation in the north. Also, we~~ find SUB regions at the cloud top level, ~~in part,~~ located above DEP regions in the ~~northern part north~~ of the cloud (Fig. 5). ~~Similarly, Our findings fit well into the perspective of~~ model simulations, investigating the influence of dynamics on the evolution of a cirrus cloud (Spichtinger and Gierens, 2009) ~~, that~~
- 10 also found more complex horizontal distributions deviating from a simplistic cirrus evolution pattern comprising ice nucleation at cloud top level, a crystal growth in the middle ~~an and~~ sublimation at the bottom. Thus our classification illustrates how the large-scale meteorological context, wind and gravity wave fields can effect the structure of individual clouds.

- ~~Unfortunately, no dropsonde temperature measurements were conducted in this cirrus case. Thus we must assume a relative uncertainty of  $RH_i$  between 10 % and 15 % at cirrus level (see Sect. 2). This directly translates into an uncertainty in locating~~
- 15 ~~the exact border between evolution stages. However, we found that the horizontal distribution order and the existence of SUB regions at the cloud top continue to exist even after offsetting the temperature field by  $\pm 0.8$  K. Therefore we are confident that ECMWF temperature data is suitable for this kind of analysis.~~

### 4.3 Investigating the influence of lee waves

- A special feature of this case study is the presence of lee wave patterns in the cloud region. Fig. 6 gives a close-up view of
- 20 BSR data in the cloud-free area south to the cirrus cloud. A layer of slightly higher BSR ( $> 1.2$ ) is located above an altitude of 9500 m. It shows clear oscillations at its boundary to a lower, cleaner layer of air. One period extends over about 66 seconds in measurement time which corresponds to an apparent wavelength of 14 km with vertical displacements of up to 190 m.

- ECMWF model data, available for this cirrus case, features a horizontal grid spacing of about 16 km and thus ECMWF is not able to resolve the ~~small-scale mesoscale~~ lee waves in its temperature gradients ~~, (see Fig. 4 a).~~ As a result, temperatures in
- 25 the crests might be even lower and  $RH_i$  values therefore underestimated.

- To investigate this possible deviation and its influence on our classification results, we simulate adiabatic cooling of an air parcel along a hypothesized trajectory (Fig. 5 ~~, blue line and Fig. 6~~) in front of the cirrus cloud. The trajectory runs 200 m under and parallel to the contour line of  $BSR = 1.2$  (~~not shown~~ Fig. 6), that separates the two distinct layers of air. The wind direction in this region differs by less than  $10^\circ$  from the flight path. This makes us confident that the simulated trajectory resembles a
- 30 real trajectory reasonably well, under the assumption of a stationary air flow.

SEVIRI images from 14:00 UTC and 15:00 UTC (not shown) indicate, that the northern and southern edges of the cloud are moving only about 20 km to the north along the flight path. That corresponds to a cloud velocity of under  $6 \text{ ms}^{-1}$ , compared to wind speeds of up to  $35 \text{ ms}^{-1}$  at cirrus altitude. ~~As in a lee wave cloud~~ Similar to real lee wave clouds (Field et al., 2012), air is flowing through the region, becoming part of the cloud in the south and leaving the cloud in the north. In the confined

area and time frame of our simulation, we consider the underlying wind and wave fields to be quasi-stationary. However this might certainly not be true for the duration of the whole flight leg (14:18 - 14:41 UTC). Also our simulation is not intended to provide corrected temperature data but to illustrate the general influence of gravity waves on cloud formation and of unresolved temperature fluctuations on our classification.

- 5 Along this trajectory we calculated peak vertical velocities of up to  $120 \text{ cm s}^{-1}$ , more than twice as high as the average vertical velocity of  $50 \text{ cm s}^{-1}$  in the initial updraft region, estimated in Sect. 4.2. During nucleation, such peaks may lead to even smaller ice crystals and higher ice crystal number densities of up to  $50 \text{ cm}^{-1}$ .

In Fig. 7, ECMWF temperature and relative humidity calculated with ECMWF temperature along the trajectory are plotted (blue) as a function of measurement time. As the trajectory follows the vertical displacements of the gravity wave along the  
10 wind direction, i. e. from right to left in Fig. 3 to 7, relative humidity and ECMWF temperature show oscillations. They stem from the undisturbed temperature gradient, as ECMWF does not resolve the ~~small-scale~~ mesoscale lee waves. Besides clear oscillations, a development towards higher  $RH_i$  and lower temperature, approaching the cloud, is discernible. From 14:36:20 UTC on,  $RH_i$  shows supersaturation and reaches values of 120 % and 130 % in the following two crests, surpassing the HET threshold.

15 Now we start a parcel at the beginning of the trajectory (measurement time: 14:40 UTC) initialized with the ECMWF temperature at this point. As it follows the trajectory, the temperature is calculated from its vertical displacement using the dry adiabatic temperature gradient. The simulated temperature and relative humidity is plotted in green. Compared to ECMWF, the temperature in the last two crests south of the cloud edge is more than 0.5 K lower and values of  $RH_i$  are higher by 10 %. The deviations result from the non-adiabatic temperature gradient in the ECMWF data.

20 These results emphasize the role of lee waves in the most southern part of the studied cirrus cloud. Comparably cooler temperatures due to adiabatic cooling in the wave crests facilitate the early nucleation of ice crystals. We find that our classification, using original ECMWF data with relatively coarse spatial resolution (horizontal grid spacing: 16 km), is able to reveal the relevant classification features within the gravity wave region. The classification quality will be even better in cases where latest ECMWF data with an improved grid spacing of 9 km, or output from regional models is available.

25 Overall our classification proved to be applicable in a meteorological context that comprises both ~~small-scale~~ mesoscale and large-scale dynamics. It identifies all relevant stages of cirrus evolution and their detailed spatial distribution and thus, also reveals the influences of gravity waves and large-scale atmospheric motion on the studied cirrus cloud.

## 5 Summary and conclusions

We presented a novel cirrus classification scheme capable of identifying all ~~evolution stages of the cirrus life-cycle~~ relevant  
30 stages of cirrus cloud evolution. It is based on airborne Lidar measurements with high spatial resolution of water vapor, backscatter and aerosol depolarization. This data is used together with ECMWF model temperature fields and knowledge and assumptions about onset conditions for homogeneous and heterogeneous freezing, to retrieve a cross-section of the cloud, revealing the detailed distribution of evolution stages.

In cloud-free air ( $BSR < 2$ ), ice supersaturated regions ( $ISSR_{out}$ ) as well as regions of homogeneous ( $HOM_{out}$ ) and heterogeneous freezing ( $HET_{out}$ ) are determined. These indicate favorable areas for cirrus cloud formation. Inside of a cloud, ice nucleation ( $HET_{in}$ ,  $HOM_{in}$ ), depositional growth (DEP) and sublimation regions (SUB) are distinguished. They represent the formation, growing and break up phase of a cirrus cloud, respectively.

5 We demonstrated the applicability of our classification in a first case study of a cirrus cloud that was observed in a complex meteorological situation comprising a thick aerosol layer, large-scale dynamics and ~~small-scale~~-mesoscale gravity lee waves. Here it revealed a non-standard horizontal order of the aforementioned ~~life-cycle~~-evolution stages and helped to identify the influence of underlying wind and gravity wave conditions as well as large-scale dynamics on individual parts of the cloud.

With this valuable tool at hand, we are investigating in our ongoing research the large airborne Lidar data set obtained during  
10 the ML-CIRRUS campaign. This classification scheme facilitates the study of the spatial distribution of evolution stages and can be used to set in-situ and other remote sensing data, obtained during the campaign, in perspective to cirrus evolution. By ~~combining~~-bringing together those data sources, the specific optical and microphysical properties of different cirrus stages now can be explored. The possibility to combine our classification with trajectory based methods promises to reveal details of the temporal succession of evolution stages. Thus we aim to achieve more detailed insights in radiative properties of cirrus clouds  
15 under various formation and life-cycle conditions.

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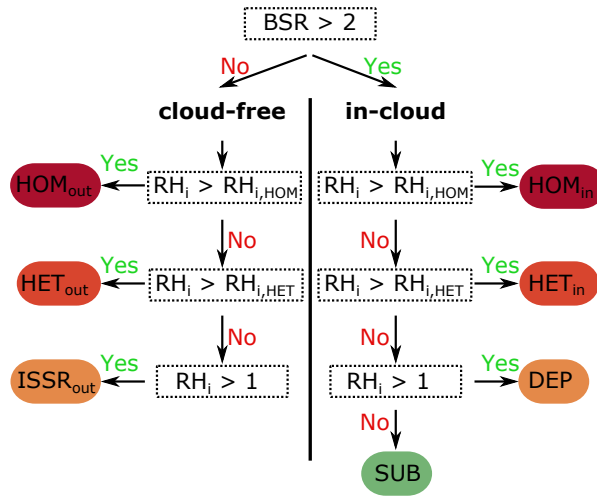
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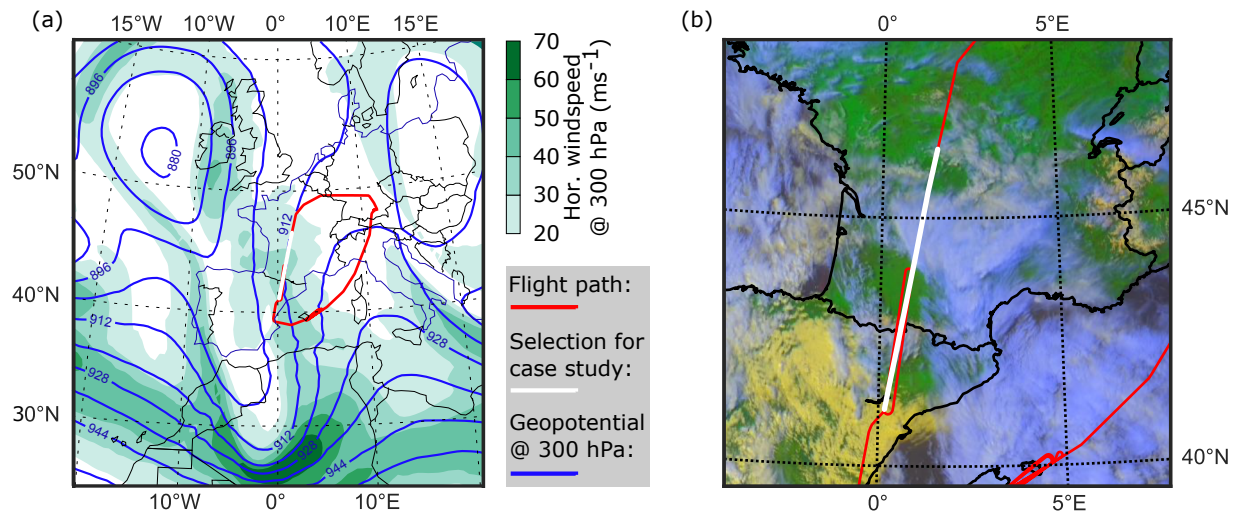
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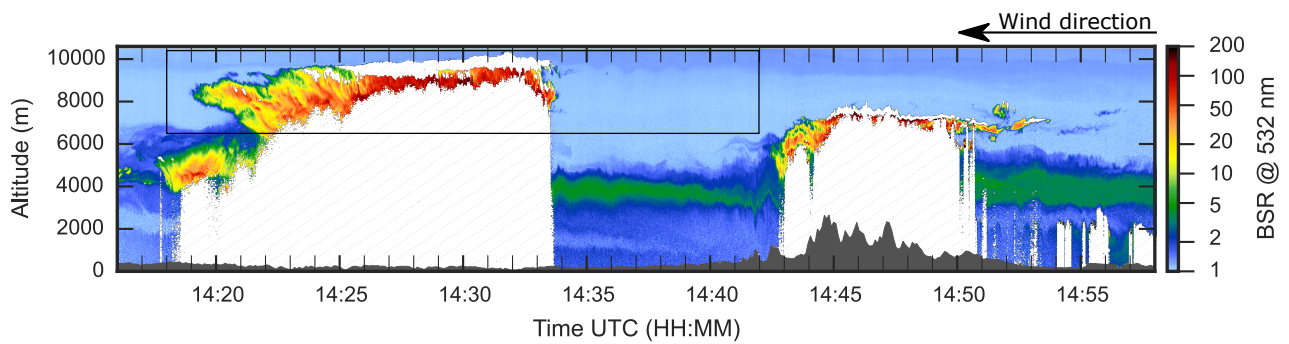




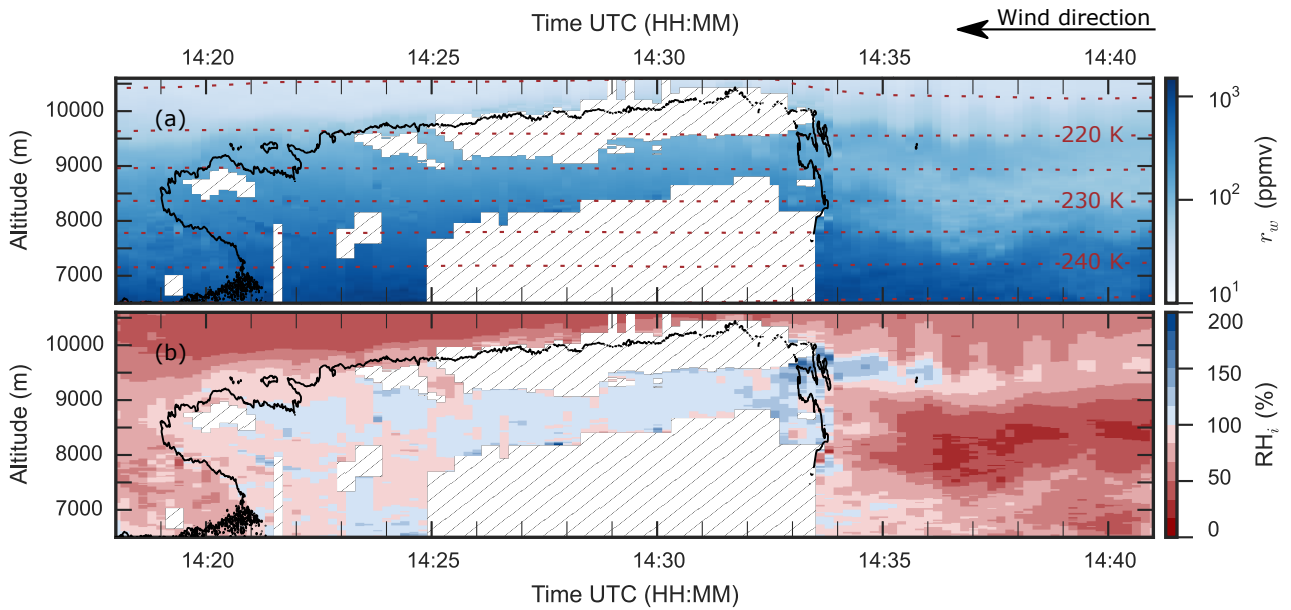
**Figure 1.** Cirrus life-cycle-evolution classification scheme based on WALES backscatter ratio (BSR) and relative humidity ( $RH_i$ ) derived from WALES humidity and ECMWF temperature field (description see text).



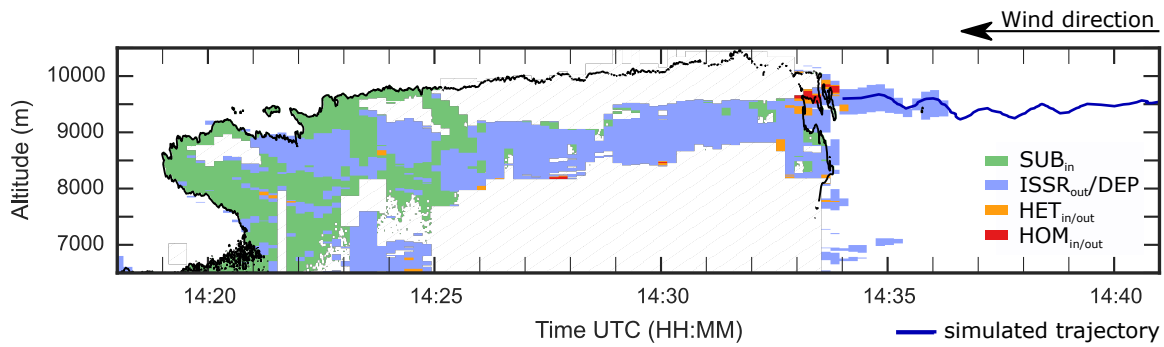
**Figure 2.** Meteorological situation on 29 March 2014 over Western Europe: (a) ECMWF analysis data (12:00 UTC) at 300 hPa of geopotential (blue isolines) and horizontal wind speed (shaded green). (b) SEVIRI false color image taken at 14:30 UTC ([see text](#)). High ice clouds have a blue color, lower liquid clouds are yellow, [cloud-free ground has green color](#). The path of the research flight is plotted in red and the flight leg used in the case study ([14:16 - 14:58 UTC](#)) is marked white.



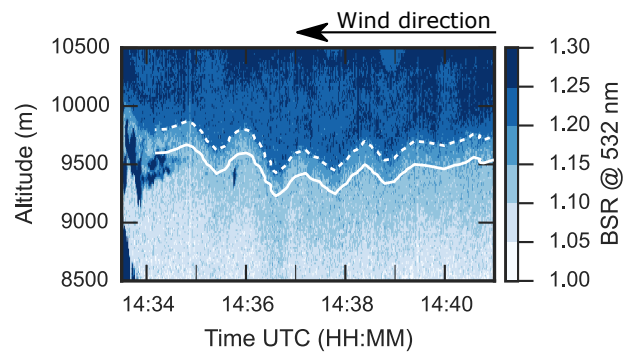
**Figure 3.** Backscatter ratio (BSR) at 532 nm measured along the flight path (white line in Fig. 2). Hatched areas indicate data that was excluded due to detector saturation or low signal to noise ratio and the terrain profile is shown in dark gray. The black rectangle marks the cirrus region that is studied further (see Fig. 4, 5). The arrow shows the main wind direction at cirrus level.



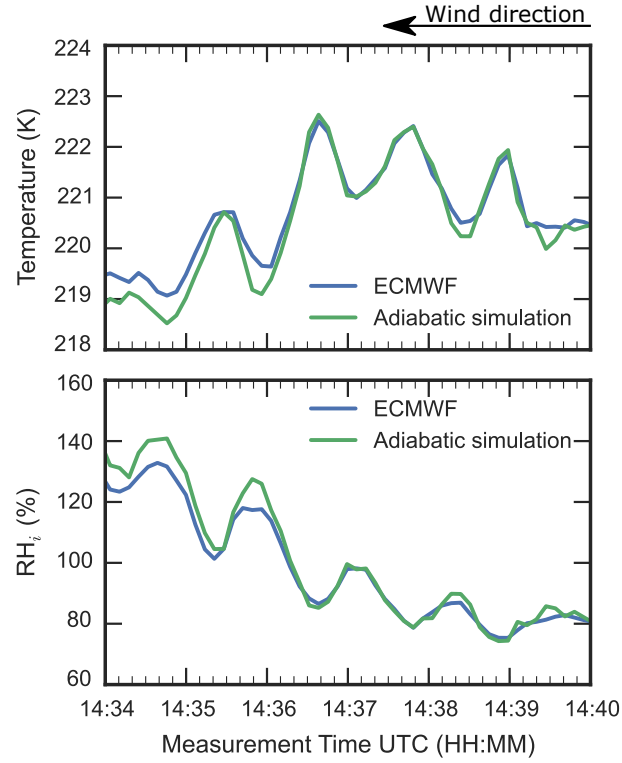
**Figure 4.** Humidity data of cirrus marked in Fig. 3: (a) Water vapor mixing ratio  $r_w$  as measured with WALES [and ECMWF temperature contour lines](#). (b) Relative humidity with respect to ice  $RH_i$  derived from WALES data and ECMWF temperature field. The cirrus is outlined by a black contour line (BSR=2) and invalid data is marked by black hatching.



**Figure 5.** Classified life-cycle evolution regions of cirrus cloud. In-cloud and cloud-free regions can be distinguished by the black contour line of  $BSR = 2$ . Regions of ice nucleation ( $HOM_{in/out}$ ,  $HET_{in/out}$ ), ice supersaturation outside ( $ISSR_{out}$ ), and depositional growth (DEP) and subsaturation (SUB) inside the cloud are visualized. Subsaturated regions outside the cloud are left blank and invalid data is marked by black hatching. The blue line shows a trajectory used for simulation of adiabatic cooling (see Sect- 4.3 and Fig. 7).



**Figure 6.** Close-up view of BSR data in the lee wave region south to the cirrus cloud. Dotted: contour line (BSR=1.2); solid: trajectory used for simulation (contour line shifted by 200 m).



**Figure 7.** Simulation of temperature  $T$  and relative humidity  $RH_i$ . Green lines are simulated values along a derived trajectory (Fig. 5) assuming adiabatic cooling and heating within the gravity wave, respectively. In blue are ECMWF values interpolated to the trajectory location.



**Table 1.** Freezing onset humidity parameterizations

$RH_{i,HOM}(T)$	$= \cancel{237 \text{ K}} - \cancel{0.4} \cdot \cancel{T} = \underline{232 - 0.37 \text{ K}^{-1} \cdot T}$	(Koop et al., 2000)
$RH_{i,HET}^{MD}(T)$	$= \cancel{134 \text{ K}} - \cancel{0.1} \cdot \cancel{T} = \underline{134 - 0.1 \text{ K}^{-1} \cdot T}$	(Krämer et al., 2016)
$RH_{i,HET}^{CS}(T)$	$= \cancel{230 \text{ K}} - \cancel{0.43} \cdot \cancel{T} = \underline{230 - 0.43 \text{ K}^{-1} \cdot T}$	(Krämer et al., 2016)